

1 **Fires, Asian, and Stratospheric Transport-Las Vegas Ozone Study**
2 **(FAST-LVOS)**

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1 **DISCLAIMER**

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3 The statements and conclusions in this Report are those of the author from the National
4 Oceanic and Atmospheric Administration, and not necessarily those of the Clark County
5 Department of Air Quality. The views, opinions, and findings contained in this report are those
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39

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20
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24
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28
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13

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1 **Abstract**

2 The Fires, Asian, and Stratospheric Transport-Las Vegas Ozone Study (FAST-LVOS) was
3 conducted in the late spring and early summer of 2017 to investigate the impact of transported
4 ozone (O₃) originating from non-controllable sources on surface concentrations in Clark County,
5 NV, including the greater Las Vegas area. This 6-week (May 17-June 30) study built on the
6 findings of the first Las Vegas Ozone Study (LVOS), which brought a more limited measurement
7 suite to Clark County in May and June of 2013. FAST-LVOS brought an improved version of the
8 TOPAZ lidar and a micro-Doppler lidar to the North Las Vegas Airport (NLVA), and deployed a
9 mobile laboratory with in-situ measurements of NO, NO_x, NO_y, CO, CO₂, CH₄, N₂O, and O₃ to
10 Angel Peak in order to investigate the interannual variability of these transport processes, and
11 more closely examine the influence of wildfires on surface O₃ in Clark County. The quasi-
12 continuous lidar and in-situ measurements were supplemented by ozonesondes and aircraft
13 sampling during 4 intensive operating periods (IOPS). In this report, we summarize these
14 measurements and assess the impact of non-controllable ozone sources (NCOS) on surface O₃
15 and NAAQS attainment during the 6-week campaign, and compare the FAST-LVOS results to the
16 earlier LVOS campaign. The measurements were augmented by a variety of models including
17 the NOAA/NESDIS RAQMS model, the NOAA/ESRL/GSD RAP-Chem and HRRR-smoke models,
18 and the FLEXPART particle dispersion model. Our analyses are also compared to the global
19 model analyses conducted by NOAA GFDL and Princeton University under a separate contract.
20 In this report, we summarize the FAST-LVOS measurements, and use these measurements
21 together with model analyses to assess the contributions of transported ozone to surface
22 concentrations in the Las Vegas Valley and greater Clark County in the late spring and summer
23 of 2017. We show that this finding is consistent with previous findings on the impact of ENSO
24 on stratosphere-to-troposphere transport in the western U.S.

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1 **Executive Summary**

2
3 The Fires, Asian, and Stratospheric Transport-Las Vegas Ozone Study (FAST-LVOS) was
4 conducted over a 6-week interval (May 17-June 30) in the late spring and early summer of 2017
5 to investigate the impact of transported ozone (O₃) on surface concentrations in Clark County,
6 Nevada. FAST-LVOS built on the findings from the 2013 Las Vegas Ozone Study (LVOS), which
7 deployed the truck-based TOPAZ (Tunable Optical Profiler for Aerosols and oZone) lidar along
8 with in-situ O₃ and CO measurements from the NOAA ESRL Chemical Sciences Division to the
9 Clark County communications facility on Angel Peak, but brought a much larger suite of
10 instrumentation to Clark County.

11
12 The primary goals of FAST-LVOS were:

- 13
14 • Assess the representativeness of the 2013 LVOS measurements and better quantify the
15 contributions of stratospheric intrusions and Asian pollution to surface ozone in Clark
16 County.
- 17
18 • Estimate the importance of these sources relative to wildfires and regional pollution.
- 19
20 • Investigate the processes by which ozone is transported downward from the free
21 troposphere to the surface.

22
23 FAST-LVOS brought a much larger suite of instrumentation to Clark County including the
24 upgraded TOPAZ lidar which was deployed to the Clark County Department of Air Quality
25 measurement site at the North Las Vegas Airport (NLVA). A continuously operating vertically-
26 staring micro-Doppler lidar was also deployed to the NLVA to measure boundary layer heights
27 and vertical winds and mixing. The TOPAZ lidar acquired data on all 45 days of the field
28 campaign, operating for a total of 523 hours or an average of nearly 12 hours per day.
29 NOAA/ESRL/CSD also deployed a mobile laboratory with nearly continuous sampling of NO,
30 NO_x, NO_y, CO, CO₂, CH₄, N₂O, and O₃ to the summit of Angel Peak.

31
32 The daily lidar and in-situ measurements were supplemented by ozonesondes launched near
33 the Clark County Joe Neal monitoring site, and airborne sampling from a single engine aircraft
34 operating out of the NLVA during 4 intensive operating periods (IOPs). Personnel from
35 NOAA/ESRL/GMD launched a total of 30 ozonesondes over 13 days (1-4 per day) and the
36 Scientific Aviation Mooney TLS Bravo logged about 90 hours over 15 days during the 4 FAST-
37 LVOS IOPs. The IOPs were conducted when synoptic conditions appeared conducive to the
38 formation of stratospheric intrusions above the Western U.S.

39
40 Planning for the IOPs relied on forecasts from the NCEP Global Forecast System (GFS)) and the
41 NOAA/ESRL/GSD RAP-Chem and NOAA/NESDIS RAQMS models, and interpretation of the
42 measurements was aided by the FLEXPART particle dispersion model and the NOAA GFDL AM4
43 model. The AM4 modelling was directly supported by Clark County under a separate project
44 (CBE605334-19) and is described in more detail within a separate report.

1
2 The FAST-LVOS measurements found high levels of O₃ between 4 and 6 km in the free
3 troposphere above Clark County on nearly all of the 45 measurement days. The measurements
4 and model analyses show that the O₃ aloft originated primarily from stratospheric intrusions,
5 transported Asian pollution, and wildfire plumes. Entrainment of free tropospheric O₃ layers
6 into the mixed layer was directly observed on multiple occasions, confirming the hypothesis
7 inferred from the LVOS 2013 measurements.

8
9 The 2015 NAAQS of 70 ppbv was equaled or exceeded by one or more of the Clark County
10 Department of Air Quality monitors on 10 of the 45 FAST-LVOS measurement: May 23(1), June
11 3, (1), June 14(2), June 16(3) and 17(3), June 22(4) and 23(3), and June 28(2), 29(1), and 30(3),
12 where the numbers in parentheses show the number of affected monitors. Our analyses
13 suggest that transported pollution played a role in at least half of these exceedances. We found
14 that Asian pollution may have contributed to the exceedance on June 3, and the major
15 stratospheric intrusion that passed over Clark County on June 11-13 contributed to the total
16 background O₃ concentrations of 45-50 ppbv that comprised the bulk of the 62-73 ppbv
17 measured in the Las Vegas Valley on June 14. Our analyses also show that wildfires played a
18 major role in the exceedances of June 22, 23, and 28, and may also have contributed to the high
19 ozone on June 29, and 30.

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1 **Definitions and Abbreviations**

2

3	agl	above ground level
4	asl	above mean sea level
5	AS	Asian
6	BB	biomass burning
7	CC	Clark County
8	CIMSS	Cooperative Institute for Meteorological Satellite Studies
9	CIRES	Cooperative Institute for Research in Environmental Sciences
10	CSD	Chemical Sciences Division
11	DAQ	Department of Air Quality
12	ESRL	Earth System Research Laboratory
13	FAST-LVOS	Fires, Asian, and Stratospheric Transport- Las Vegas Ozone Study
14	GFDL	Geophysical Fluid Dynamics Laboratory
15	GFS	Global Forecasting System
16	GMD	Global Monitoring Division
17	GSD	Global Systems Division
18	HRRR	High-resolution rapid refresh
19	IOP	Intensive operating period
20	LVOS	Las Vegas Ozone Study
21	LVV	Las Vegas Valley
22	MDA8	daily maximum 8-h average
23	NARR	North American Reanalysis
24	NCEP	National Centers for Environmental Prediction
25	NCOS	non-controllable ozone source
26	NESDIS	National Environmental Satellite, Data, and Information Service
27	NOAA	National Oceanic and Atmospheric Administration
28	NPS	National Park Service
29	RAP-Chem	Rapid refresh with chemistry
30	RAQMS	Realtime Air Quality Modeling System
31	NAAQS	National Ambient Air Quality Standard
32	ppbv	parts-per-billion-by-volume
33	SMYC	Spring Mountain Youth Camp
34	SoCAB	South Coast Air Basin
35	ST	stratospheric
36	STT	stratosphere-to-troposphere transport
37	TOPAZ	Tunable Optical Profiler for Aerosols and oZone
38	UT/LS	upper troposphere/lower stratosphere
39	WRCC	Western Regional Climate Center
40		

1 1. Introduction

2
3 High concentrations of ozone (O₃) are harmful to human health [*U.S. Environmental Protection Agency, 2014*] and impair plant growth and productivity [*Avnery et al., 2011a; b*]. The U.S. Clean Air Act (CAA) accordingly designated ground-level ozone a criteria pollutant and established
4 primary and secondary National Ambient Air Quality Standards (NAAQS) to protect human
5 health and welfare, respectively. Compliance with the CAA requires that the design value (DV),
6 i.e. the 3-yr running average of the 4th highest maximum daily 8-h average mixing ratio (MDA8)
7 O₃ concentration measured annually within an air basin, to be less than or equal to the NAAQS.
8 The CAA also requires that the NAAQS be reviewed periodically to consider the latest research
9 on the health effects, and adjusted, if necessary, to ensure an adequate margin of safety for the
10 public. The primary and secondary NAAQS for ozone were lowered from 75 to 70 parts-per-
11 billion by volume (ppbv) for the in October 2015 following the most recent such review [*U.S.*
12 *Environmental Protection Agency, 2015*].
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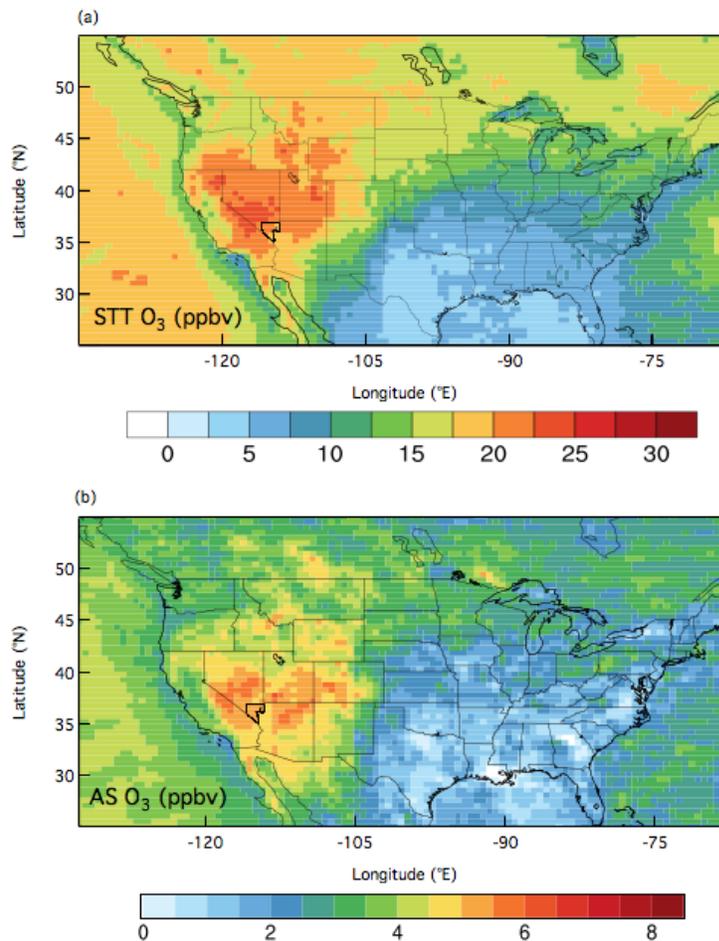
15
16 Ozone is a secondary pollutant formed by photochemical reactions between nitrogen oxides
17 (NO_x) and volatile organic compounds (VOCs) emitted by both natural and anthropogenic
18 sources. Surface O₃ has declined across much of the U.S. over the last two decades as more
19 stringent emission controls have led to substantial reductions in anthropogenic NO_x precursors
20 [*Gaudel et al., 2018*]. The mean DV for the contiguous U.S. decreased by 17% from 82 to 67
21 ppbv between 2000 and 2017 (<https://www.epa.gov/air-trends/ozone-trends#ozreg>). The
22 largest declines occurred in the Southeastern (83 to 62 ppbv or 26% at 145 sites) and South-
23 Central (87 to 66 ppbv or 24% at 104 sites) U.S., with much smaller changes in the West (82 to
24 75 ppbv or 8% at 143 sites) and Southwest (77 to 72 ppbv or 7% at 58 sites)¹. The 4th highest
25 MDA8 O₃ concentrations measured in the West (California and Nevada) and Southwest
26 (Arizona, New Mexico, Utah and Colorado) during 2017 were higher than those measured in
27 any other region of the U.S.
28

29 The weaker response of ground-level O₃ to declining NO_x emissions in the western U.S. can be
30 explained, at least in part, by the presence of much higher background concentrations [*Jaffe et*
31 *al., 2018*]. Ozone is the only one of the six criteria pollutants (CO, NO₂, SO₂, O₃, Pb, and
32 particulate matter) with a significant background in the troposphere. This background is
33 created primarily by downward transport of O₃ from the stratosphere or production of O₃ by
34 reactions of NO_x and VOCs from non-anthropogenic sources such as lightning, wildfires, soils,
35 and vegetation. Since O₃ is relatively long-lived in the free troposphere (≈1 month), the
36 concentrations measured at any given location may include O₃ formed from natural or
37 anthropogenic precursors emitted hundreds or thousands of kilometers upwind and even a
38 small amount of O₃ formed from local or regional precursors that has circumnavigated the
39 globe.
40

41 The high western background can be explained by several factors. First, the Southwestern U.S is
42 a global hotspot for deep tropopause folds [*James et al., 2003; Langford et al., 2012; Skerlak et*

¹ The 2000-2017 decline in Las Vegas (78 to 71 ppbv, or about 9%) was similar to the rest of the Southwest

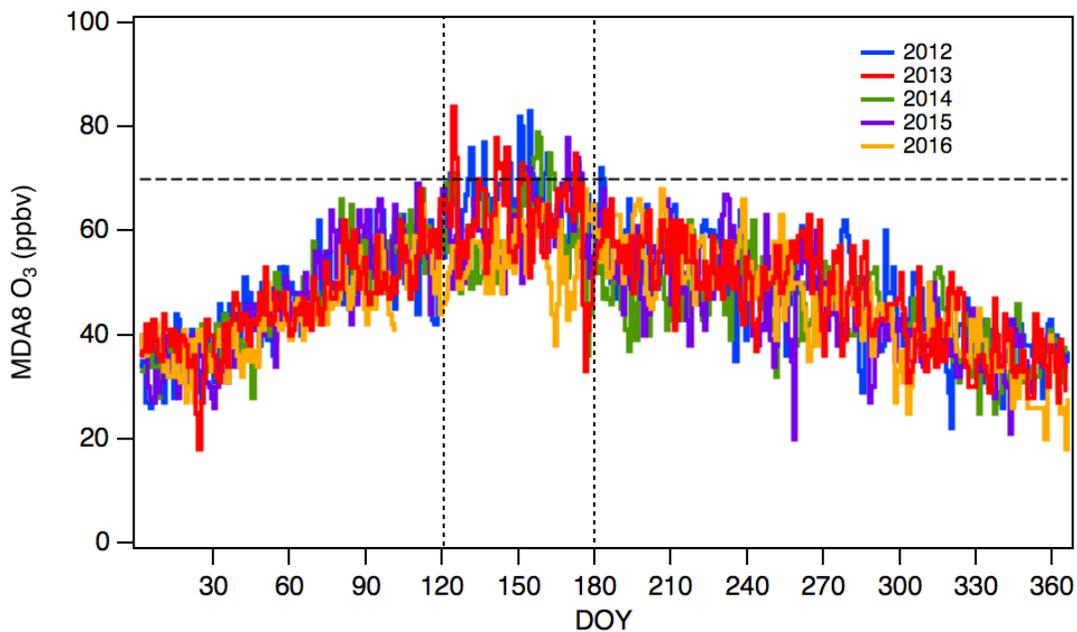
1 *al.*, 2014; Sprenger and Wernli, 2003; Wernli and Bourqui, 2002], intrusions of lower
2 stratospheric air that penetrate to within 3 km of the ground. These folds are created by jet
3 streaks in the winds circulating around mid-latitude cyclones and stratosphere-to-troposphere
4 transport (STT) peaks in late spring when cyclonic activity is still high and the O₃ content of the
5 lower stratosphere at a maximum. Second, many intrusions entrain wildfire or pollution plumes
6 transported across the Pacific Ocean from Asia as they descend through the troposphere and
7 bring this O₃-enriched air toward the surface with them. Finally, the high average elevation of
8 the Southwest and Intermountain West increase the likelihood that descending O₃-rich air will
9 reach the surface. Global model simulations (**Figure 1-1**) show the largest influxes of
10 stratospheric O₃ and Asian pollution are in the Intermountain West including Clark County, NV.
11



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13
14 **Figure 1-1.** Mean contributions (in ppbv) of (a) STT, and (b) Asian pollution, to MDA8 surface O₃
15 calculated by the NOAA GFDL AM3 model for May and June 2010. The resolution is 50 km x 50
16 km. Note the different color scales. Clark County, NV is outlined in black. Adapted from [Lin et
17 *al.*, 2012a; Lin et *al.*, 2012b].

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1 Photochemical production of O₃ depends primarily on sunlight and high temperatures, and
2 measurements in the LA Basin accordingly show a broad O₃ peak in midsummer. Some of this
3 O₃ must be transported inland by the prevailing southwesterly winds along the West Coast and
4 it might be expected that measurements in rural areas of the Desert Southwest would also
5 show a midsummer peak, particularly since long-range transport of O₃ and other pollutants
6 from the LA Basin in the lower free troposphere begins with thermally driven orographic lifting
7 by the mountain ranges surrounding the Basin. Surprisingly, however, five years (2012-2016) of
8 measurements from the southernmost Clark County monitor at Jean, which lies between Los
9 Angeles and Las Vegas (cf. **Figure 2-1b**) show that the highest concentrations are typically
10 measured earlier during May and June (**Figure 2-2**). Similar patterns are found in the
11 measurements at many of the remote USNPS sites surrounding Clark County (cf. **Figure 2-1**).
12



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14
15 **Figure 2-2.** MDA8 O₃ measured by the Clark County DAQ monitor at Jean in southwestern Clark
16 County for the years 2012-2016. The horizontal dashed line marks the 70 ppbv NAAQS. May and
17 June lie between the two vertical dotted lines.

18
19 The late spring peak in **Figure 2-2** coincides with the time of year when transport of O₃ from
20 Asia and the lower stratosphere is greatest. The apparent connection between STT and high
21 springtime O₃ in Clark County motivated Dr. Zheng Li of the Clark County Department of Air
22 Quality to propose the Las Vegas Ozone Study (LVOS), which was subsequently conducted in
23 the late spring and early summer of 2013 [Langford et al., 2015b]. The NOAA/ESRL/CSD
24 deployed the Tunable Optical Profiler for Aerosol and oZone (TOPAZ) mobile lidar and in-situ
25 CO and O₃ measurements to the Clark County communications facility on the summit of Angel
26 Peak (2.7 km asl) about 45 km northwest of Las Vegas in the Spring Mountains [Langford et al.,
27 2015b].

1 The study lasted for a period of 6 weeks (May 20 to June 28) during which TOPAZ detected
2 numerous layers of high O₃ between 4 and 6 km asl consistent with intrusions of lower
3 stratospheric air with or without admixed Asian pollution [Langford et al., 2018], and the direct
4 descent of one of these intrusions to the summit of Angel Peak coincided with a simultaneous
5 increase in surface O₃ and decrease in CO and H₂O. This interpretation was supported by
6 simulations from the NOAA GFDL AM3 global model and the FLEXPART particle dispersion
7 model.

8 On three occasions, the appearance of high O₃ layers above Angel Peak during LVOS were
9 followed by exceedances of the 2008 (75 ppbv) NAAQS in the Las Vegas Valley the next day,
10 suggesting [Langford et al., 2017] that the elevated layers were entrained and mixed to the
11 surface in the LVV by the exceptionally deep convective mixed layers that form above the
12 Mojave Desert [Seidel et al., 2012].

13 14 **3. The Fires, Asian, and Stratospheric Transport-Las Vegas Ozone Study (FAST-LVOS)**

15
16 The FAST-LVOS measurement campaign was designed primarily to assess the
17 representativeness of the 2013 LVOS findings and test the entrainment hypothesis [Langford et
18 al., 2017]. Another important goal was to better quantify how the contributions of STT and
19 Asian pollution compare to other sources including local photochemical production, wildfires,
20 and regional transport from California. A third goal was to examine the interannual variability of
21 deep STT in the U.S. Southwest, which is influenced by ENSO [Lin et al., 2015] and other large-
22 scale circulations [Albers et al., 2018].

23
24 FAST-LVOS brought a greatly expanded suite of measurements to Clark County, including a core
25 set of lidar and in-situ instruments based at the North Las Vegas Airport (NLVA) and Angel Peak
26 (AP), respectively, that operated throughout the campaign (May 17 to June 30, 2017). These
27 core measurements complemented the routine measurements from the Clark County DAQ
28 monitoring network, and were supplemented by ozonesonde and aircraft sampling during four
29 2-to-4 day long intensive operating periods (IOPs) called when the synoptic conditions
30 appeared favorable to the formation of tropopause folds above the western U.S.

31
32 Daily measurements of ozone and backscatter profiles were made by TOPAZ and a vertically-
33 staring Doppler lidar, which also measured vertical wind speed and variance. These lidars were
34 co-located at the NLVA with the Clark County DAQ profiling radiometer and wind profiler.
35 Nearly continuous in-situ measurements of NO, NO_x, NO_y, CO, CO₂, CH₄, N₂O, and O₃ were
36 made by mobile sampling laboratory that primarily remained on Angel Peak.

37 These measurements from NOAA/ESRL/CSD were augmented by ozonesondes from
38 NOAA/ESRL/GMD launched ≈7.3 km NW of TOPAZ at the Estelle Neal park adjacent to the
39 CCDAQ Joe Neal monitoring site, and by aircraft sorties between NLVA and Big Bear, CA by the
40 Scientific Aviation Mooney single engine research aircraft during the IOPs.

41

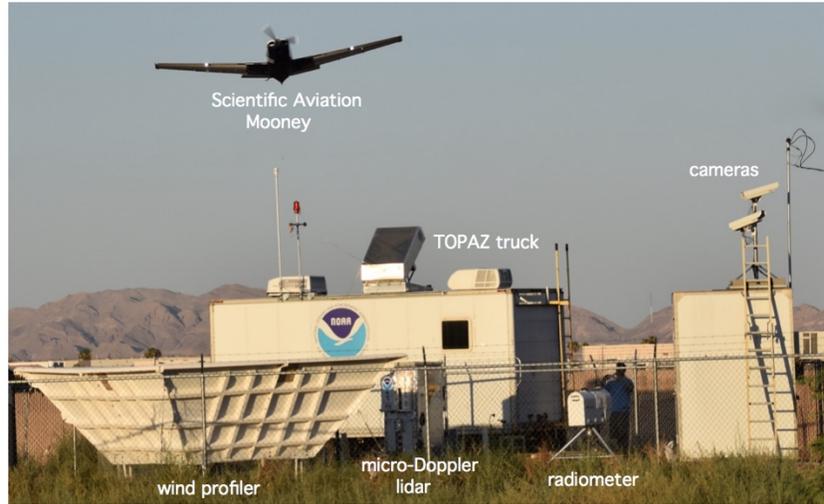
1 The FAST-LVOS study relied on several atmospheric models and tools for forecasting and
2 planning during the measurement campaign, and for interpretation of the measurements
3 afterwards. The former group included both operational (e.g. NCEP Global Forecast System
4 (GFS)) and “quasi-operational” models (e.g. RAP-Chem () and RAQMS ()). The GFS forecasts and
5 NWS radiosonde profiles were obtained from the Wyoming Weather Web maintained by the
6 University of Wyoming (<http://weather.uwyo.edu/index.shtml>). The NOAA HRRR-smoke model
7 () was also used to assess wildfire influences. We also made use of NOAA satellite products
8 archived by the Cooperative Institute for Research in the Atmosphere (CIRA) at the Colorado
9 State University (<http://rammb.cira.colostate.edu/ramsd/online/index.asp>) for forecasting
10 purposes. The post-campaign analysis relied heavily on the FLEXPART particle dispersion model
11 and the NOAA GFDL AM4 model. The NOAA GFDL AM4 modeling efforts were directly
12 supported by Clark County (CBE605334-19) and are described in detail in a separate report.
13

14 **4. FAST-LVOS Measurement Suite**

15

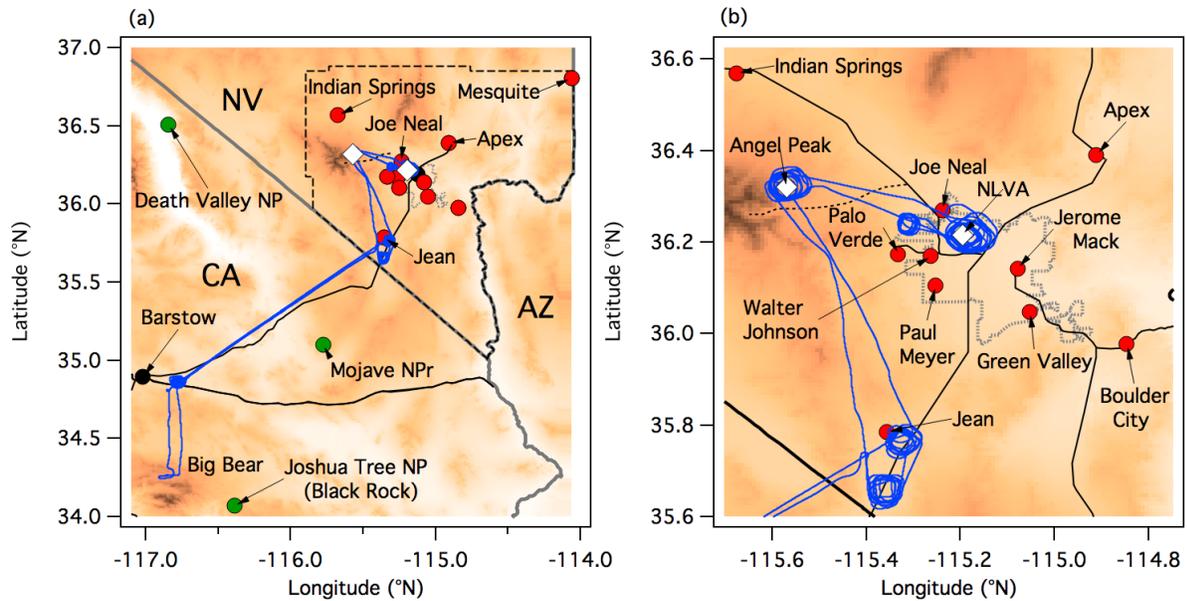
16 The FAST-LVOS field campaign was based primarily at the North Las Vegas Airport (36.2°N, -
17 115.2°E, 681 m asl), which hosted the NOAA TOPAZ and micro-Doppler lidars (**Figure 4-1**).
18 These instruments were deployed within the Clark County DAQ enclosure near the radar wind
19 profiler and profiling radiometer comprising the CCDAQ Integrated Upper-Air Station
20 (http://airquality.clarkcountynv.gov/cgi-bin/DAQ/RAOB_charts.pl). The Scientific Aviation
21 Mooney TLS Bravo aircraft operated out of the NLVA during the FAST-LVOS IOPs and the
22 NOAA/ESRL/GMD Ozonesondes were launched adjacent to the nearby (7 km) Joe Neal
23 monitoring station. The NOAA mobile laboratory spent most of the campaign on the summit of
24 Angel Peak (Appendix C) where TOPAZ was located during the first LVOS campaign. **Figure 4-2**
25 shows these sites along with the regulatory O₃ monitors operated by Clark County and other
26 agencies, and the June 28 flight track of the Mooney.
27

28 The daily lidar and in-situ measurements at NLVA and Angel Peak were supplemented by
29 ozonesonde and aircraft measurements during four IOPs (May 23-25, May 31-June 2, June 10-
30 14, and June 28-30). These intensives were conducted during periods when the synoptic
31 conditions appeared favorable for the development of tropopause folds above the western U.S.
32 Planning for the IOPs relied primarily on the NCEP Global Forecast System (GFS) long range
33 weather forecasts obtained from the University of Wyoming
34 (<http://weather.uwyo.edu/models/fcst/gfs003.shtml>) and NOAA RAP-Chem
35 (<https://rapidrefresh.noaa.gov/RAPchem/>) and RAQMS (<http://raqms-ops.ssec.wisc.edu/>)
36 models which were reviewed daily by the FAST-LVOS Principal Investigator. The ozonesonde
37 and aircraft teams were typically alerted 48 to 72 hours before the start of the IOP.
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Figure 4-1. The Scientific Aviation Mooney makes a low overpass of the NLVA measurement site. The NOAA TOPAZ lidar truck and micro-Doppler lidar can be seen in the background behind the Clark County DAQ radar wind profiler, profiling radiometer, and visibility cameras. The 2B ozone monitor sampling mast is seen to the left of the red warning light on top of the TOPAZ truck. Sunrise Mountain can be seen just above the wind profiler (photo by A.O. Langford)



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Figure 4-2. Topographical maps showing the FAST-LVOS measurement domain. The filled red, green, and black circles locate the regulatory O₃ monitors operated by Clark County, the U.S. National Park Service, and the State of California, respectively. The white diamonds show the FAST-LVOS measurement sites at NLVA and AP. The blue trace shows the Scientific Aviation flight track from June 28, 2017.

4.1 NOAA TOPAZ ozone and aerosol lidar

The truck-mounted NOAA/ESRL/CSD TOPAZ mobile differential absorption lidar (DIAL) system was located at the North Las Vegas Airport (NLVA) for the duration of the FAST-LVOS campaign. TOPAZ was originally developed for profiling O₃ and particulate backscatter in the boundary layer and lower free troposphere from NOAA Twin Otter aircraft [Alvarez *et al.*, 2011; Langford *et al.*, 2012; Langford *et al.*, 2010; Langford *et al.*, 2011; Senff *et al.*, 2010]. The lidar was reconfigured and installed in a medium-duty 4-wheel drive truck after CalNex, and subsequently deployed in this configuration to several field campaigns including the 2013 Las Vegas Ozone Study (LVOS) [Langford *et al.*, 2015b].

Improvements in the TOPAZ data acquisition system during 2015 greatly extended the operational range of TOPAZ from the 2.5 to 3 km available during LVOS, to 6-8 km during the 2016 CABOTS [Faloona *et al.*, 2019] and 2017 FAST-LVOS campaigns. This improvement allowed the deployment of the lidar at the lower elevation NLVA site without any loss in altitude coverage compared to LVOS.

The TOPAZ lidar uses a low energy ($\approx 50 \mu\text{J}$), high pulse rate (1 kHz) quadrupled Nd:YLF pumped Ce:LiCAF laser that is re-tuned between each pulse to generate light at three different wavelengths ranging from 287 to 294 nm reducing the effective repetition rate to 333 Hz [Alvarez *et al.*, 2011]. The low pulse energies and tunability of TOPAZ allow the system to be operated from airports and other locations where eye safety is a major concern and facilitate optimization of the wavelengths for lower tropospheric measurements. The laser pulses are transmitted and backscatter signals collected by a zenith viewing coaxial transmitter/receiver beneath an opening in the truck roof. A large scanning mirror inserted above the transceiver allows profile measurements at different slant angles. These slant profiles can be combined to create vertical profiles that begin much closer to the ground than conventional vertically staring lidar systems [Proffitt and Langford, 1997]. During FAST-LVOS, the scanning mirror was moved sequentially between elevation angles of 90, 20, 6, and 2° with a 225-s averaging time at 90° and 75-s averaging times at the other 3 angles. The cycle was repeated approximately every 8 minutes and the vertical projections combined to create a composite vertical profile beginning 27.5 ± 5 m above ground level (agl).

In addition to the lidar, the TOPAZ truck is equipped with an *in-situ* O₃ monitor (2B Technologies Model 205) that samples air 5 m above the surface and an Airmar 150WX weather station to measure temperature, pressure, relative humidity, and wind speed and direction.

The O₃ and aerosol profiles were computed using the iterative technique described in [Alvarez *et al.*, 2011]. The O₃ profiles were retrieved using the longest and shortest wavelengths (≈ 287 and 294 nm) with 30 m range gates and a smoothing filter with a varying width that increased from 270 m at the minimum range (800 m) up to 1400 m at the maximum range (8 km). The aerosol backscatter and extinction wavelength dependences, which are needed to correct differential aerosol effects in the O₃ calculation, are assumed to follow a power law. In most cases, a power-law exponent of 0 (no wavelength dependence) was used for the aerosol

1 backscatter with an exponent of -0.5 for aerosol extinction. These values seem to be a good
2 compromise for a wide range of aerosol types [Völger *et al.*, 1996]. In the second iteration step,
3 O₃ profiles are computed with an aerosol correction based on the aerosol backscatter and
4 extinction profiles from the first step. These O₃ profiles are then used to provide a more
5 accurate O₃ extinction correction of the signal data at 294 nm, which in turn results in more
6 accurate aerosol profile retrieval. This iteration procedure is repeated until the O₃ profiles
7 produced in successive iteration steps converge. Convergence is reached when the absolute
8 difference between successive O₃ profiles is less than $2.5 \times 10^{15} \text{ m}^{-3}$ (corresponding to about
9 0.1 ppbv) at all range gates.

10
11 The DIAL calculations used the temperature dependent O₃ absorption cross-sections from
12 Malicet *et al.* [Malicet *et al.*, 1995]. The National Centers for Environmental Prediction (NCEP)
13 North American Regional Reanalysis (NARR) temperature and pressure profiles from the grid
14 point closest to the TOPAZ lidar location were used to account for the temperature dependence
15 of the O₃ cross sections and to convert O₃ number densities to mixing ratios. The 3-hourly NARR
16 temperature and pressure profiles were interpolated to the exact time each individual O₃
17 profile was recorded. The total uncertainties in the 8-min ozone retrievals are estimated to
18 increase from ± 3 ppbv below 4 km, to ± 10 ppbv at the top of the profile. Profiles of the
19 backscatter from aerosols, smoke, and dust were retrieved with 7.5 m resolution at 294 nm.
20 These unpolarized single wavelength backscatter profiles provide a semiquantitative measure
21 of the particulate distribution, but do not distinguish between different particle types.

22
23 An intercomparison at the 2014 Colorado DISCOVER-AQ campaign [Wang *et al.*, 2017] found
24 that the performance of the original analog data acquisition system developed at NOAA for
25 aircraft operations in 2004 had deteriorated. This system was accordingly replaced with a new
26 commercial dual analog/photon counting system prior to CABOTS. This upgrade improved the
27 stability of the lidar retrievals and greatly expanded the useful altitude range of TOPAZ
28 compared to earlier studies. The maximum altitude achieved during CABOTS ranged from
29 ≈ 6 km during the day to more than 8 km at night depending on the lidar return signal-to-noise
30 ratios, which varied with laser power, atmospheric extinction, and solar background light.
31 The improved performance was verified during the California Baseline Ozone Transport Study
32 (CABOTS) conducted in the summer of 2016 [Langford *et al.*, 2019] and the TOLNet multi-lidar
33 Southern California Ozone Observation Project (SCOOP) intercomparison at the NASA Jet
34 Propulsion Laboratory (JPL) Table Mountain Facility (TMF) in the San Gabriel Mountains
35 immediately after CABOTS [Leblanc *et al.*, 2018].

36 37 **4.2 NOAA micro-Doppler lidar**

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39 The autonomous micro-Doppler lidar (μ DL) used during FAST-LVOS was developed by the
40 NOAA/ESRL Atmospheric Remote Sensing Group in 2016 and field tested in Visalia, CA during
41 the CABOTS campaign [Faloona *et al.*, 2019]. This near-infrared eye safe coherent lidar
42 determines wind speeds from phase shifts in the light backscattered from moving aerosol
43 particles [Grund *et al.*, 2000]. The μ DL was installed adjacent to the CCDAQ radar wind profiler

1 during FAST-LVOS (**Figure 4-3**) and operated in a vertically-staring mode to continuously
2 measure vertical wind velocities above the NLVA. The variance in the vertical velocities was
3 then used in conjunction with the aerosol backscatter to infer boundary layer heights and
4 mixing [*Bonin et al., 2017; Tucker et al., 2009*].

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Figure 4-3. The NOAA autonomous micro-Doppler lidar stands on the former RASS pad adjacent to the TOPAZ lidar truck. (photo by A.O. Langford)

4.3 NOAA mobile sampling laboratory

The Chemical Sciences Division also brought a van-based mobile laboratory (https://esrl.noaa.gov/csd/groups/csd7/measurements/csd_mobilelab/) for in-situ measurements to FAST-LVOS [Wild *et al.*, 2017] (**Figure 4-4**). The van was outfitted with instrumentation to measure O₃, CO, CO₂, CH₄, NO, NO₂, NO_y, N₂O, H₂O, and meteorological parameters during FAST-LVOS (**Table 4-1**). All of the measurements can be made while moving with 1-s sampling, but the van can also be connected to external power for stationary operations. The van stayed on Angel Peak (**Figure 4-5**) during most of the study, but also executed a series of Angel Peak-Lee Canyon-Kyle Canyon drives on May 24-25, and relocated to the NLVA for an intercomparison with the SA Mooney on June 15 (**Figure 4-6**).



Figure 4-4. Jeff Peischl of NOAA/ESRL/CSD and CIRES works on the mobile lab instrumentation at Angel Peak (photo by A.O. Langford).

The in-situ measurements of O₃, CO, and H₂O made at Angel Peak during LVOS provided a simple matrix to help identify air mass origins. Ozone derives from both tropospheric and stratospheric sources, but CO and H₂O originate at the earth's surface and have greatly decreased concentrations in the stratosphere. Thus, air descending from the upper troposphere/lower stratosphere (UT/LS) leads to enhancements in surface O₃ that are coupled with sharp decreases in CO and H₂O. Asian pollution plumes transported across the Pacific Ocean in the middle troposphere are nearly as dry, but typically contain elevated levels of CO. Local photochemistry and regional biomass burning plumes typically have both high CO and N₂O. Marine boundary layer tends to have low O₃ and CO, but high H₂O. The FAST-LVOS mobile

1 lab measurements provide a more complete matrix to identify different air sources.
 2 Stratospheric air can be identified by enhancements in O₃ correlated with decreased CO, N₂O,
 3 and H₂O since these species originate at the Earth's surface and have much lower
 4 concentrations in the stratosphere. The NO, NO₂, (NO_x=NO+NO₂) and NO_y measurements can
 5 be used along with CO and CO₂ to identify biomass burning and other combustion sources.
 6 Methane (CH₄) provides a marker for oil and gas activities. Summary plots showing the mobile
 7 laboratory measurements are plotted in Appendix C.

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10 **Table 4-1. Mobile laboratory payload during FAST-LVOS**

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Measured Parameter	Method	Time Resolution	Detection Limit
CO ₂ and CH ₄	Wavelength scanned cavity ring-down spectroscopy	1 sec	0.2 ppmv for CO ₂ 2 ppbv for CH ₄
N ₂ O, CO, H ₂ O	Integrated cavity output spectroscopy	1 sec.	0.2 ppb for N ₂ O 0.2 ppb for CO 100 ppm for H ₂ O
NO, NO ₂ , NO _y , O ₃	Cavity ring-down spectroscopy	1 sec	0.1 - 0.001 ppbv
O ₃	2B Technologies, Model 205	10 sec	≈1 ppbv
Position and Met Data	Differential GPS, Airmar Wx and Young 2-D sonic	1 sec	

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Ozone was measured both by cavity ringdown spectroscopy (CRDS) and by a conventional UV absorption analyzer (2B Model 205).



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Figure 4-5. Aerial view of Angel Peak from the SA Mooney showing the NOAA mobile laboratory. The TOPAZ mobile lidar was parked in the same spot during LVOS, and the LVOS in-situ measurements made from the roof of the Clark County building. (photo by D. Caputi).



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Figure 4-6. NOAA mobile laboratory and Scientific Aviation aircraft during standards comparison at the NLVA. (photo by R. Langston)

1 **4.4 NOAA ECC ozonesondes**

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3 NOAA/ESRL GMD launched up to 4 electrochemical cell (ECC) [Johnson et al., 2002].
4 ozonesondes per day (30 total) on 13 days during the 4 FAST-LVOS IOPs. The balloons were
5 launched from the Estelle Neal Park (adjacent to the CCDAQ Joe Neal monitoring site) during
6 the four FAST-LVOS IOPs. These balloon-borne packages combined the ozone sensor with a
7 GPS-equipped meteorological radiosonde and transmit ozone partial pressures, pressure,
8 temperature, relative humidity, and winds to a ground station throughout the flight. The
9 balloons typically ascended to more than 20 km and acquired both tropospheric and
10 stratospheric O₃ profiles. In some cases, useful data were also obtained during the parachute
11 descent, although these profiles were often displaced from the launch site by 10-50 km.
12 Although the ozonesondes lack the temporal resolution of the lidar, they can profile O₃ from
13 the surface to about 30 km, along with temperature, frost point, relative humidity, and winds.
14 These measurements provide important information about the atmospheric stability, and the
15 frost point measurements can help distinguish upper tropospheric from lower stratospheric air.
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21 **Figure 4-7.** Patrick Cullis (L) and Chance Sterling (R) of NOAA/ESRL/GMD and CIRES, U. of
22 Colorado prepare to launch one of the FAST-LVOS ozonesondes from the Estelle Neal municipal
23 park (photo by R. Langston).

1
 2 **Table 4-2** lists the ozonesonde launch times. Multiple launches were conducted at 3-hour
 3 intervals: 0900, 1200, 1500, 1800 PDT. Four ozonesondes were launched on June 11 and 12
 4 during the major stratospheric intrusion event, and one each during the last three days of the
 5 study. All of the ozone profiles are shown in Appendix D.

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 10 **Table 4-2. NOAA/ESRL/GMD ozonesonde launch times**
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Date	Launch 1 (PDT)	Launch 2 (PDT)	Launch 3 (PDT)	Launch 4 (PDT)
May 23	9:05	12:03		
May 24	9:03	12:01	15:12	
May 25	11:57			
May 31	9:18	12:05		
June 1	9:18	11:56	15:03	
June 2	9:14	12:09	15:01	
June 10	9:08	11:57		
June 11	8:54	11:54	14:53	17:54
June 12	9:02	11:58	14:58	17:56
June 13	8:57	12:02	15:00	
June 28	11:59			
June 29	12:03			
June 30	11:59			

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4.5 Scientific Aviation Mooney aircraft

Scientific Aviation, Inc. (<http://www.scientificaviation.com>) conducted research flights during the FAST-LVOS intensives using a single-engine Mooney TLS Bravo aircraft (**Figure 4.8**). The aircraft carried a pilot and technician, along with a 2B Technologies Model 205 O₃ monitor, an Aerodyne Research Cavity Attenuated Phase Shift (CAPS) NO₂ monitor, and a Picarro 2301f Wavelength-scanned Cavity Ring-Down Spectrometer (WS-CRDS) to measure CO₂, CH₄, C₂H₆, and H₂O [Trousdel et al., 2016]. The 2B O₃ data were sampled every 2 s, which corresponds to a mean distance of 150 m at the typical level leg flight speed of 75 m s⁻¹. The airborne 2B O₃ measurements were compared to TOPAZ lidar profiles during the California Baseline Ozone Transport Study (CABOTS) in the summer of 2016 [Langford et al., 2019].



Figure 4-8. Scientific Aviation 1998 Mooney TLS Bravo M20M (N2132X) on landing approach at NLVA (photo by A. Langford).

The typical research flight (cf. **Figure 4-2**) included a spiral climb to ≈ 6 km after takeoff from the NLVA followed by transit to Angel Peak and a spiral descent from there into the western valley. The aircraft then flew to Jean and conducted another spiral profile near the CCDAQ monitor. The aircraft then transited to the airport at Big Bear Lake, CA at various altitudes, where it landed and re-fueled. The aircraft conducted a spiral near Barstow after taking off from Big Bear and reversed the outbound flight plan on the return leg to NLVA. **Table 4-3** lists the 90 hours of Scientific Aviation research flights made during FAST-LVOS. All of the aircraft measurements are shown in Appendix E.

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Table 4-3. Scientific Aviation research flights during FAST-LVOS

Date	Takeoff (PDT)	Landing (PDT)	Duration (hours)
May 23	11:53	16:16	4:23
May 24	10:05	16:12	6:07
May 25	9:49	15:19	5:30
			0:00
May 31	10:15	16:19	6:04
June 1	12:13	18:27	6:14
June 2	12:30	18:47	6:17
			0:00
June 10	11:54	18:06	6:12
June 11	11:51	13:13	1:22
June 12	12:08	19:26	7:18
June 13	12:15	18:44	6:29
June 14	12:17	19:00	6:43
			0:00
June 27	12:10	18:43	6:33
June 28	12:11	18:50	6:39
June 29	12:21	19:24	7:03
June 30	12:22	19:12	6:50

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5. FAST-LVOS Modeling Support

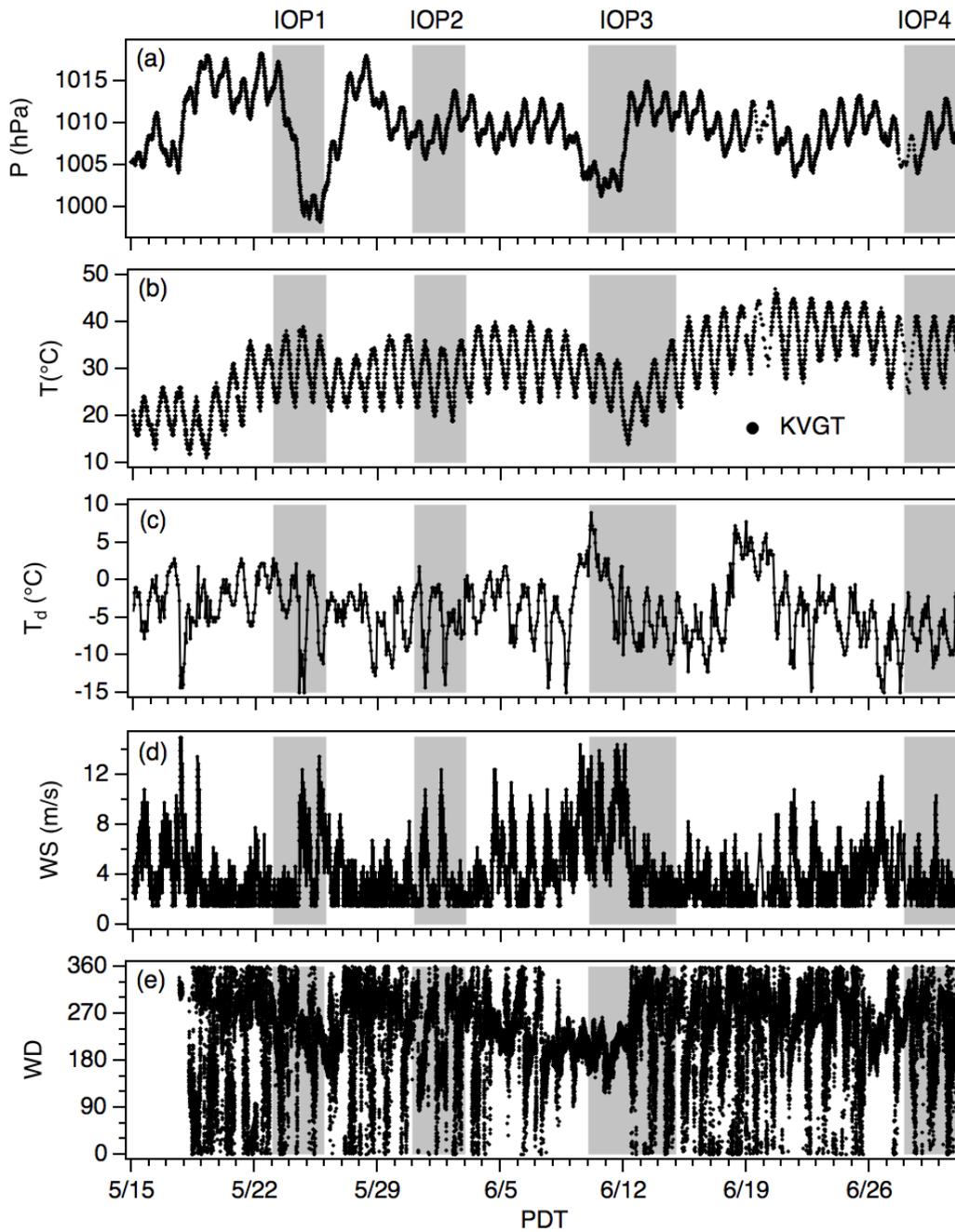
The FAST-LVOS project relied on several atmospheric models for planning purposes during the measurement phase of the campaign. Synoptic weather forecasts, radiosonde profiles, and upper air analyses from the operational NOAA NCEP Global Forecast System (GFS) were obtained from the Wyoming Weather Web maintained by the University of Wyoming (<http://weather.uwyo.edu/index.shtml>). These sources were used in conjunction with O₃ and CO forecasts from the NOAA/ESRL/GSD RAP-Chem (<https://rapidrefresh.noaa.gov/RAPchem/>) provided by Mariusz Pagowski, and the NOAA NESDIS RAQMS (<http://raqms-ops.ssec.wisc.edu/>) model developed by R. Bradley Pierce to plan the daily operations and IOPs. This information was assimilated by the Principal Investigator who recalled the ozonesonde and aircraft teams to Las Vegas when synoptic conditions favorable for STT were developing.

Several other models aided the interpretation of the measurements. Meiyun Lin and Alex Zhang of NOAA GFDL and Princeton University used two global models (GFDL-AM4 and GEOS-Chem) to investigate the various sources for the high O₃ observed above Clark County during FAST-LVOS. These efforts were directly supported by Clark County under a separate contract (CBE-605334-19) and are described in more detail in a separate report. Jerome Brioude and Stephanie Evan of the Université de La Réunion provided FLEXPART particle dispersion model simulations of stratospheric ozone, Asian pollution, and biomass burning influences during the campaign. Wildfire influences were also assessed using an early version of the NOAA/ESRL/GSD Rapid Refresh-Smoke (HRRR-Smoke) air quality modeling system (<https://rapidrefresh.noaa.gov/hrrr/HRRRsmoke/>) developed by Ravan Ahmadov, which simulated the emissions and transport of smoke from wildfires detected by the VIIRS/JPSS satellite fire product in high spatial resolution (3km) over the CONUS domain. Appendix A describes these models in more detail.

6. Meteorological Context

Summer-like conditions arrive early in the Desert Southwest and the warmer temperatures lead to the development of regional-scale southwest-northeast plains-mountain circulations and locally-driven valley and slope flows in the Las Vegas Valley [Stewart *et al.*, 2002]. These diurnal wind patterns lead to southeasterly to southerly flow during the morning transition, but the winds shift to the southwest by mid-afternoon as the mixed layer grows in depth, and plains-mountain winds driven by the thermal contrast between the land and Gulf of California develop. This regional-scale flow converges with southeasterly up-valley flow in the LVV, and these winds typically persist until well into the night when downslope flows from the east and southwest converge in Las Vegas. This diurnal pattern can be seen in the wind measurements at both the NLVA and AP plotted in the lower panels of **Figures 6-1** and **6-2**, respectively.

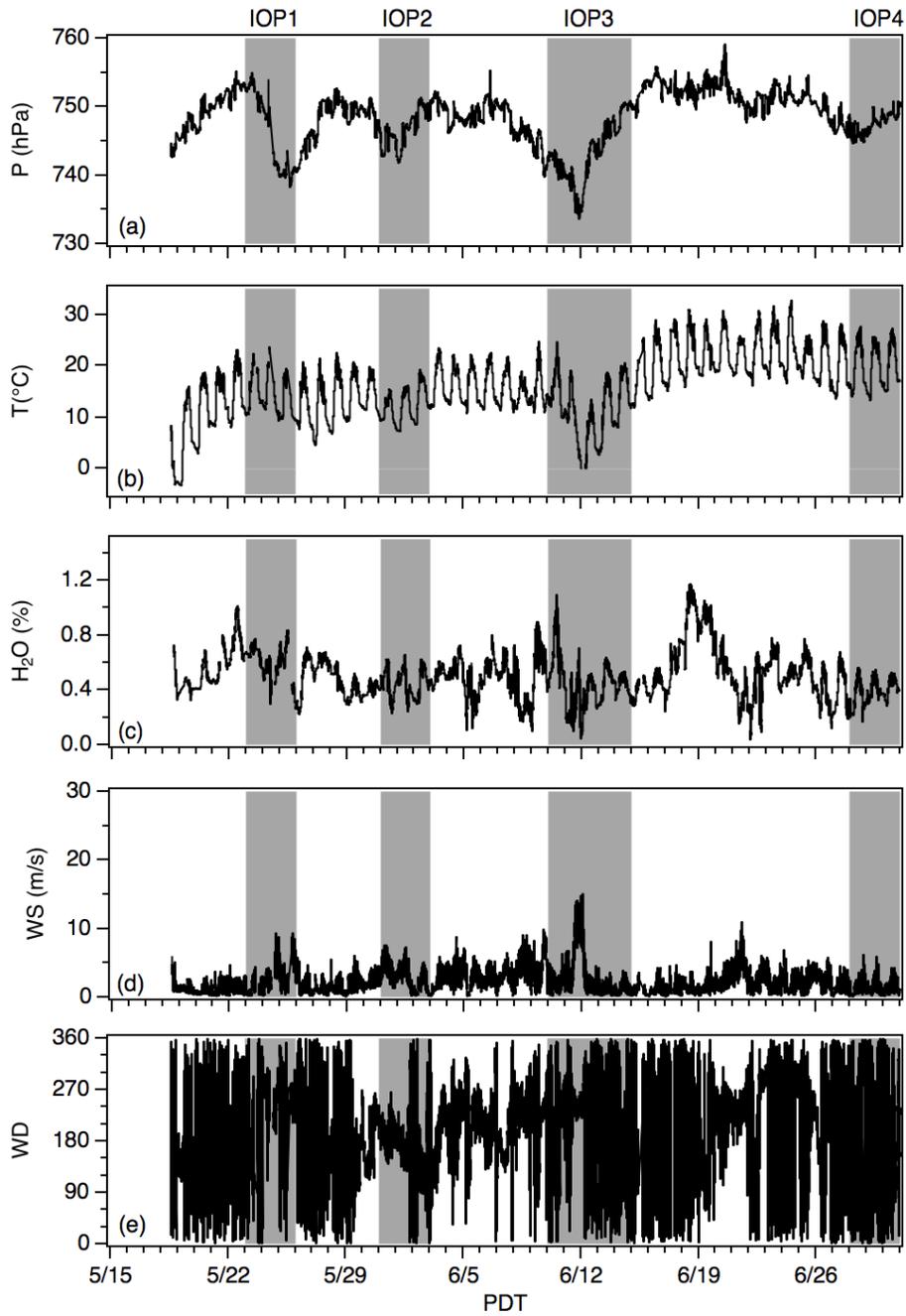
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Figure 6-1. Time series showing the 5-min measurements from the NWS (KVG T) station located near the TOPAZ truck at the NLVA. The gray bands mark the four FAST-LVOS IOPs.

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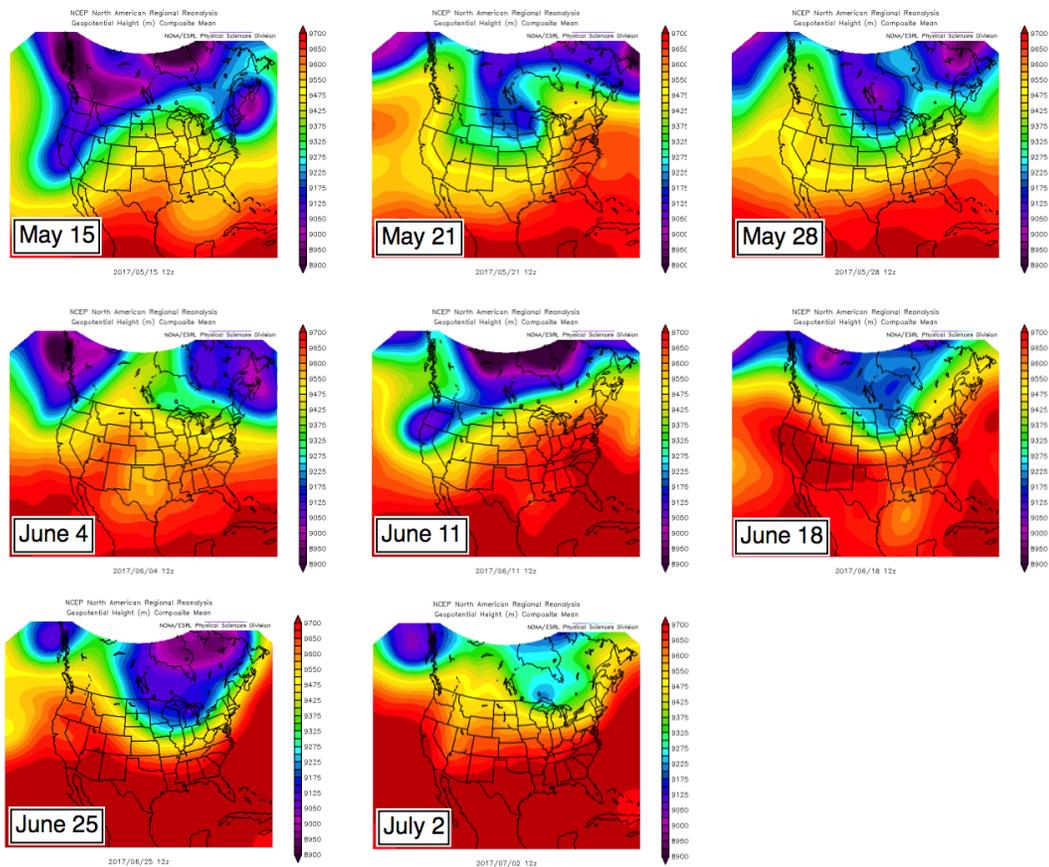
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Figure 6-2. Time series of the 1-min meteorological measurements from the mobile laboratory on Angel Peak (black). The gray bands mark the four FAST-LVOS IOPs.

1 This diurnal pattern is periodically interrupted by the passage of deep midlatitude cyclones that
2 gradually become shallower and less frequent as the jet stream migrates northward into
3 Canada in June. This gradual shift is seen in **Figure 6-3**, which shows 300 hPa geopotential plots
4 from the NCEP NARR Reanalysis at weekly intervals from May 15 to July 2, 2017. The changes in
5 pressure, temperature, and wind speed and direction during the passage of the upper level
6 troughs and their associated surface cold fronts can be seen in the NLVA and AP
7 measurements. There was no measurable precipitation at the NLVA during the entire 6-week
8 campaign.

9
10 The temperatures in the LVV climbed dramatically after the passage of the last deep low on
11 June 12, as strong ridging brought an incursion of moist subtropical air into southern Nevada.
12 The maximum high temperature of 117°F (47.2°C) on June 20 tied the all-time Las Vegas record
13 high. These high temperatures were briefly moderated by the passage of the shallow trough
14 seen above northern California on June 11, and returned to near climatological values during
15 the passage of the final low-pressure system in the last days of the campaign.

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20 **Figure 6-3.** NCEP NARR Reanalysis 300 hPa 12UT geopotential maps at weekly intervals during
21 FAST-LVOS.

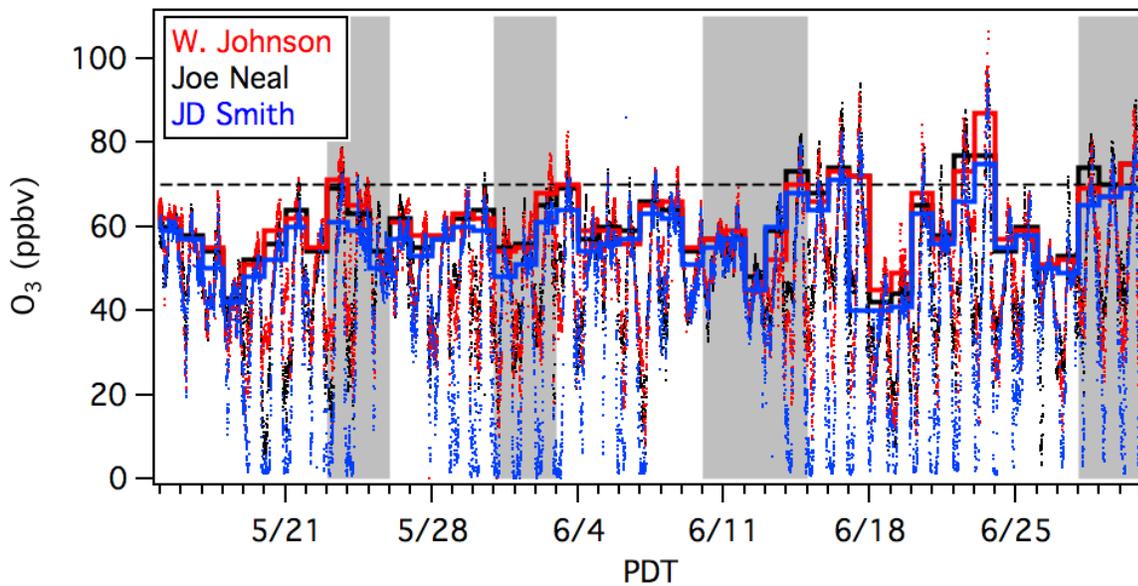
1 **7. Overview of the FAST-LVOS Measurements**

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3 **7.1 Clark County measurements**

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5 The Clark County Department of Air Quality maintains a network of continuous air monitoring
6 sites (CAMS) in the Las Vegas Valley and surrounding areas. This network included 11 active
7 ozone monitors (cf. **Figure 2-1**) during the FAST-LVOS campaign, including two (JD Smith and
8 Apex) that have subsequently been deactivated. The measurements from these regulatory
9 monitors provide context for the following discussions of the FAST-LVOS measurements. **Figure**
10 **7-1** plots time series of the 5-min and MDA8 O₃ measurements from the Walter Johnson, Joe
11 Neal, and JD Smith monitors. These monitors were all located about 7 km from the NLVA, which
12 lies near the center of an equilateral triangle formed by these three sites. The Walter Johnson
13 and Joe Neal monitors recorded the largest number of ozone exceedances during FAST-LVOS
14 and **Figure 7-1** shows that most of those high ozone days occurred during or immediately after
15 one of the FAST-LVOS IOPs.
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19 **Figure 7-1.** MDA8 (steps) and 5-min (dots) O₃ mixing ratios measured by the Walter Johnson
20 (red), Joe Neal (black), and JD Smith (blue) monitors during FAST-LVOS. The horizontal dashed
21 line shows the 70 ppbv NAAQS.

22

23 **Table 7-1** lists all of the instances when one or more of the regulatory monitors equaled or
24 exceeded the 2015 NAAQS during FAST-LVOS. This occurred on 10 of the 45 measurement days,
25 with 3 or more monitors at or above the NAAQS on 6 of those 10 days. These days (June 16, 17,
26 22, 23, 28, and 30) all occurred during the last two weeks of the field campaign.

27

28

Table 7-1. MDA8 O₃ on high ozone days during FAST-LVOS

	CAMS	May 23	June 3	June 14	June 16	June 17	June 22	June 23	June 28	June 29	June 30
Mesquite	23	49	52	54	54	50	61	51	62	59	60
Paul Meyer	43	65	66	67	67	64	64	71	61	63	68
Walter Johnson	71	71	70	70	73	72	73	87**	69	67	75
Palo Verde	73	68	66	68	75	67	73	76	69	65	72
Joe Neal	75	69	69	73	74	72	77	77	74	70	75
Green Valley	298	57	64	65	68	72	63	65	65	68	69
Jerome Mack	540	53	60	62	65	66	60	65	59	61	63
Boulder City	601	51	58	63	66	64	66	62	69	65	66
Jean	1016	51	54	57	63	63	62	60	65	63	69
JD Smith*	2002	61	64	68	71	70	66	75	65	67	69
Apex*	22	55	62	65	68	68	70	64	71	66	65
NLVA#	N/A	72	70	74	76	73	76	83	74	71	76
AP#	N/A	64	60	71	74	61	63	53	68	71	74

* now inactive

** MDA8 based on 6-hour average because of maintenance

NOAA/ESRL/CSD measurements

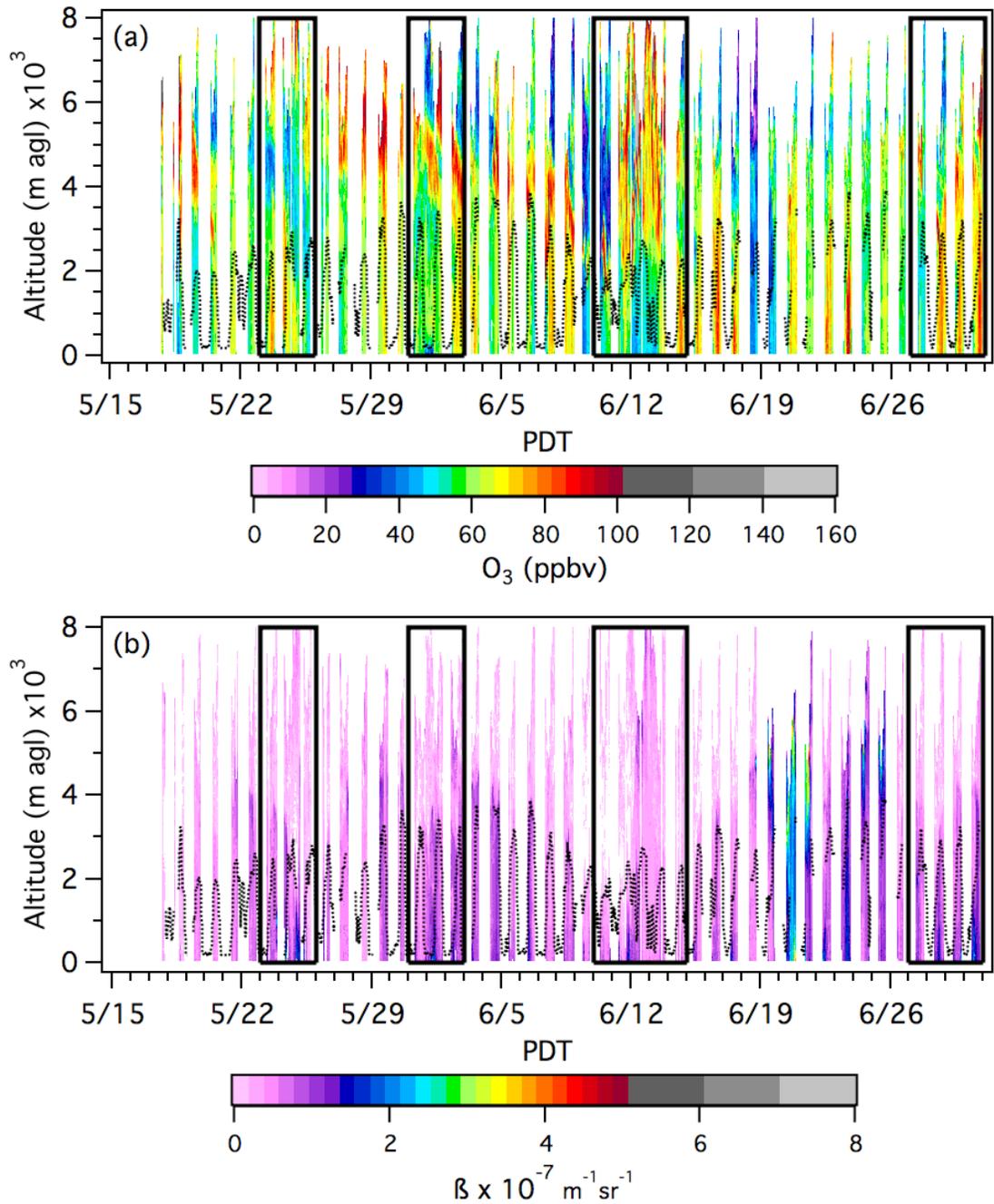
7.2 TOPAZ lidar measurements

The TOPAZ lidar acquired ozone and backscatter data on all 45 days (May 17-June 30) of the FAST-LVOS campaign, operating a total of 523 hours for an average of nearly 12 hours per day. The O₃ and backscatter (β) measurements are summarized as a series of time-height curtain plots in **Figure 7-2**. The dotted black lines show the mixed layer heights inferred from the co-located micro-Doppler lidar and the heavy black boxes show the IOP days when aircraft and/or ozonesonde measurements were also made.

The upper panel of **Figure 7-2** shows high O₃ aloft during much of FAST-LVOS including the slowly descending band of high O₃ at the end of May and beginning of June, and the deep tropopause fold with mixing ratios exceeding 270 ppbv at 6 km asl (5.3 km agl) on June 11-13. This elevated O₃ sometimes reached the top of the mixed layer, which was usually more than 3 km deep during June. These general observations are consistent with the measurements from the first LVOS campaign.

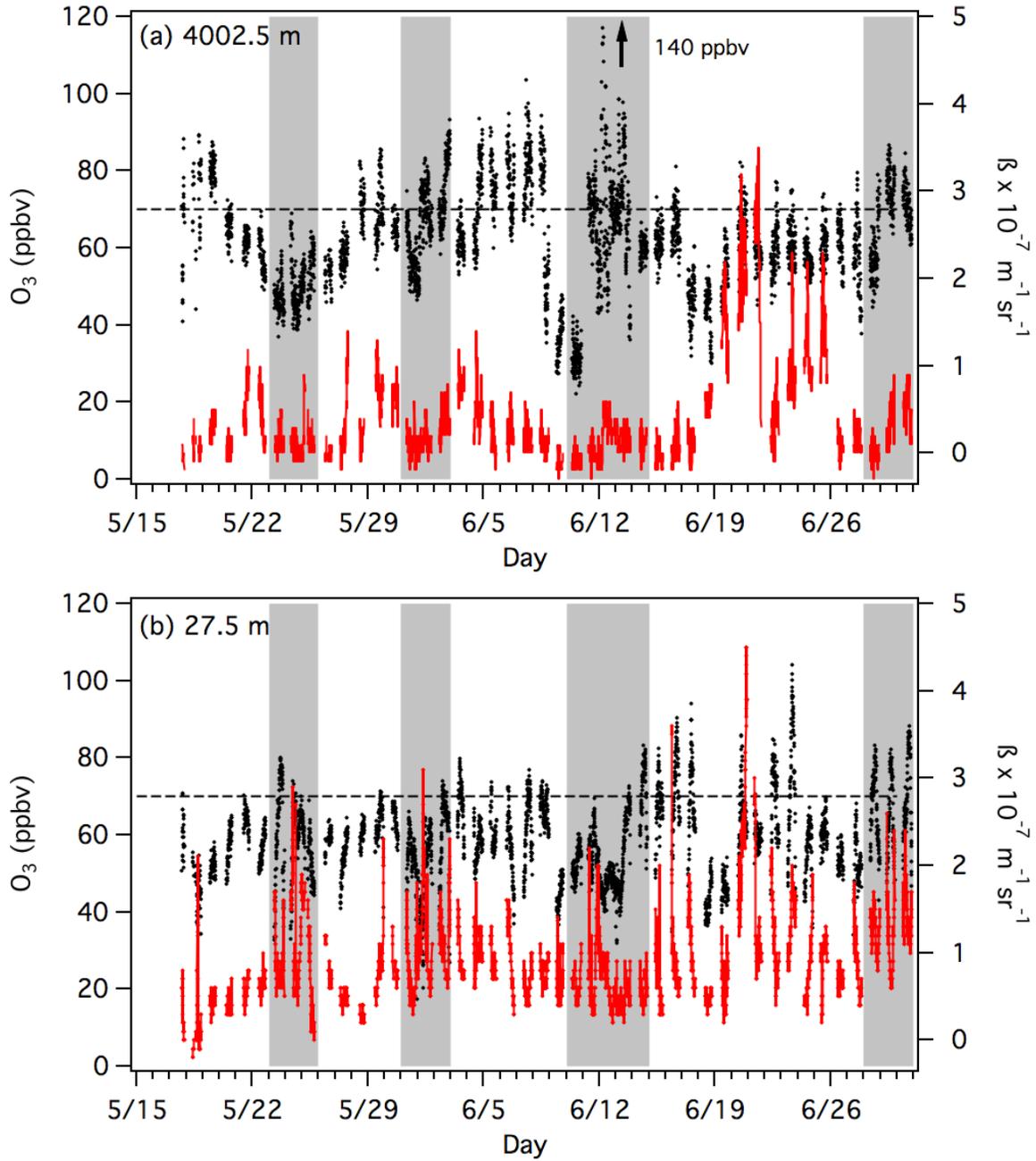
Figure 7-3 plots time series of the O₃ mixing ratios and backscatter coefficients measured above the top of the mixed layer at 4002.5 m agl, and just above the surface at 27.5 m agl. The O₃ mixing ratios at both altitudes often exceeded 70 ppbv, but higher O₃ was measured aloft more

- 1 frequently during the first two weeks of June, and near the surface during the last two weeks.
- 2 The highest backscatter was measured between June 19 and 21 at both altitudes.
- 3



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- 6 **Figure 7-2.** Time-height curtain plots of the TOPAZ (a) O₃ mixing ratios and (b) particulate
- 7 backscatter coefficients. The dotted black line shows the mixed layer heights inferred from the
- 8 micro-Doppler lidar measurements and the heavy black boxes enclose the IOP days.
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Figure 7-3. Time series of the TOPAZ O₃ mixing ratios and particulate backscatter coefficients at (a) 4002.5 and (b) 27.5 m agl. Note that the peak O₃ value of 140 ppbv on June 12 is off scale. The dashed black lines correspond to the 70 ppbv NAAQS.

1 **7.3 Comparison between lidar and surface measurements**
2

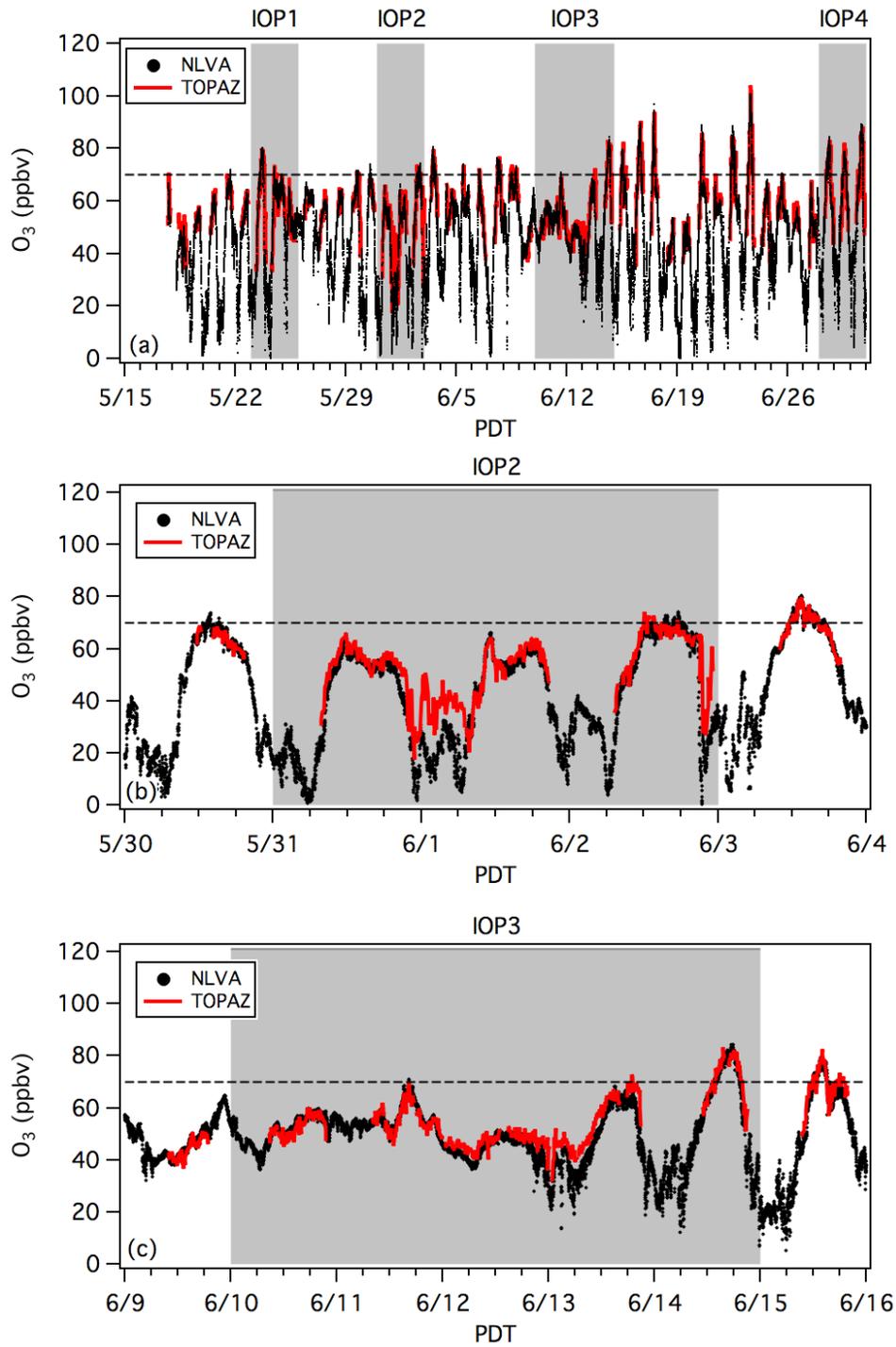
3 The variations in surface O₃ at the NLVA are more easily seen in **Figure 7-4**, which compares the
4 lowest elevation (27.5 m agl) TOPAZ O₃ retrievals with the in-situ measurements made at the
5 TOPAZ truck (5 m agl). The time series are in excellent agreement during the day when the
6 boundary layer was well-mixed, but usually diverged at night (e.g. June 1-2) when a shallow
7 nocturnal boundary layer formed and surface O₃ was destroyed by deposition and NO_x titration.
8 The latter was particularly significant at the NLVA sampling site which was located next to a
9 staging area for heavy construction vehicles that were often left idling during the early morning
10 hours. These local NO_x sources did not affect the TOPAZ mixing ratios, which were measured
11 about 800 m downrange of the TOPAZ truck. Note that the in-situ and lidar measurements did
12 not diverge during the night of June 11-12 when high winds associated with the cold front and
13 descending tropopause fold kept the lower troposphere well mixed.
14

15 **Figure 7-5** is similar to **Figure 7-4**, but compares the 27.5 m agl TOPAZ measurements with the
16 5-min O₃ mixing ratios measured by the Joe Neal monitor located 7.3 km to the NNW. The
17 agreement between these measurements is also excellent with much less divergence at night,
18 confirming that the TOPAZ measurements were representative of the North Las Vegas area.
19

20 **Figure 7-6** compares the in-situ surface measurements from the NLVA with those from the
21 CRDS instrument in the NOAA ESRL mobile laboratory on Angel Peak (2.68 km asl). The absence
22 of a strong diurnal variation in the AP measurements shows that mobile laboratory instruments
23 usually sampled free tropospheric air during the night. Surface O₃ was usually lower on AP in
24 the afternoon, but the AP and NLVA time series often converged (e.g. May 31-June 1) when the
25 mixed layer was sufficiently deep. The daytime O₃ concentrations on AP exceeded those at the
26 NLVA on only a few days (June 5 and June 9-11). **Figure 7-6** also compares the AP
27 measurements with the TOPAZ measurements made at the same altitude (2002.5 m agl or 2.68
28 km asl). These measurements were usually in good agreement despite the 36 km separation,
29 but TOPAZ measured higher concentrations when ozone was descending from the upper
30 troposphere (e.g. June 10-12).
31

32 **Figures 7-7** and **7-8** compare the AP O₃ measurements to the other in-situ measurements from
33 the mobile laboratory. As noted above and shown below, the correlations between these
34 species and O₃ can help elucidate the origins of the sampled airmasses. Particularly striking are
35 the high H₂O and low O₃, CO, and CH₄ concentrations during the subtropical incursion on June
36 18-19 and the off scale (450 ppbv) CO peak on June 21-22 coinciding with the enhancements in
37 TOPAZ backscatter.
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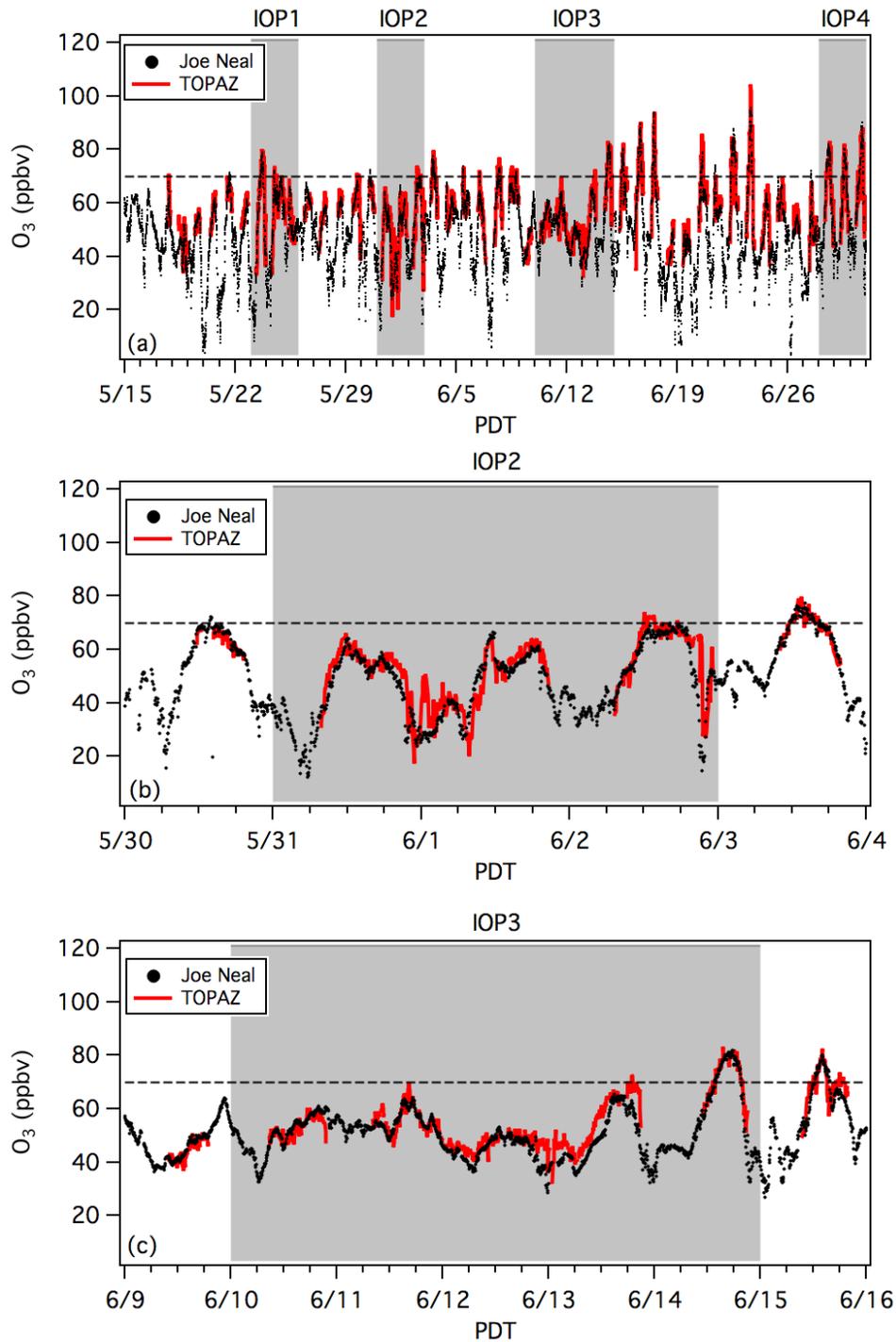
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Figure 7-4. (a) Time series of the 27.5 m TOPAZ O₃ measurements and the 1-min in-situ measurements at the NLVA. The gray bands identify the FAST-LVOS IOPs. The lower panels are expanded views of the measurements from (b) IOP2 and (c) IOP3. The dashed black lines correspond to the 70 ppbv NAAQS.

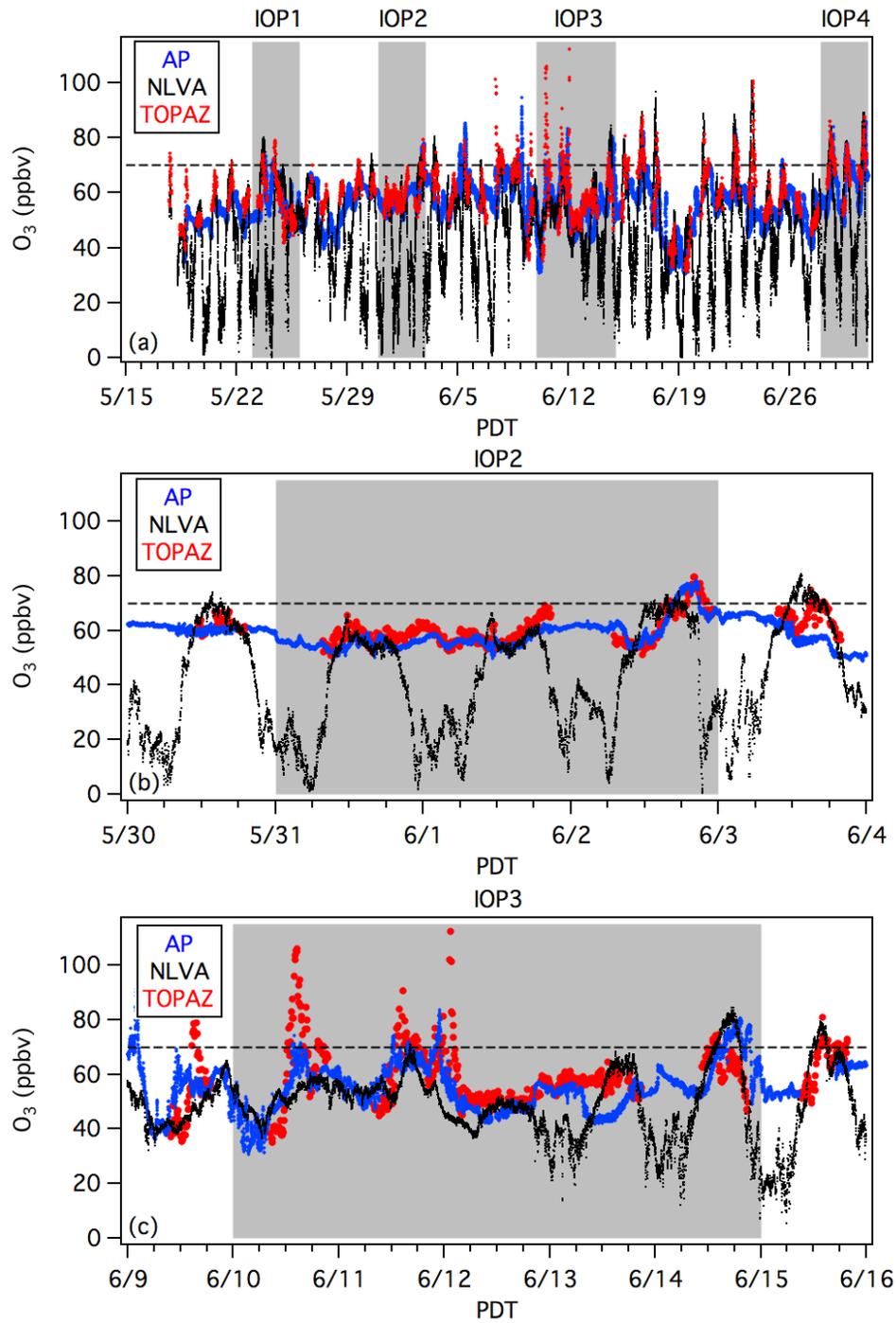
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5 **Figure 7-5.** (a) Time series of the 27.5 m TOPAZ O₃ measurements and the 5-min in-situ O₃
6 measurements from Joe Neal (C75). The gray bands identify the FAST-LVOS IOPs. The lower
7 panels are expanded views of the measurements from (b) IOP2 and (c) IOP3. The dashed black
8 lines correspond to the 70 ppbv NAAQS.

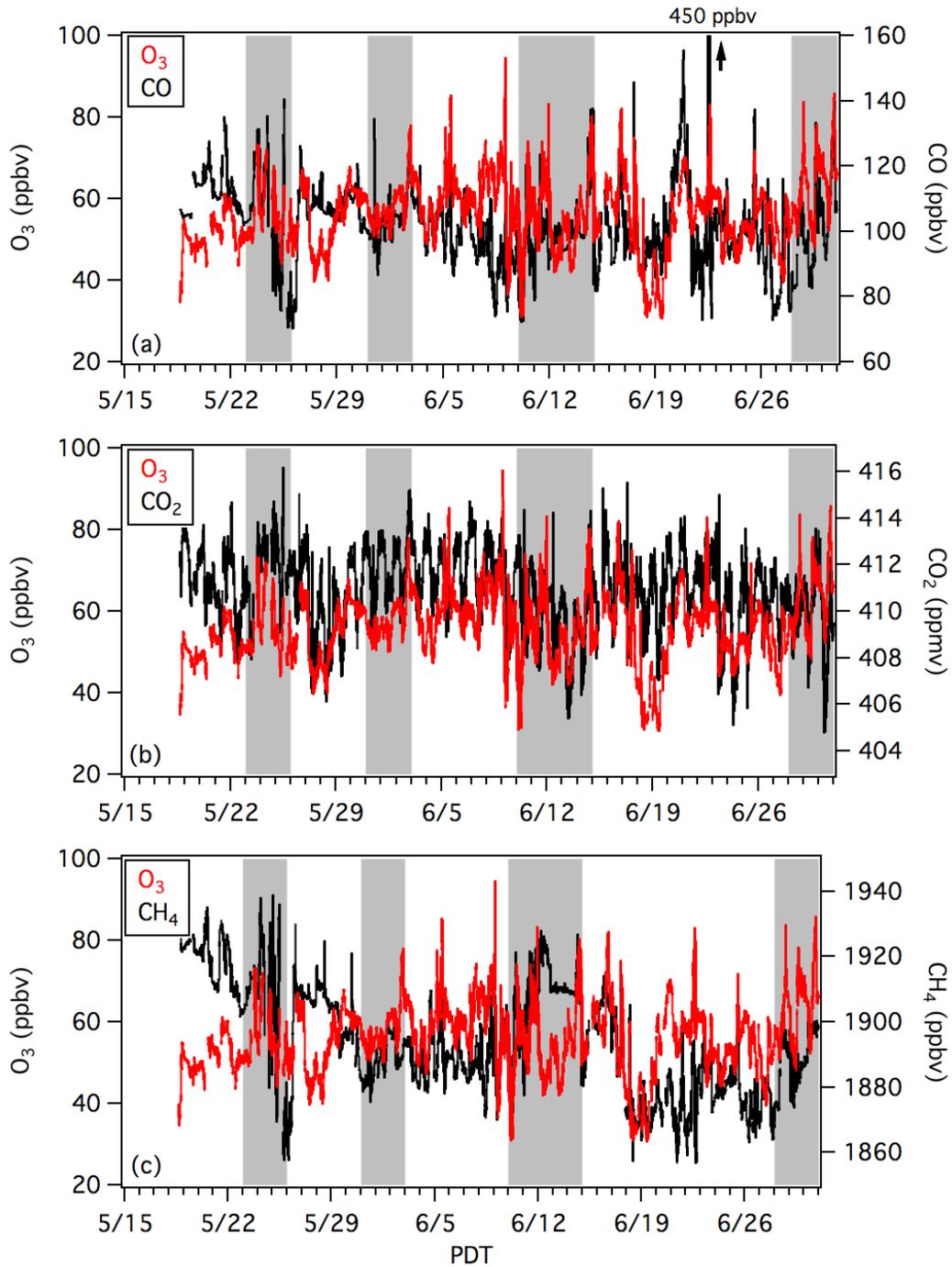
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5 **Figure 7-6.** (a) Time series of the 1-min in-situ O_3 measurements from the NLVA (black) and AP
6 (blue) compared to the 2002.5 m agl (2683.5 km asl) TOPAZ measurements (red). The gray
7 bands identify the FAST-LVOS IOPs. The lower panels are expanded views of the measurements
8 from (b) IOP2 and (c) IOP3. The dashed black lines correspond to the 70 ppbv NAAQS.

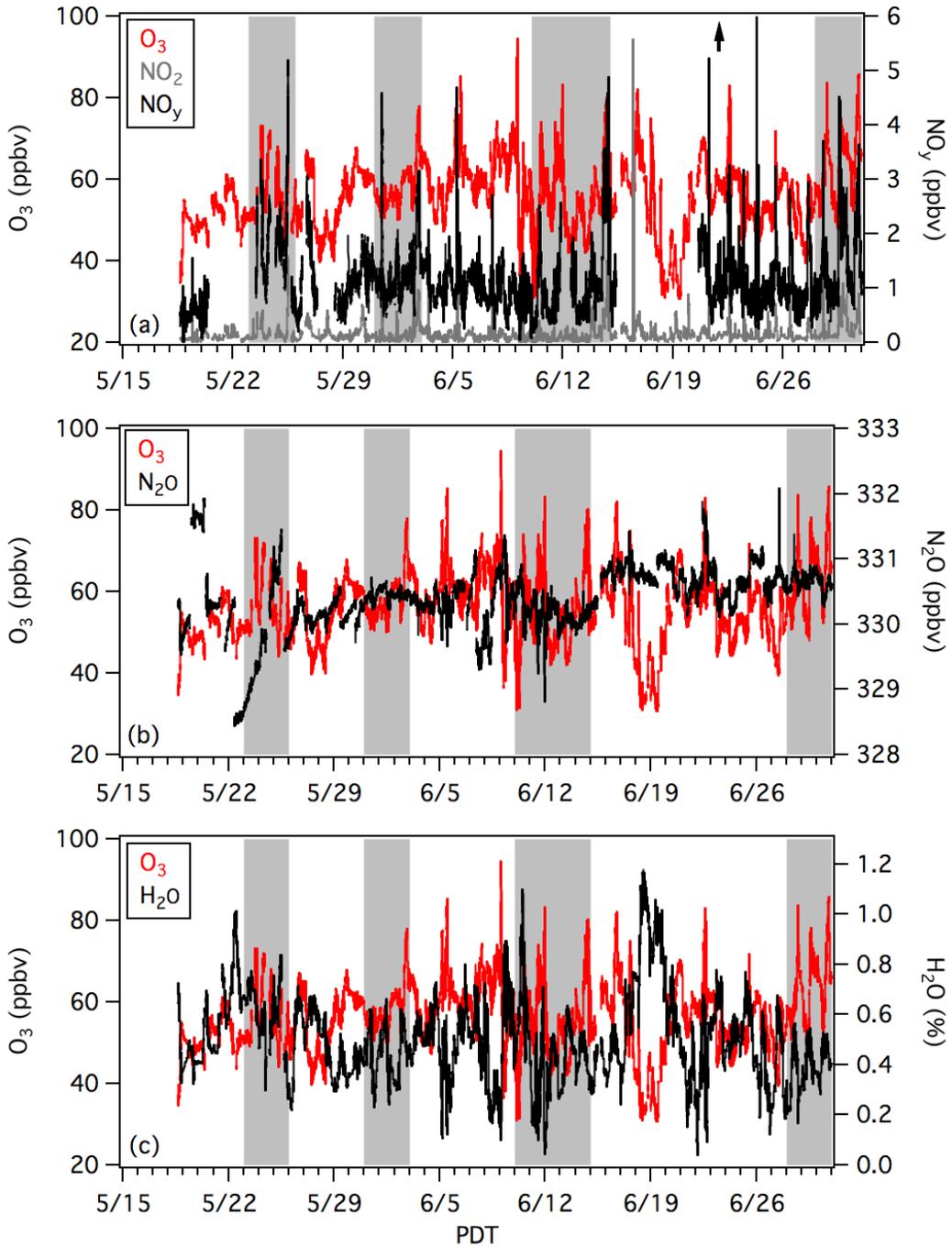
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Figure 7-7. Time series of the 1-min O₃, CO, CO₂, and CH₄ measurements from the mobile laboratory on Angel Peak (black). The gray bands mark the four FAST-LVOS IOPs. The CO concentrations on June 22 reached 450 ppbv (arrow).

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4 **Figure 7-8.** Time series of the 1-min O₃, NO_y, N₂O, and H₂O measurements from the mobile
5 laboratory on Angel Peak (black). The gray bands mark the four FAST-LVOS IOPs.
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8. Comparison to FLEXPART tracer distributions

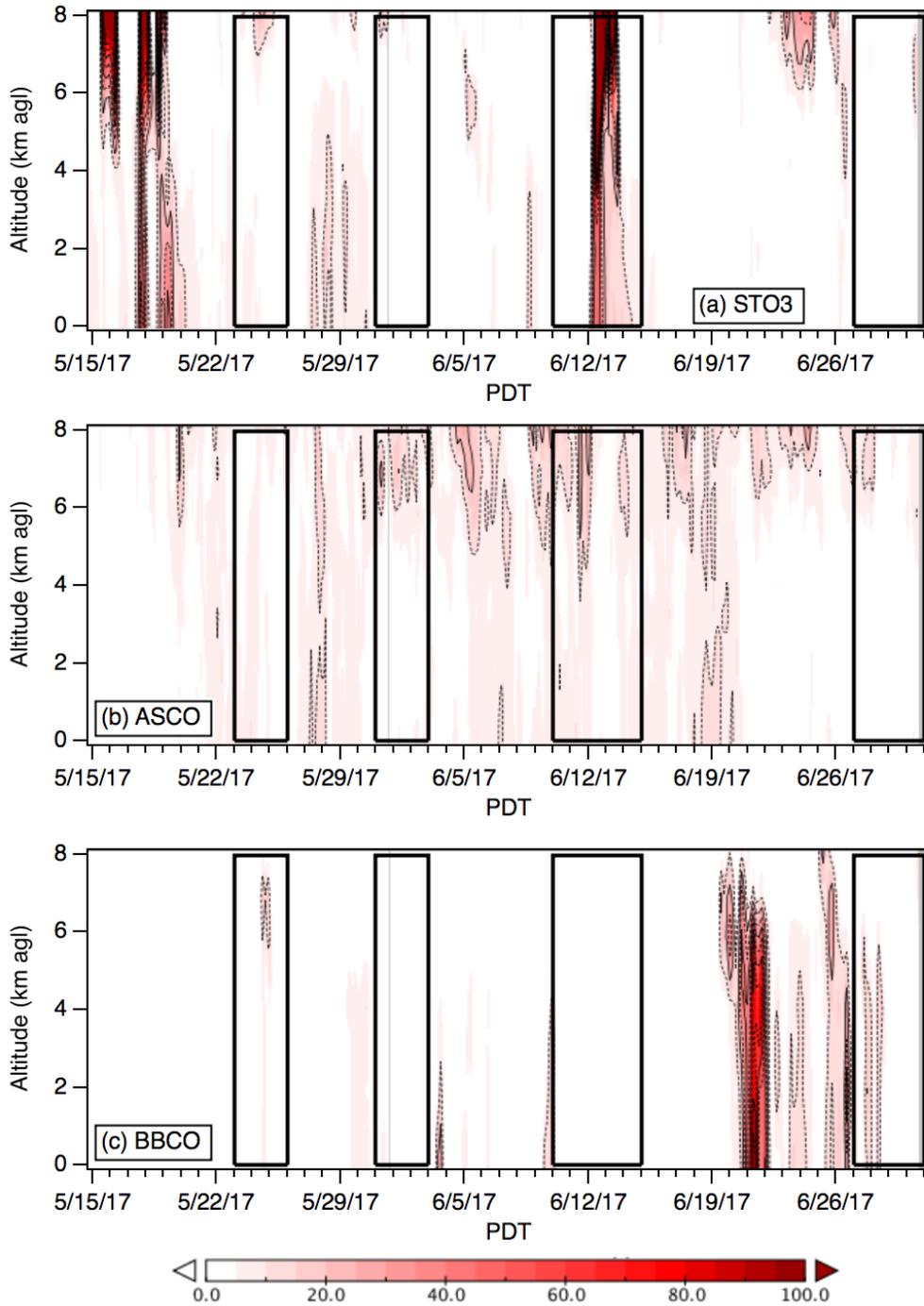
Figure 8-1 shows time-height curtain plots of the FLEXPART tracer distributions (**Appendix A**) over the $0.5^\circ \times 0.5^\circ$ grid cell centered at 36.25°N , -119.25°E . This grid cell encompasses most of the Las Vegas Valley, with Angel Peak lying just beyond the western edge. These plots are analogous to the lidar ozone and backscatter curtains shown in **Figure 7-2**, and qualitatively reproduce many of the features seen in these plots. All three of the tracers are plotted using a 0-100 ppb concentration scale. The ASCO and BBCO tracers can be compared directly, but neither can be quantitatively compared with the STO3 tracer because of the wide range of O_3/CO ratios observed in biomass burning and pollution plumes. An approximate range is on the order of 0.2 to 0.5 {Langford, 2012 #5717}, suggesting that the ASCO and BBCO tracer concentrations can be considered very conservative upper limits for ASO3 and BBO3.

The STO3 tracer plot in **Figure 8-1a** shows significant amounts of stratospheric O_3 penetrating all the way to the surface during two major (May 16-21 and June 11-15) and one minor (May 28-31) intrusion events. The tracer surface concentrations peaked at nearly 60 ppbv at 0200 UT on May 18 and 50 ppbv at 1400 UT on June 12. Note that the first intrusion event was well underway when TOPAZ first started measuring at 0037 UT on June 18 (1737 PDT on June 17).

The ASCO curtain plot in **Figure 8-1b** shows a more continuous influx of Asian pollution in the middle and upper troposphere, with the largest surface contributions (>10 ppbv of CO) on May 28, June 7, and June 18. The biomass burning tracer in **Figure 8-1c** shows little enhancement prior to June 20, but BBCO tracer surface enhancements of up to 100 ppbv occurred during the last 10 days of the study. Note that the BBCO tracer plot appears qualitatively similar to the TOPAZ backscatter curtain (**Figure 7-2**).

Figure 8-2 shows the STO3 tracer distributions for 12UT on May 18, May 27, and June 12 on the 500 and 700 hPa surfaces (roughly 5.6 and 3.1 km asl). The tracer distributions in the May 18 and June 12 plots have the classic appearance of recent tropopause folds curling cyclonically around the low-pressure centers. The older intrusion on May 27 has started to dissipate and is less well-defined. The ASCO tracer distributions in **Figure 8-3** are generally more diffuse, but sometimes appear to be intertwined with descending stratospheric air as on May 27-28. The BBCO distributions in **Figure 8-4** suggest that large backscatter enhancement on June 20 was caused by diffuse smoke from widespread agricultural burning.

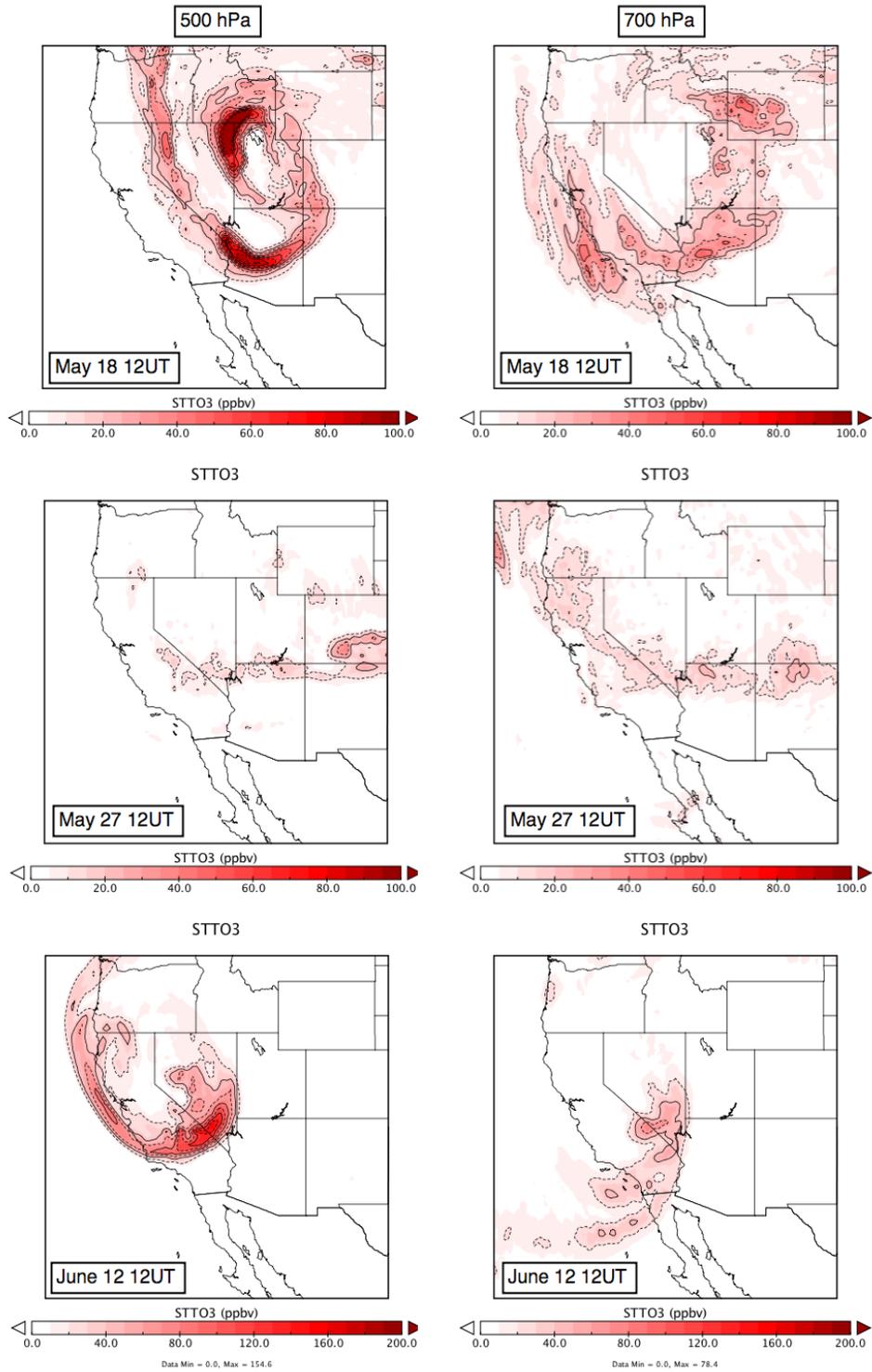
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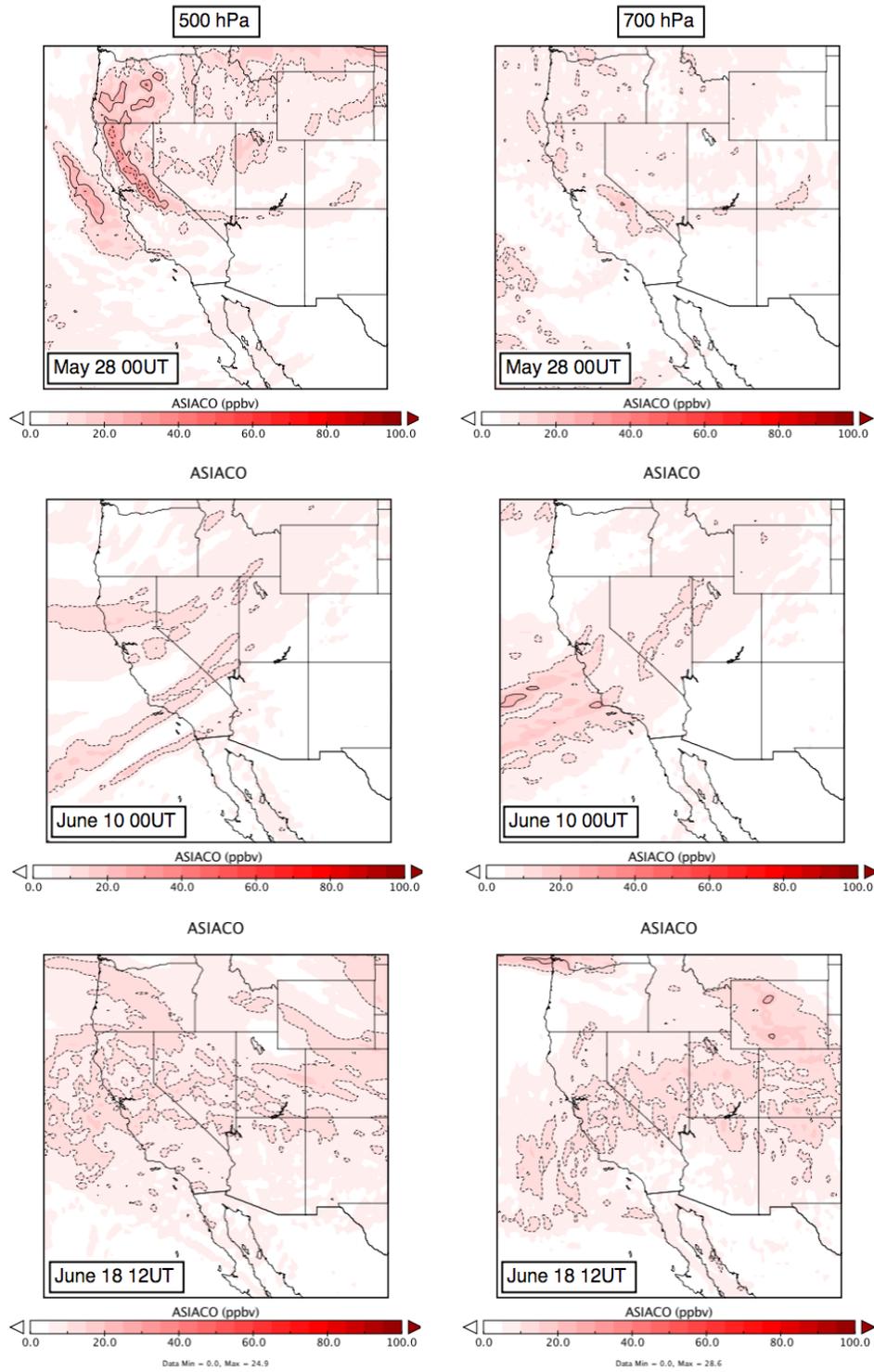
Figure 8-1. Time-height curtain plots of the (a) STT O3, (b) ASCO, and (c) BBCO FLEXPART tracers above the Las Vegas Valley. All three tracers are plotted with a 0-100 ppb concentration scale.

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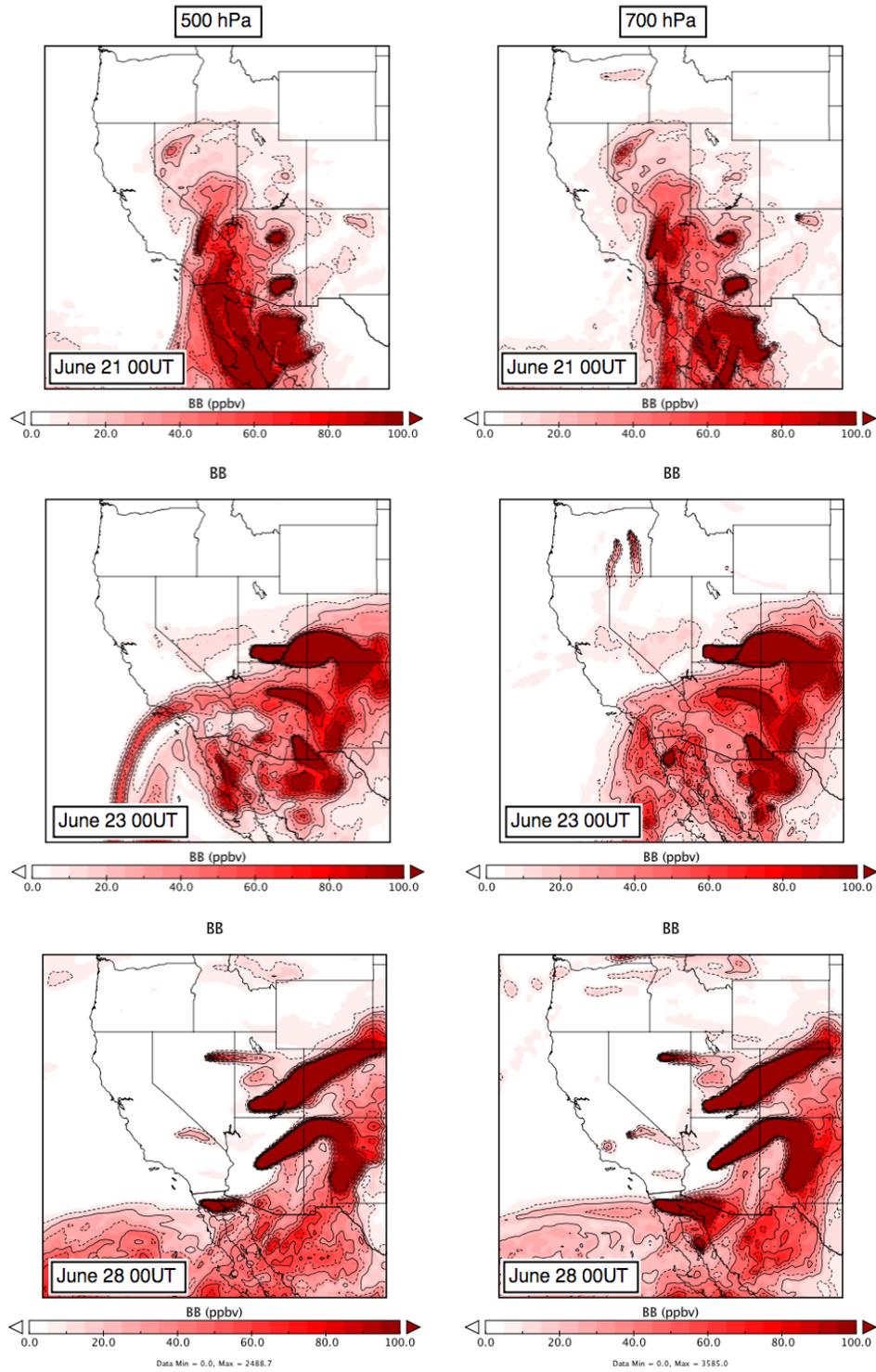
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Figure 8-2. Maps showing the FLEXPART STO3 tracer distribution at 12UT on May 18, May 27, and June 12 on the (left) 500 hPa and (right) 700 hPa surfaces.



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Figure 8-3. Maps showing the FLEXPART ASIACO tracer distributions at 00UT on May 28 and June 10, and 12UT on June 18 on the (left) 500 hPa and (right) 700 hPa surfaces.



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Figure 8-4. Maps showing the FLEXPART BBCO tracer distribution at 00UT on June 21 23, and 28 on the (left) 500 hPa and (right) 700 hPa surfaces.

9. Weekly summaries of the FAST-LVOS Results

The TOPAZ summary plots in **Figure 7.2** hide a wealth of detail in the lidar and in-situ measurements. In the following sections, we take a closer look at the TOPAZ and in-situ measurements in week long segments. We compare the enlarged TOPAZ curtain plots with the corresponding FLEXPART tracer distributions, and with selected plots from the other FAST-LVOS measurements to assess the potential contributions of transport to the high ozone days in Clark County. The remaining FAST-LVOS measurements are summarized in the Appendices.

9.1 Week 1: May 17-20

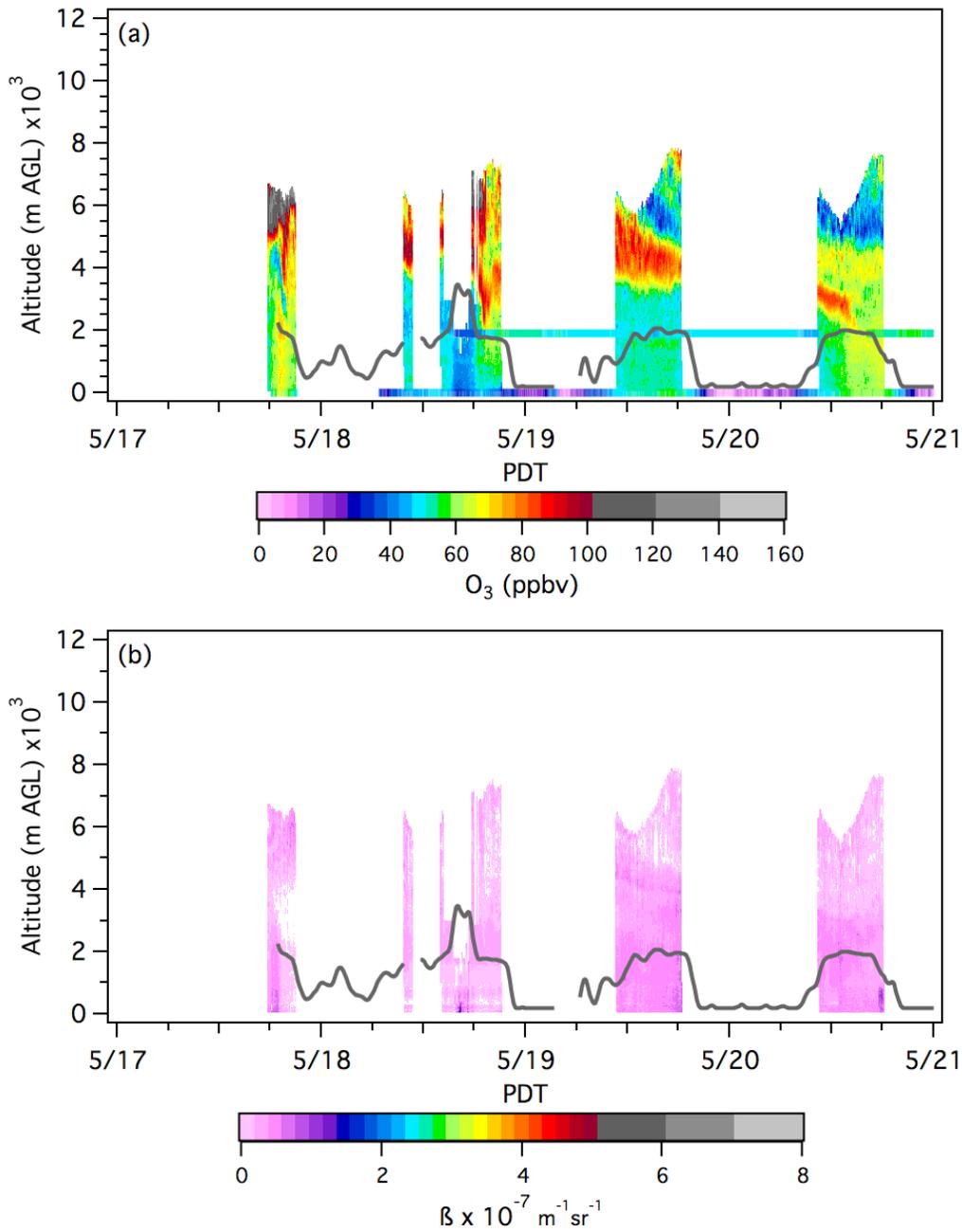
Temperatures in the Las Vegas Valley were well below normal at the start of the FAST-LVOS measurement campaign as a spring storm generated by the first upper level trough shown in **Figure 6-3** passed through Nevada. This trough also spawned the stratospheric intrusion shown in **Figure 8-2**. The TOPAZ truck drove through late season snow in Utah and arrived at the North Las Vegas Airport around 1300 PDT on Wednesday May 17. The first TOPAZ and in-situ measurements at the NLVA were made later that afternoon. The intrusion appears as a layer with more than 100 ppbv of O₃ between 5 and 6 km agl on May 17 in **Figure 9-1**, an expanded version of the curtain plots in **Figure 7-2**. TOPAZ also measured elevated O₃ (with low backscatter) above the top of the mixed layer on the following three days.

Figure 9-2 shows a similar view of the FLEXPART tracer distributions from **Figure 8-1**. The top panel shows a strong influx of stratospheric O₃ into the free troposphere between May 17 and 21 with up to 35 ppbv reaching the surface on the evening of May 17, and 30 ppbv reaching the surface on the late morning of May 19. The TOPAZ measurements appear to show some descent to the top of the boundary layer on May 17, but the surface concentrations remained relatively low. The ASCO tracer plot shows up to 20 ppbv of ASCO above 6 km on May 20, in qualitative agreement with the elevated O₃ layers seen in the TOPAZ measurements, and some of this elevated O₃ appears to have been entrained into the mixed layer on May 20. The FLEXPART BBCO flux was negligible, in agreement with the low backscatter.

The mobile laboratory began measurements on Angel Peak on May 18, but some of instruments were not fully operational until the following week. The 1-min in-situ O₃ measurements from the NLVA and AP are plotted along with the 5-min measurements from the Walter Johnson and Joe Neal monitors in **Figure 9-3**. The TOPAZ O₃ mixing ratios at 4.0 km agl are also plotted; these were generally larger than the surface values.

The Walter Johnson and Joe Neal measurements show two narrow peaks with up to 68 ppbv around 1705 PDT on May 17 that appear to be correlated with the ozone at 4 km. The TOPAZ curtain plot shows that this spike coincided with the apparent descent and entrainment of O₃ from a layer with more than 100 ppbv that was initially detected by between 5 and 6 km agl. The measurements from May 20 also show possible entrainment and more uniform mixing in the boundary layer, but the MDA8 O₃ concentrations at both Walter Johnson and Joe Neal remained well below the NAAQS on both days.

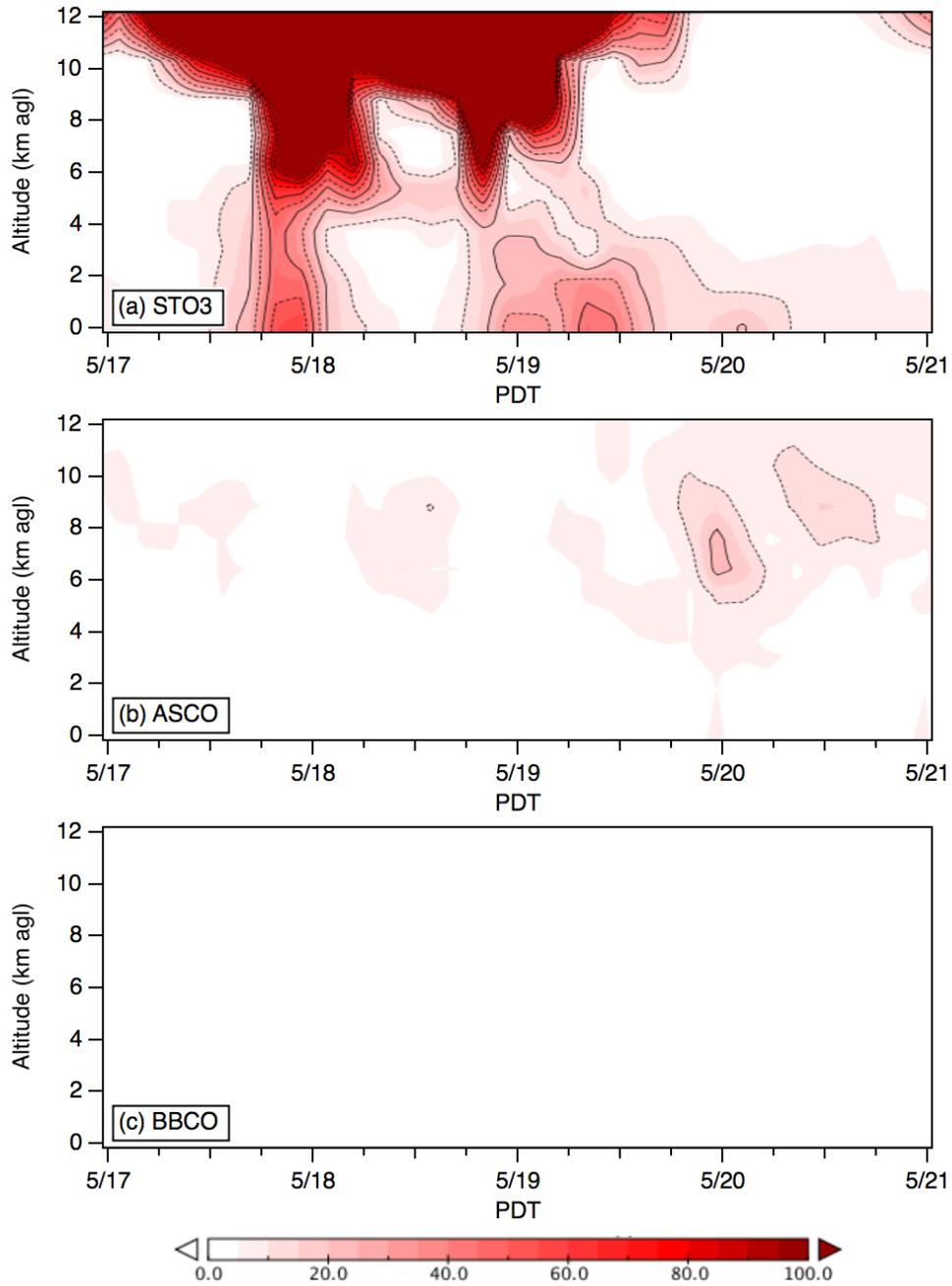
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Figure 9-1. Time-height curtain plots of the TOPAZ (a) ozone, and (b) backscatter, measured during Week 1 of FAST-LVOS. The colored horizontal bands at 2.0 and 0.0 km agl in (a) show the in-situ measurements from the NLVA and AP, respectively. The dark gray trace represents the local mixed-layer height inferred from the micro-Doppler lidar measurements. Note that the TOPAZ measurements were truncated by low clouds during the early afternoon of the 18th.

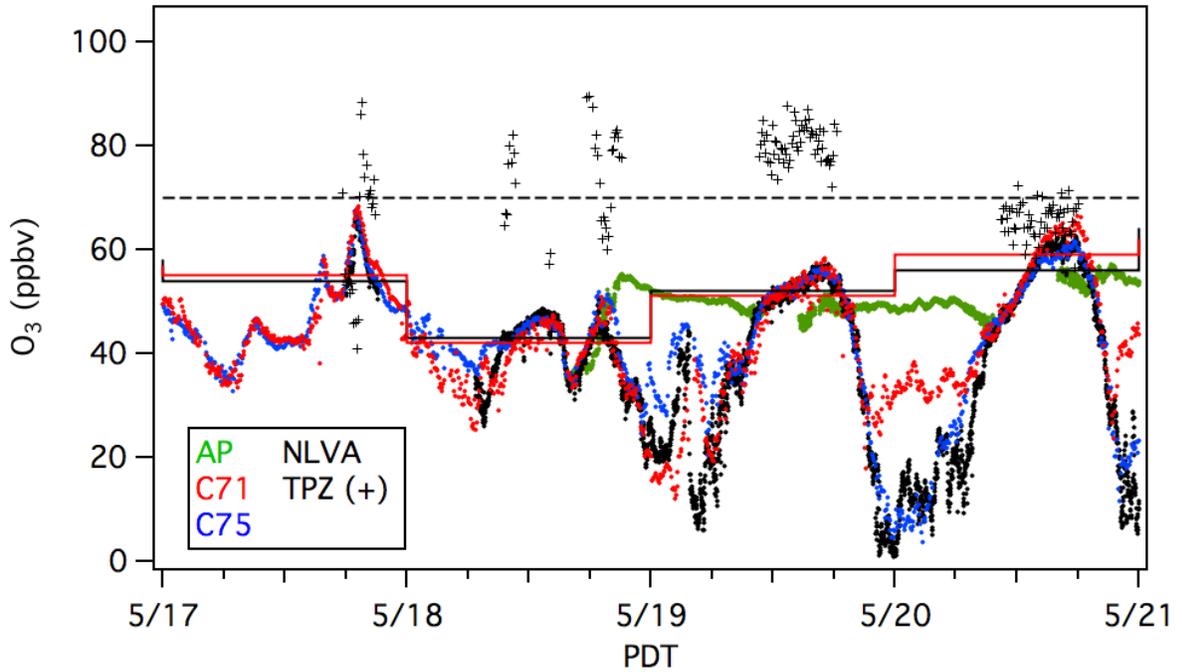
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Figure 9-2. Time-height curtain plots of the FLEXPART (a) STO3, (b) ASCO, and (c) BBCO tracer distributions above the LVV during Week 1 of FAST-LVOS.

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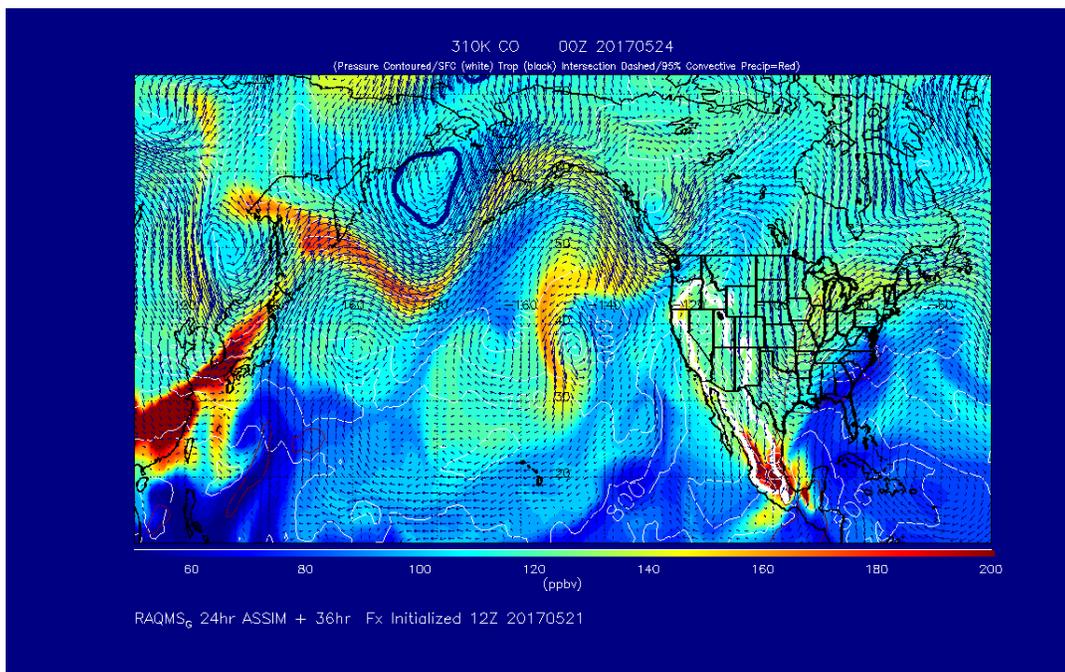
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Figure 9-3. Time series of the TOPAZ O₃ mixing ratios at 4.0 km agl (black +) during Week 1 plotted with the 1-min O₃ measurements from the NLVA (black) and AP (green) and the 5-min measurements from the Walter Johnson (C71, red) and Joe Neal (C75, blue) monitors. The red and blue steps show the corresponding C71 and C75 MDA8 measurements. The dashed black line indicates the 70 ppbv NAAQS.

1 **9.2 Week 2: May 21-28 (IOP1)**

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3 The second week of operations saw a return to normal and above average temperatures with
4 daily highs at the KVGT station reaching 37.8 and 38.9°C (100 and 102°F) on May 23 and 24,
5 respectively. The average temperatures on both days were about 5°C (9°F) above normal. The
6 NESDIS RAQMS model forecast showed a large band of CO from Asia approaching the western
7 U.S. and the first FAST-LVOS IOP was scheduled for May 23-25. Two ozonesondes were
8 launched on May 23, three on May 24, and one on May 25, and the Mooney also conducted
9 research flights to Big Bear, CA and back on all three days.

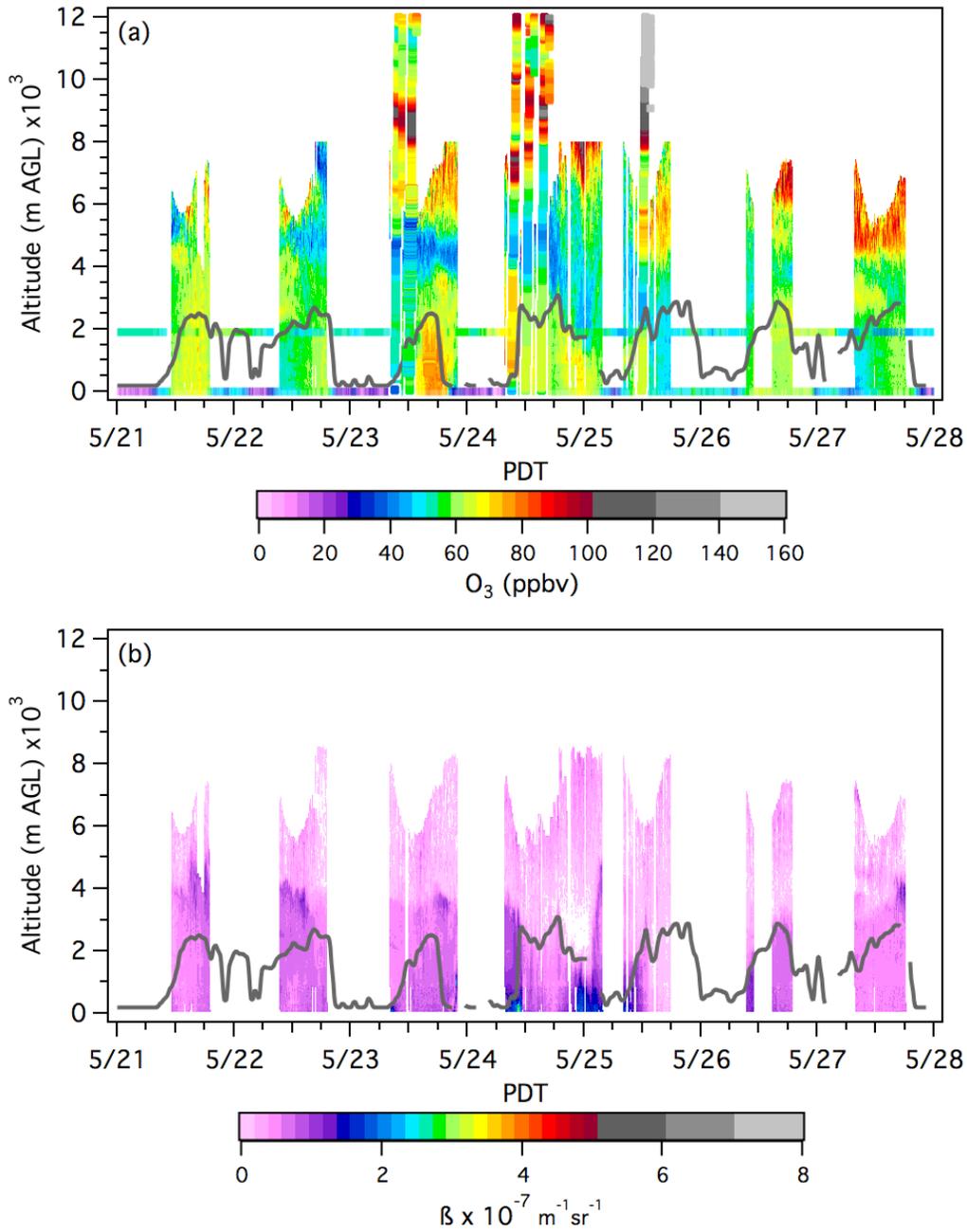
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16 **Figure 9-4.** NOAA NESDIS RAQMS model 310 K CO forecast for May 24 00UT. Forecast initialized
17 May 21 12UT.

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20 The lidar/ozonesonde curtain plots in **Figure 9-5** show that both TOPAZ and the ozonesondes
21 detected high O₃ above 6 km agl most of the week, but they also detected an intervening layer
22 of relatively moist Pacific background air (cf. **Figure 9-6**) with lower O₃ concentrations that
23 mixed down to the surface on May 25.

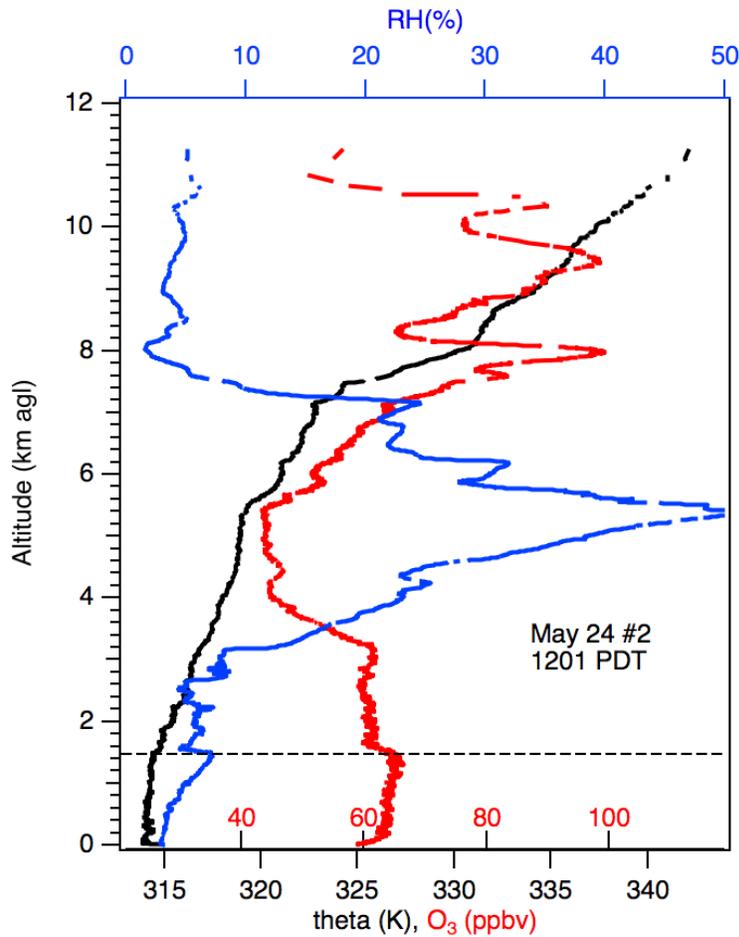
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Figure 9-5. Time-height curtain plots of the TOPAZ (a) ozone, and (b) backscatter, measured during Week 2 of FAST-LVOS. The ascending Joe Neal ozonesonde profiles from IOP1 are superimposed. The colored horizontal bands in (a) show the in-situ measurements from the NLVA and AP. The black line shows the local mixed-layer height from the micro-Doppler lidar.

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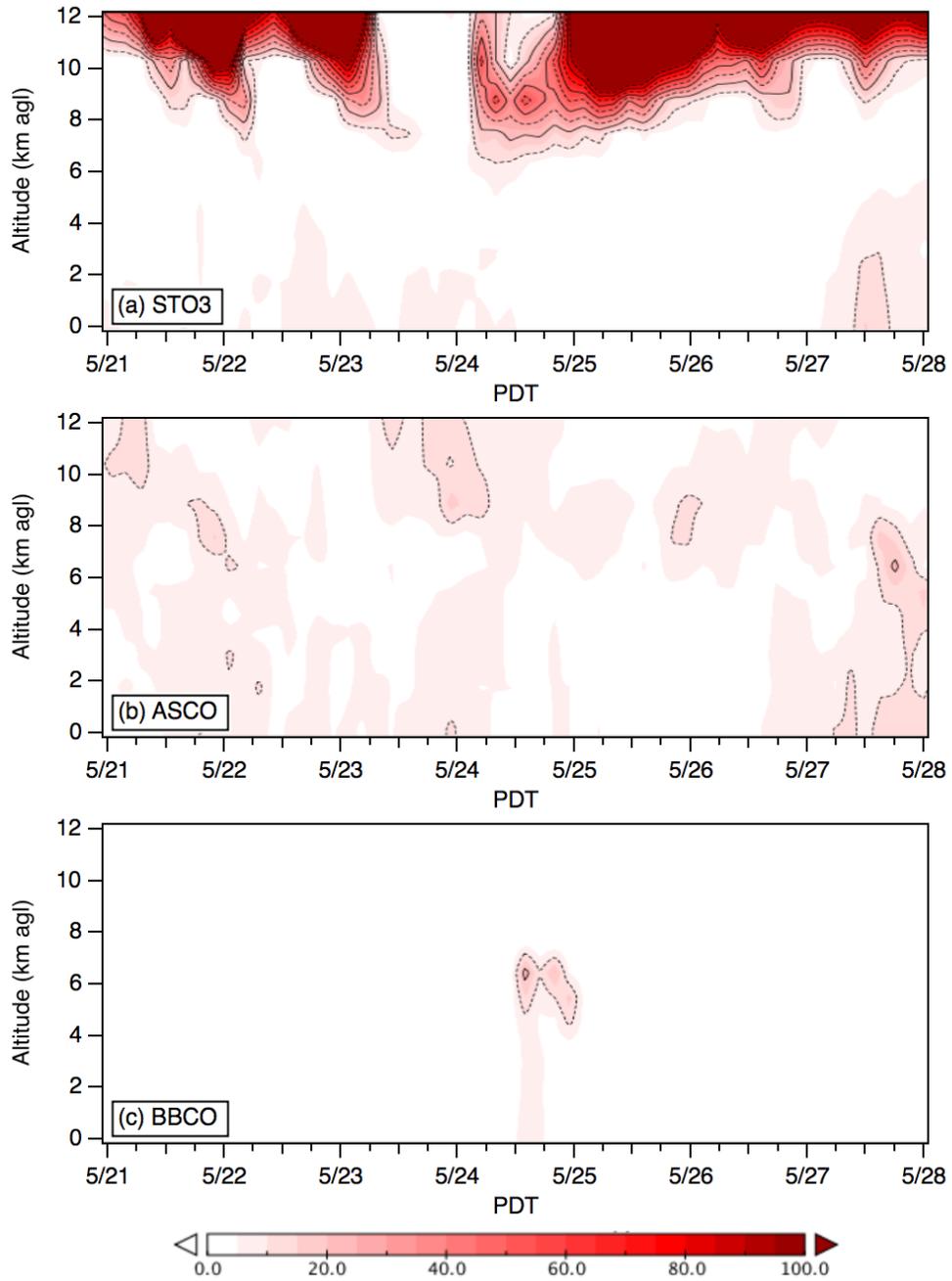
6 **Figure 9-6.** Profiles of potential temperature (black), O₃ (red), and relative humidity (blue) from
7 the Joe Neal 1201 PDT ozonesonde on May 24. The dashed black horizontal line marks the top
8 of the boundary layer.

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11 The FLEXPART STO3 tracer distributions in **Figure 9-7** show stratospheric air in the upper
12 troposphere throughout the week, and the ASCO tracer shows modest influxes of Asian
13 pollution throughout the troposphere. The BBCO tracer also shows a small wildfire plume
14 between the surface and 6 km on the evening of May 24, which is qualitatively consistent with
15 the backscatter measurements. However, none of these tracers show any significant surface
16 impacts and cannot explain the high surface O₃ measured on May 23.

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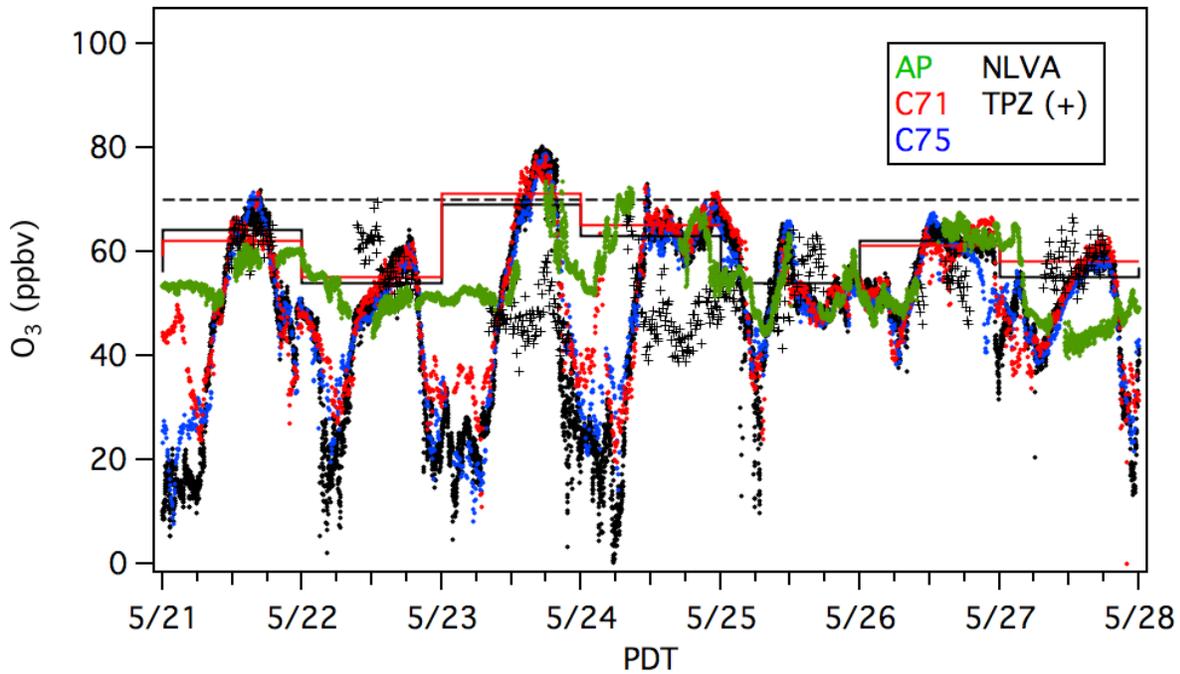
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Figure 9-7. Time-height curtain plots of the FLEXPART (a) STO3, (b) ASCO, and (c) BBCO tracer distributions above the LVV during Week 2 of FAST-LVOS.

1 The TOPAZ curtain plot also shows high O₃ in the boundary layer on May 23 and the Walter
 2 Johnson monitor recorded the first NAAQS exceedance of *FAST-LVOS* (71 ppbv) on that day
 3 with the nearby Palo Verde (68 ppbv) and Joe Neal (69 ppbv) monitors following close behind.
 4 The Paul Meyer and JD Smith monitors also measured MDA8 O₃ in excess of 60 ppbv, but the
 5 outlying monitors (e.g. Jean, Indian Springs, and Mesquite) only reported about 50 ppbv. Ozone
 6 was more uniformly distributed on May 24 when all of the Clark County monitors recorded
 7 MDA8 O₃ mixing ratios between 63 and 66 ppbv.
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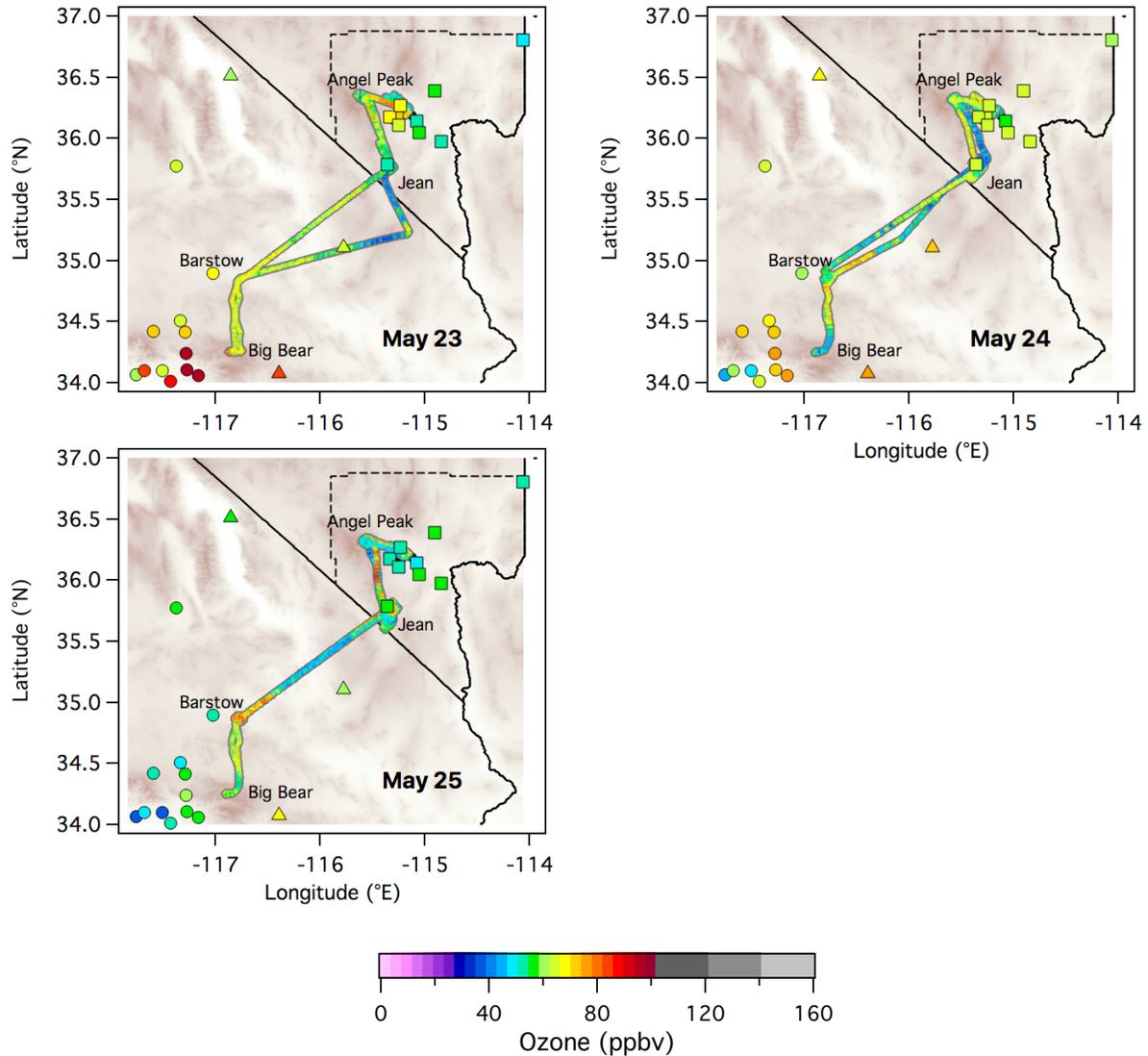


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 10 **Figure 9-8.** Time series of the TOPAZ O₃ mixing ratios at 4.0 km agl (black +) during Week 2
 11 plotted with the 1-min O₃ measurements from the NLVA (black) and AP (green) and the 5-min
 12 measurements from the Walter Johnson (C71, red) and Joe Neal (C75, blue) monitors. The red
 13 and blue steps show the corresponding C71 and C75 MDA8 measurements. The dashed black
 14 line indicates the 70 ppbv NAAQS.

15
 16 The O₃ mixing ratios at 4.0 km were similar to or less than the surface concentrations (**Figure 9-**
 17 **8**). The mixing ratios at AP were similar to those in the LVV, but the Angel Peak record has gaps
 18 during the afternoons of May 23 and 24 when the mobile laboratory descended into the
 19 western LVV.

20
 21 More insight into the high O₃ on May 23 is provided by the Scientific Aviation measurements
 22 which are summarized in **Figures 9-9** through **9-11**. The aircraft distributions above the LVV in
 23 **Figure 9-10** appear very similar to the TOPAZ and ozonesonde profiles, with the highest O₃
 24 measured below 2 km during the landing approach on the afternoon of May 23. The latitudinal
 25 transect from May 23 plotted in **Figure 9-11** shows that this high O₃ was confined to the

1 boundary layer above the LVV and therefore of local origin. The aircraft flight on May 24 shows
2 the more regional distribution also seen in the surface measurements.
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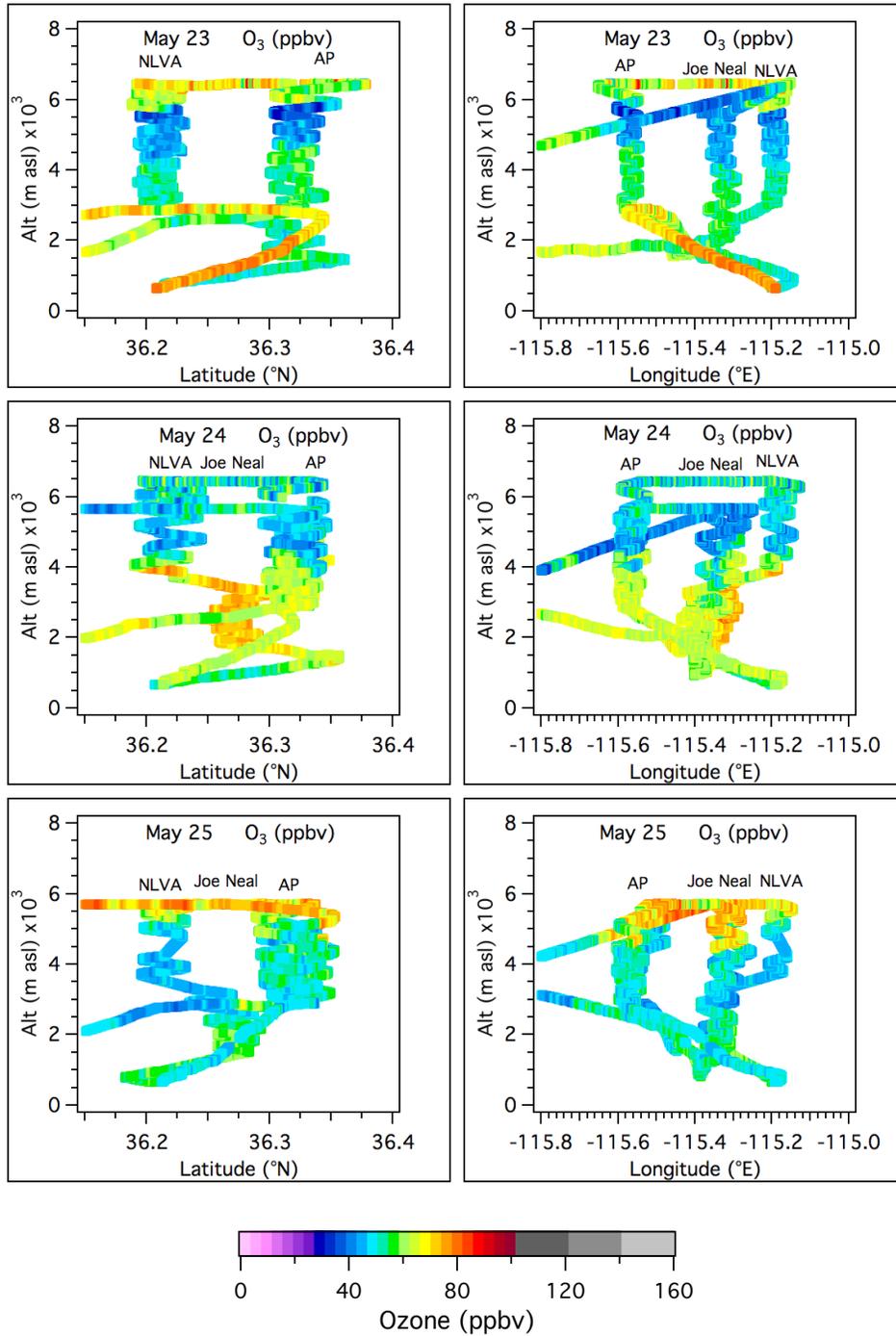


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11 **Figure 9-9.** Flight tracks for the Scientific Aviation TLS Bravo during the first FAST-LVOS IOP (May
12 23-25). The filled squares, circles, and triangles mark the locations of surface monitors operated
13 by CCDAQ, CARB, and the USNPS, respectively. The flight tracks are colored to show the in situ
14 O₃ and the symbols colored to show the reported MDA8 O₃.
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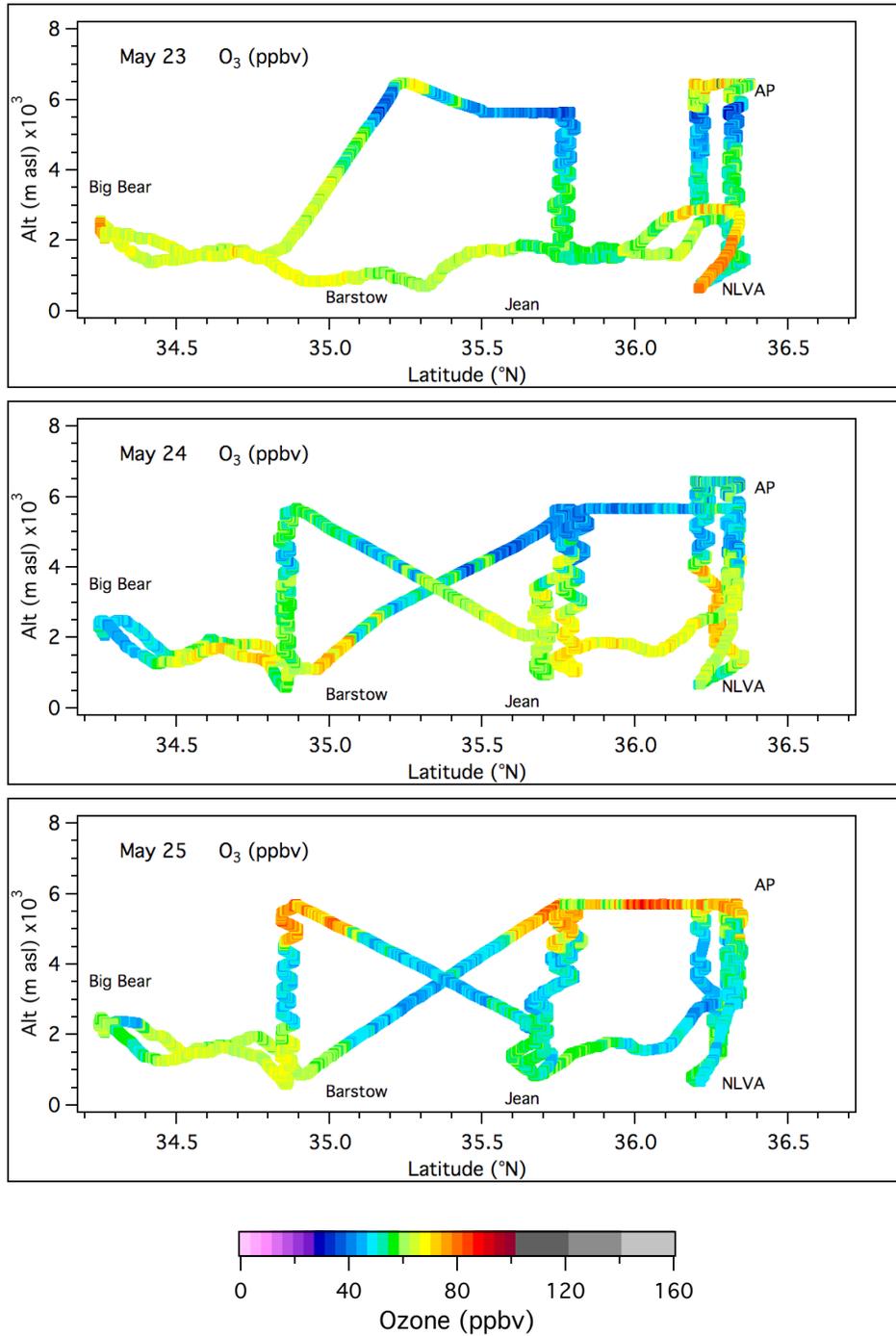


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7 **Figure 9-10.** Scientific Aviation O₃ profiles above the NLVA, Joe Neal, and Angel Peak during the
8 first IOP.

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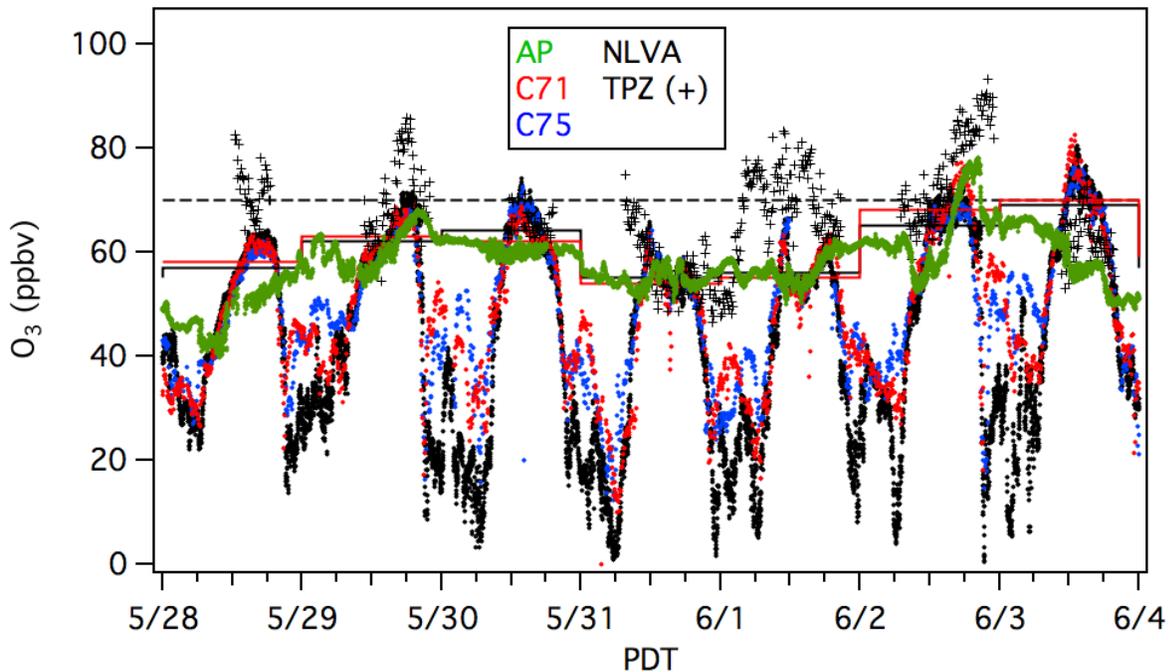
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Figure 9-11. Ozone mixing ratios measured along the flights from NLVA to Big Bear, CA and back during the first IOP.

1 **9.3 Week 3: May 28-June 3 (IOP2)**

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3 The third week of FAST-LVOS featured slightly above average temperatures. The highs at the
4 KVGT tower ranged between 34.4 and 38.9°C (94 and 102°F). The synoptic forecasts for the end
5 of May showed the approach of another upper level low and the possibility of a tropopause
6 fold. The second IOP began on May 31 with ozonesondes launched at 0900 and 1200 PDT.
7 Three more ozonesondes were launched on both June 1 and 2 at around 0900, 1200, and 1500
8 PDT. Scientific Aviation conducted research flights on all three days.
9

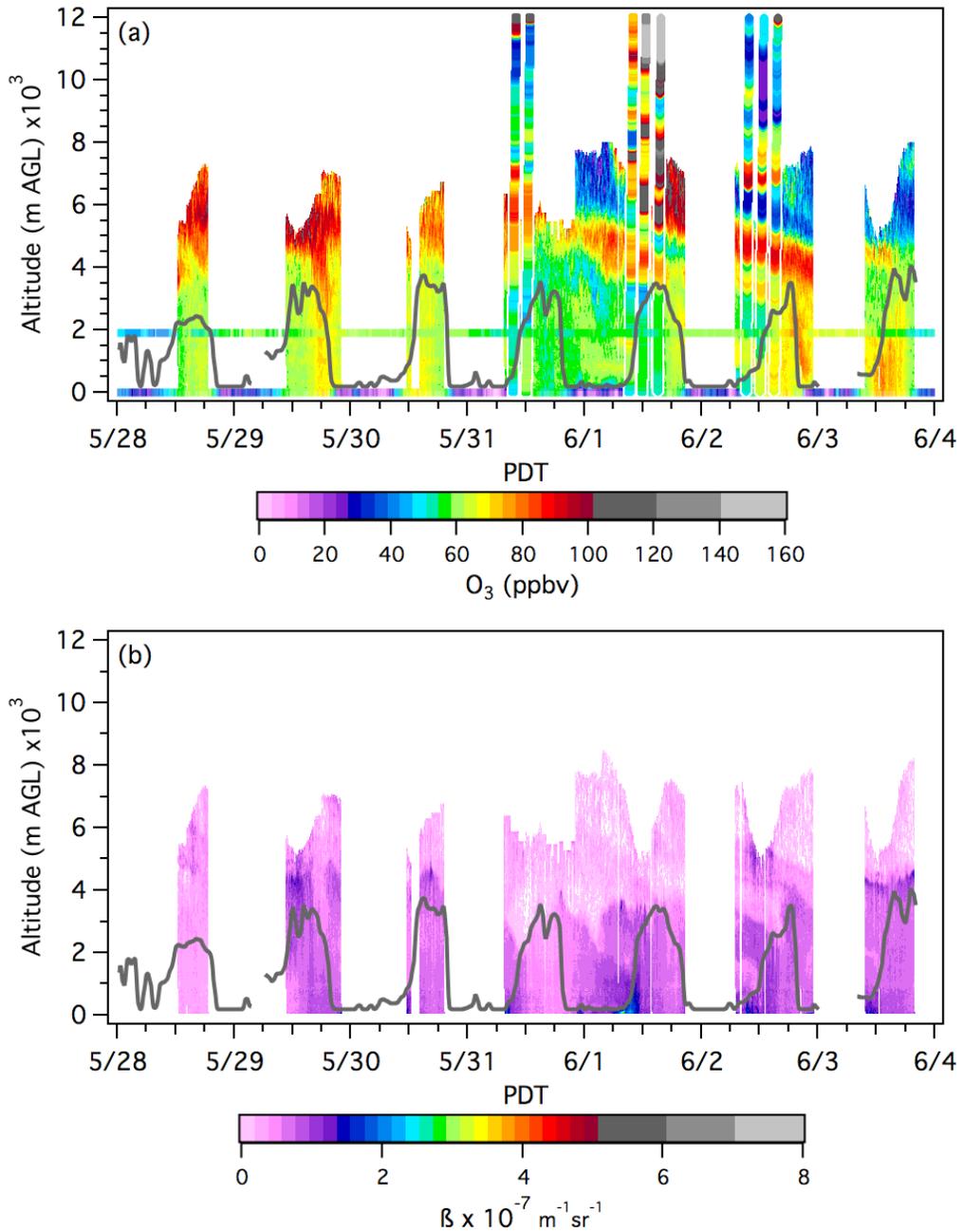
10 The surface measurements plotted in **Figure 9-12** show elevated O₃, but no exceedances at the
11 LVV surface monitors. The Walter Johnson and Joe Neal monitors recorded MDA8 O₃
12 concentrations of 68 and 65 ppbv on June 2, and 70 and 69 ppbv on June 3. TOPAZ measured
13 higher O₃ at 4 km on those days as well as on May 28-29, suggesting that there may have been
14 some transport to the surface.
15



16 **Figure 9-12.** Time series of the TOPAZ O₃ mixing ratios at 4.0 km agl (black +) during Week 3
17 plotted with the 1-min O₃ measurements from the NLVA (black) and AP (green) and the 5-min
18 measurements from the Walter Johnson (C71, red) and Joe Neal (C75, blue) monitors. The red
19 and blue steps show the corresponding C71 and C75 MDA8 measurements. The dashed black
20 line indicates the 70 ppbv NAAQS.
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23 The high O₃ aloft is more easily seen in the curtain plot of **Figure 9-13**. The lower panel **also**
24 shows higher backscatter than was observed in the previous weeks that is sometimes
25 correlated with and sometimes anticorrelated with O₃. The curtain plots appear to show
26 entrainment of O₃ from above on the afternoons of May 29 and June 2.

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Figure 9-13. Time-height curtain plots of the TOPAZ (a) ozone, and (b) backscatter, measured during Week 3 of FAST-LVOS. The ascending Joe Neal ozonesonde profiles from IOP2 are superimposed. The colored horizontal bands in (a) show the in-situ measurements from the NLVA and AP. The black line shows the local mixed-layer height from the micro-Doppler lidar.

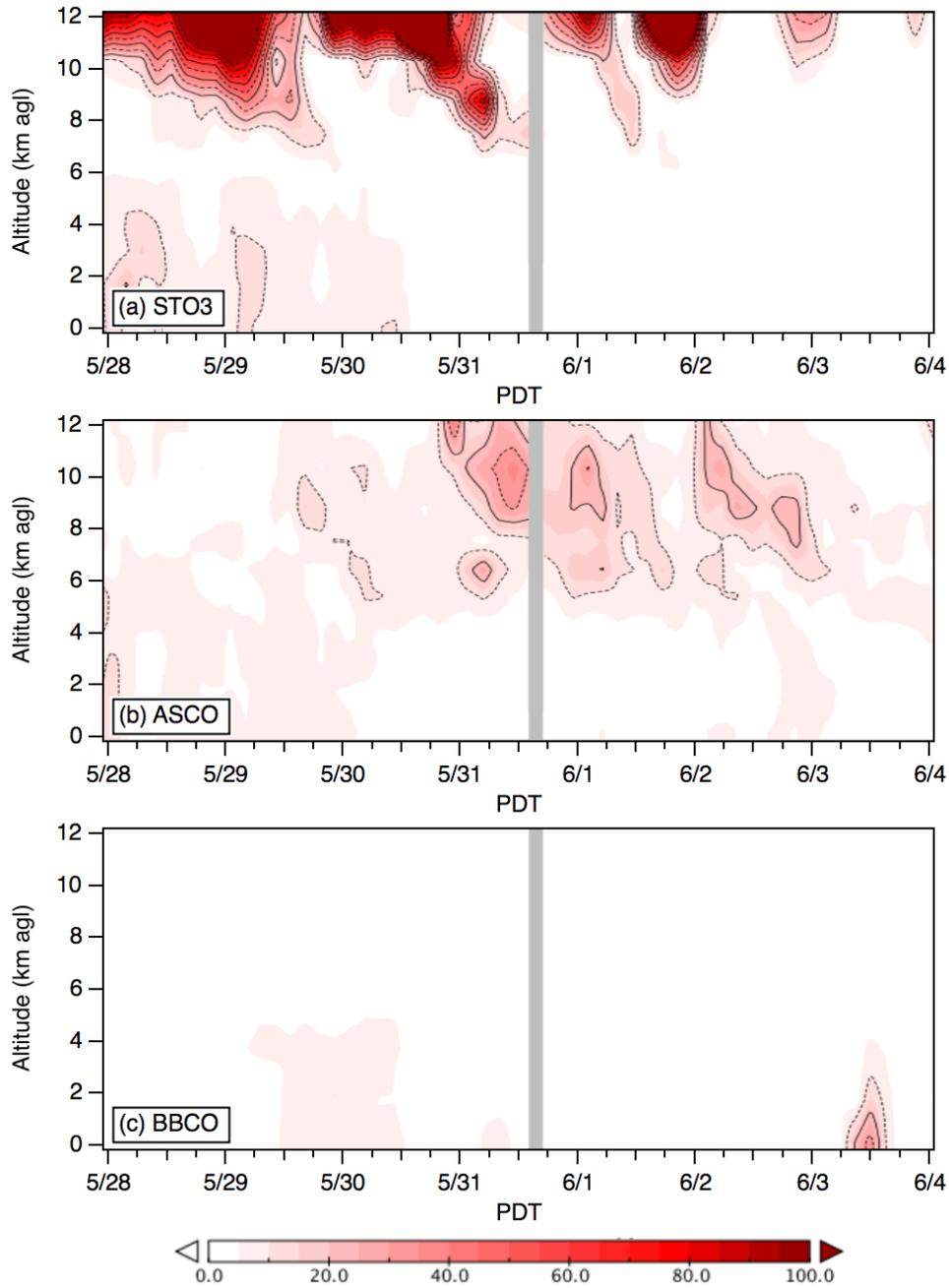
1 The FLEXPART curtain plots in **Figure 9-14** show multiple tongues of STO3 tracer descending
2 into the upper troposphere that were followed several days later by filaments of the ASCO
3 tracer (cf. **Figures 8-2** and **8-3**). Smaller concentrations of both tracers reached the surface on
4 May 28-30, and some ASCO tracer appears near the surface on June 2. The bottom panel also
5 shows small enhancements in the BBCO tracer near the surface on May 30 and June 3.

6
7 **Figure 9-15** displays enlarged curtain plots for the TOPAZ measurements from June 2 and 3,
8 which show the complex relationship between O_3 and β on these days more clearly. The NLVA
9 and AP in-situ measurements and June 2 ozonesonde profiles are superimposed as before. The
10 low-lying residual layer with about 70 ppbv of O_3 seen just above the top of the boundary layer
11 on the morning of June 2 has relatively low aerosol loading, as does the thin layer with about 80
12 ppbv of O_3 that slopes downward from 4 km to the top of the afternoon mixed layer. The
13 intervening layer with less than 60 ppbv of O_3 has higher aerosol, however, and the ozonesonde
14 profiles in **Figure 9-16** show that this layer also has high relative humidity. The highly laminar
15 band with about 100 ppbv of O_3 that slopes downward from about 5 to 4 km also has high
16 aerosol, but low relative humidity. The relatively high aerosol loading suggests that these layers
17 were caused by Asian pollution or Siberian wildfires and not stratospheric intrusions or North
18 American biomass burning plumes. This interpretation is supported by the FLEXPART tracer
19 distributions in **Figure 9-14** (the FLEXPART biomass burning tracer only includes North American
20 sources). The low relative humidity seen in the ozonesonde profiles suggests that the pollution
21 was transported across the Pacific in the upper in the upper troposphere. The regular
22 undulations seen in the ozone and relative humidity profiles and parallel sloping bands in the
23 ozone and backscatter curtain plots on June 2 are probably caused by horizontally propagating
24 gravity waves [*Langford et al.*, 1996] that may have contributed to the downward transport of
25 the Asian pollution.

26
27 The Scientific Aviation flights above the LVV and Angel Peak during IOP2 provide additional
28 evidence that the elevated ozone was primarily transported Asian pollution. **Figures 9-17** and **9-**
29 **18** show that CH_4 was also elevated in the 5 km layer with high O_3 , β , and low RH.

30
31 The Angel Peak in-situ measurements plotted in **Figure 9-19** and **9-20** show that O_3 and CO
32 were positively correlated during most of the week. The highest O_3 was measured on the
33 evening of June 2 after a period of southeasterly upslope flow, and was accompanied by
34 increases in CO, NO_2 , and NO_y . The peak in the time series coincides with the rapid descent of
35 the O_3 from the transport layer as the mixed layer subsided (**Figure 9-15**). The curtain plot
36 suggests that this high O_3 remained in the residual layer above the LVV and was entrained the
37 following morning to contribute to the high surface O_3 on June 3.

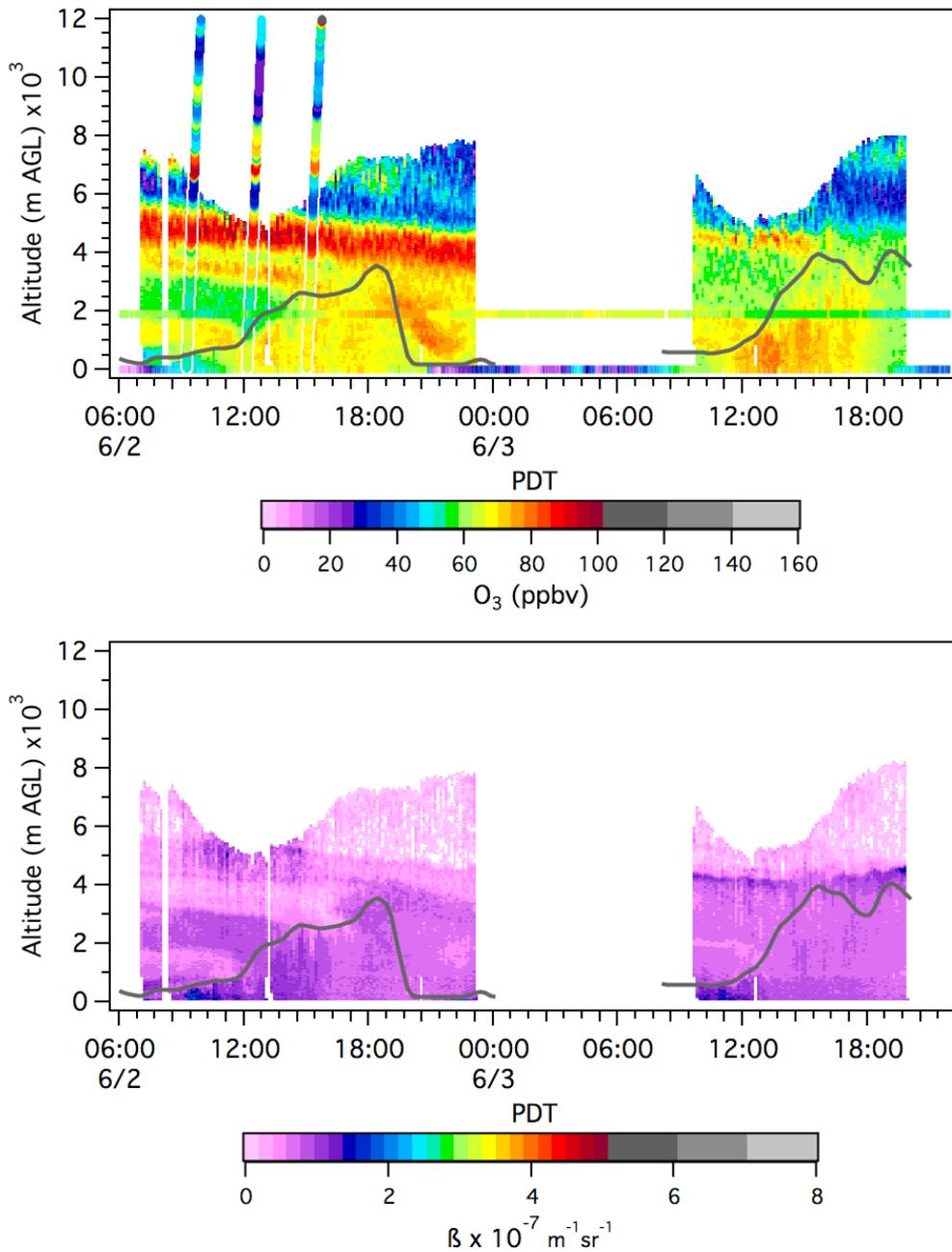
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Figure 9-14. Time-height curtain plots of the FLEXPART (a) STO3, (b) ASCO, and (c) BBCO tracer distributions above the LVV during Week 3 of FAST-LVOS.

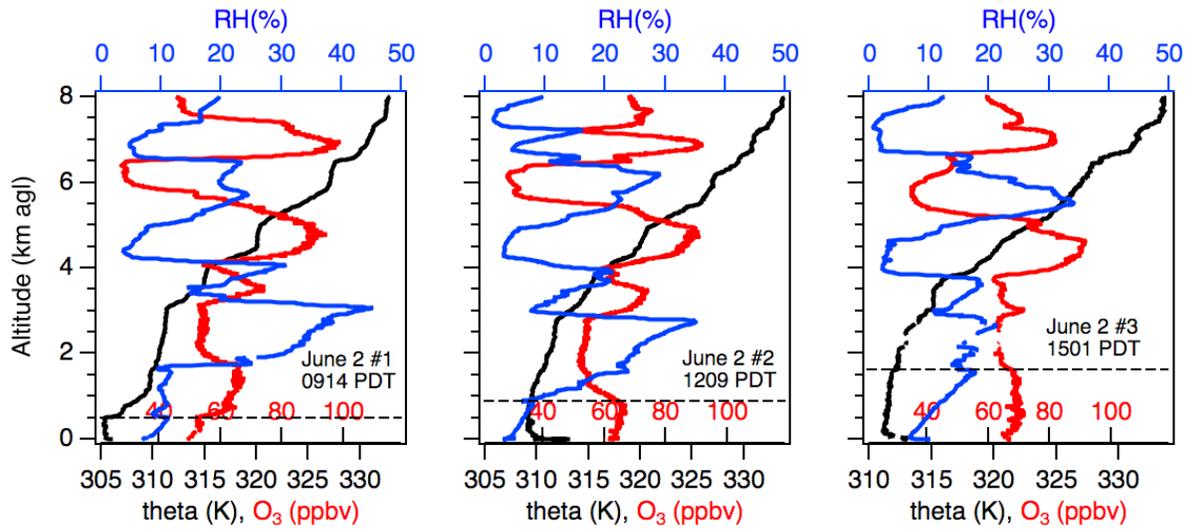
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Figure 9-15. Time-height curtain plots of the TOPAZ (a) ozone, and (b) backscatter, measured on June 2-3. The ascending Joe Neal ozonesonde profiles from IOP2 are superimposed. The colored horizontal bands in (a) show the in-situ measurements from the NLVA and AP. The black line shows the local mixed-layer height from the micro-Doppler lidar.

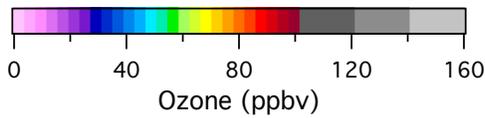
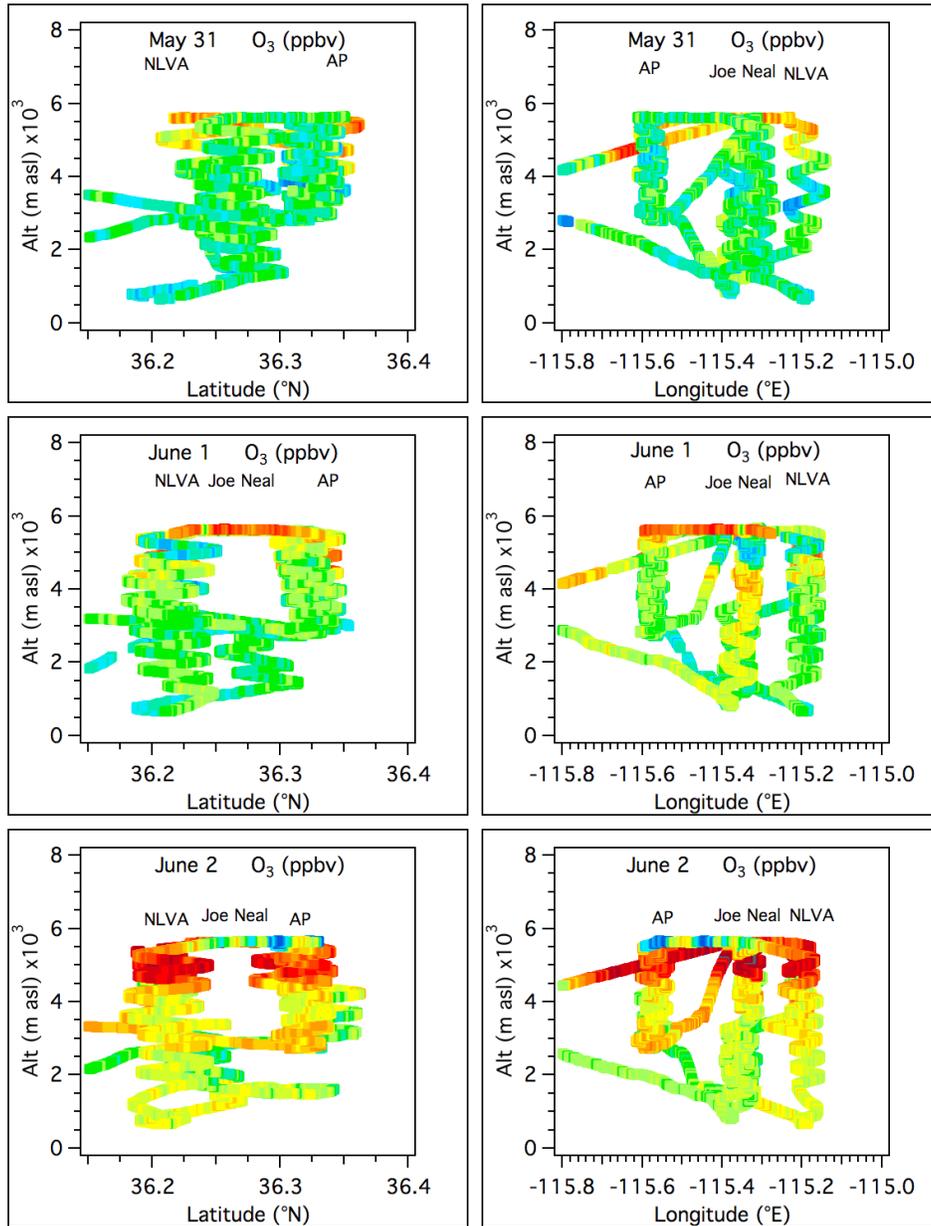
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Figure 9-16. Profiles of potential temperature (black), O₃ (red), and relative humidity (blue) from the Joe Neal ozonesondes on June 2. The dashed black horizontal line marks the top of the boundary layer.

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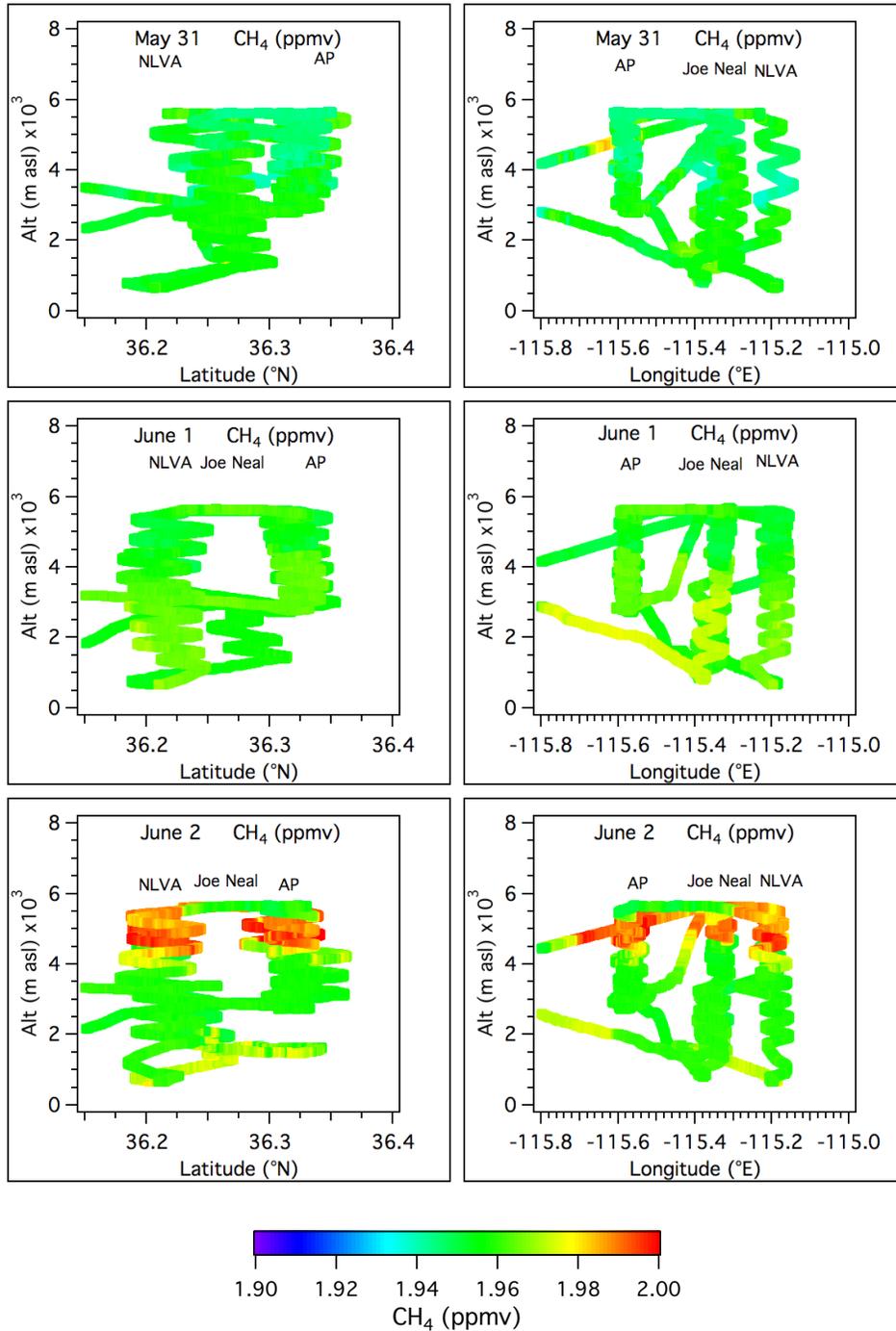
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8 **Figure 9-17.** Scientific Aviation O₃ profiles above the NLVA, Joe Neal, and Angel Peak during the
9 second IOP.

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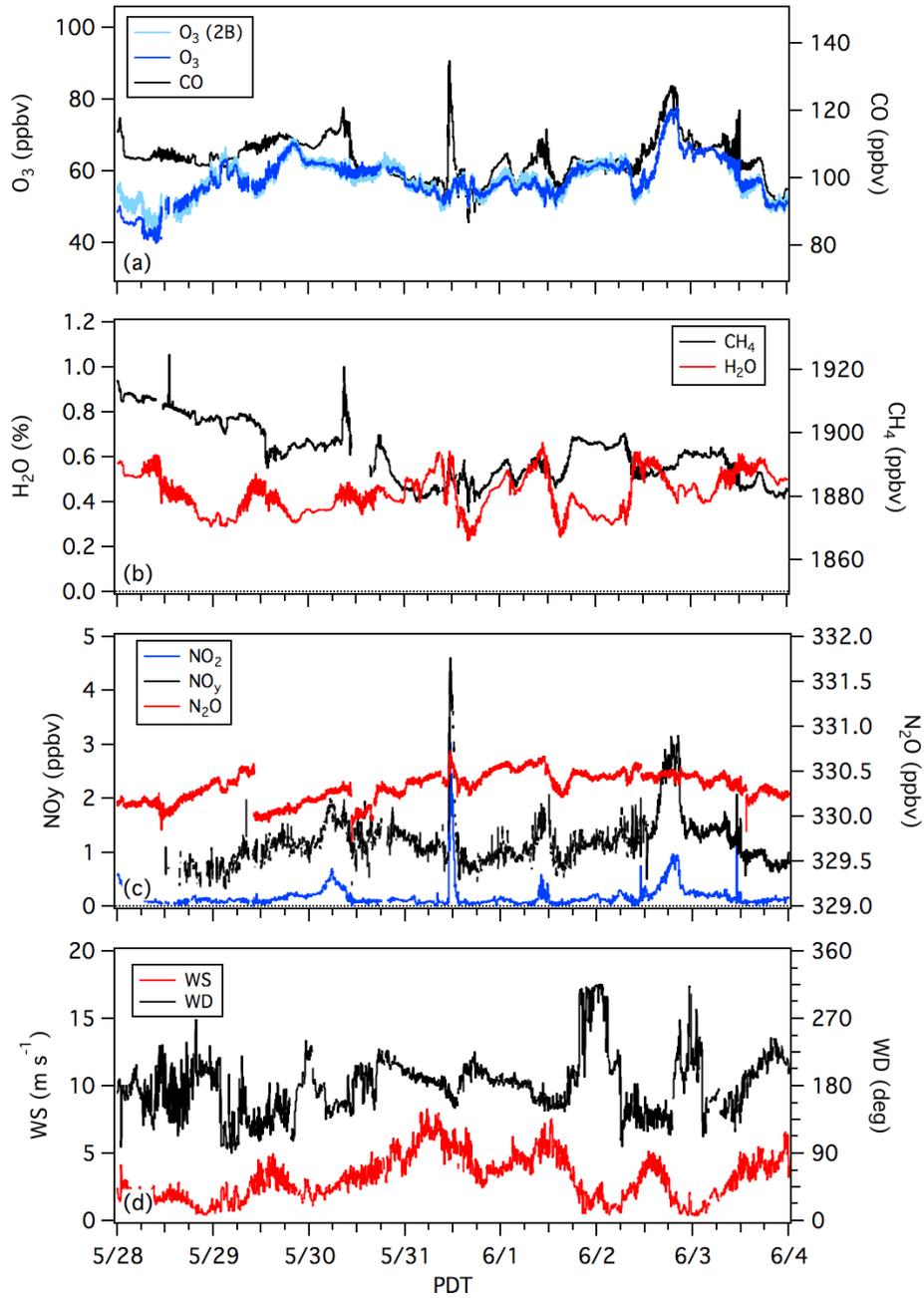
7 **Figure 9-18.** Scientific Aviation CH₄ profiles above the NLVA, Joe Neal, and Angel Peak during the
8 second IOP.

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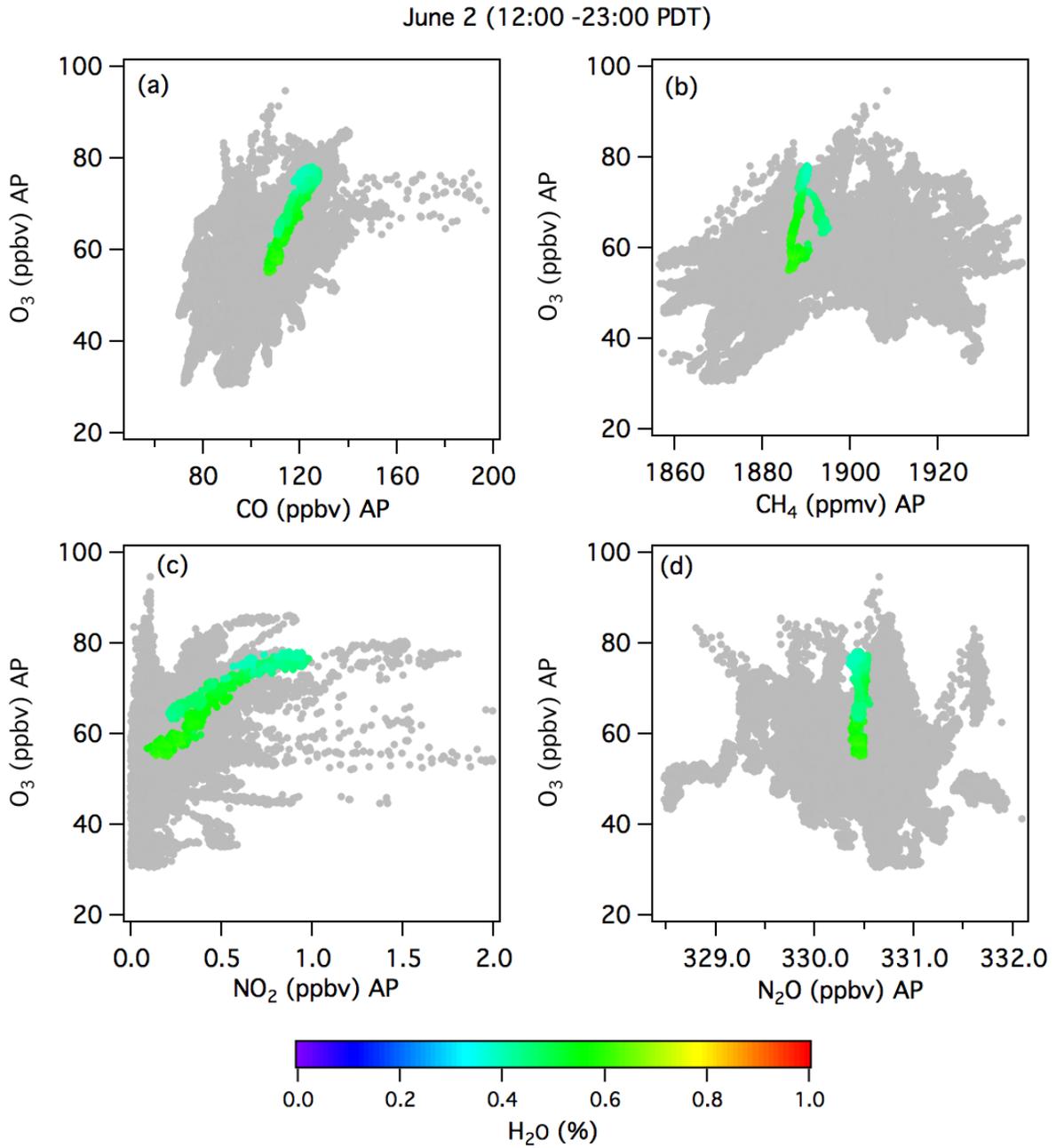
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Figure 9-19. Time series of (a) O_3 and CO, (b) H_2O and CH_4 , (c) NO_x , NO_y and N_2O , and (d) wind speed and direction measured by the mobile laboratory on Angel Peak during Week 3 of FAST-LVOS. The narrow CO and NO_y spikes near noon on May 31 are probably of local origin.

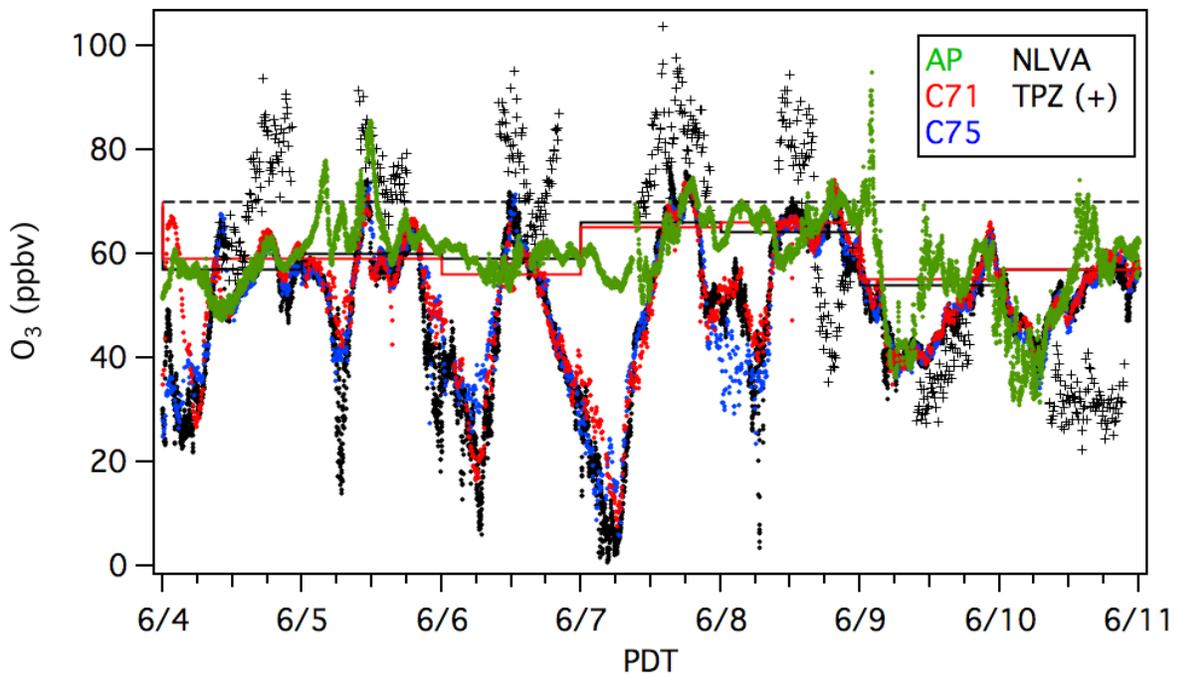
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6 **Figure 9-20.** Scatter plots showing the correlations between in-situ O₃ and (a) CO, (b) N₂O, (c)
7 CO₂, and (d) CH₄ measurements from the mobile lab on Angel Peak. The measurements from the
8 afternoon and evening (12-23 PDT) of June 2 are colored by the water vapor measurements.
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1 **9.4 Week 4: June 4-10**

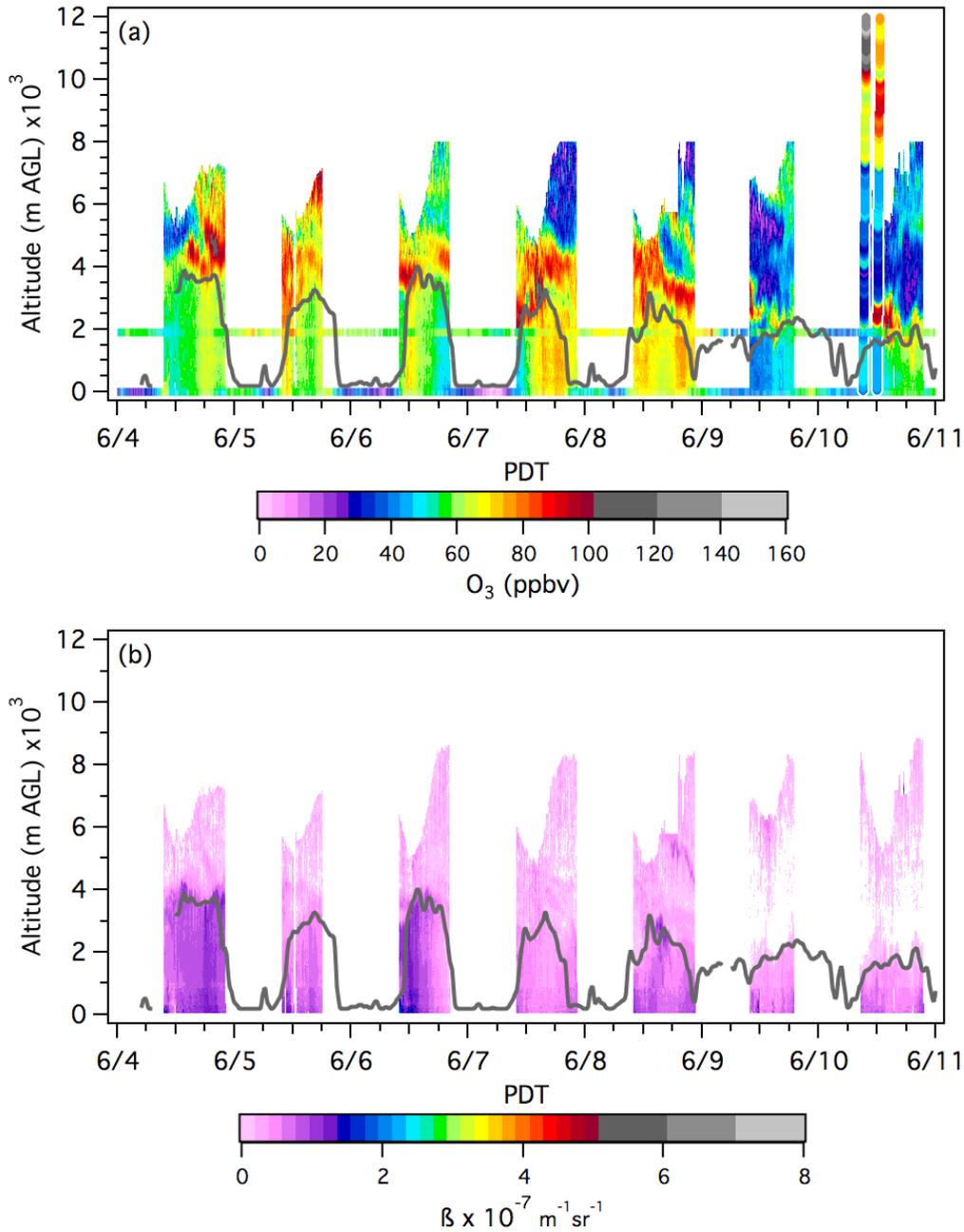
2
3 The second week of June started out very hot with the daily highs exceeding 37.8°C (100°F)
4 from the 4th to the 8th. The high dropped to 35°C (95°F) on June 9 with the arrival of the cold
5 front preceding the large tropopause fold event of the following week (cf. **Figure 8-2**). June 10
6 marked the first day of IOP3. There were no exceedances of the 2015 NAAQS during Week 4,
7 but **Figure 9-21** shows that the MDA8 O₃ at the Walter Johnson and Joe Neal monitors
8 increased from about 55 ppbv on June 4, to 65 ppbv on June 7 and 8. The concentrations fell to
9 much lower values with the arrival of the cold front. The ozone curtain plot in **Figure 9-22**
10 shows high O₃ around 4 km between June 4 and June 8, and appears to show some of this O₃
11 being entrained into the mixed layer on each day.
12



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14
15 **Figure 9-21.** Time series of the TOPAZ O₃ mixing ratios at 4.0 km agl (black +) during Week 4
16 plotted with the 1-min O₃ measurements from the NLVA (black) and AP (green) and the 5-min
17 measurements from the Walter Johnson (C71, red) and Joe Neal (C75, blue) monitors. The red
18 and blue steps show the corresponding C71 and C75 MDA8 measurements. The dashed black
19 line indicates the 70 ppbv NAAQS.

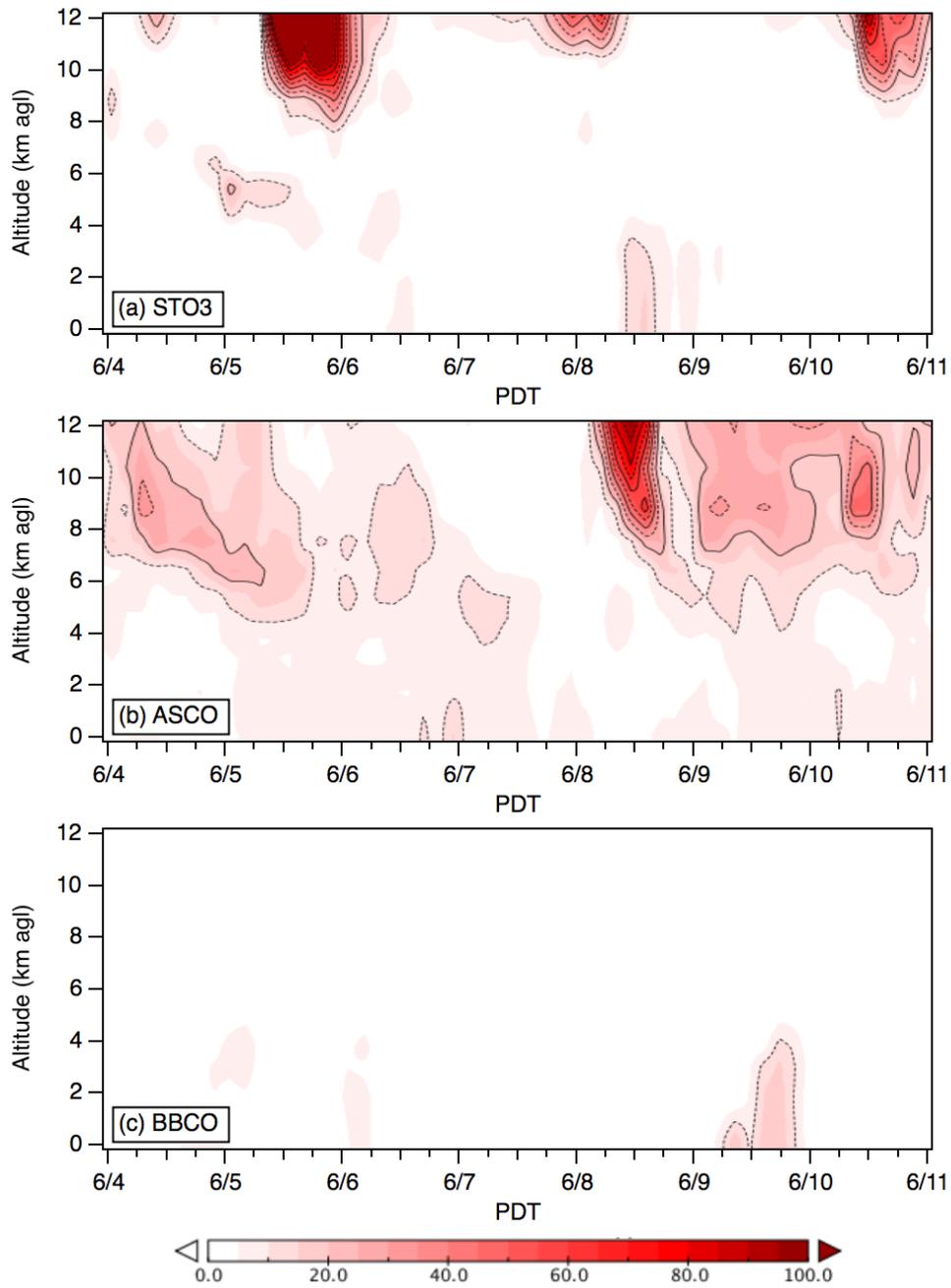
20
21 The FLEXPART curtains in **Figure 9-23** (and map in **Figure 8-3**) show some stratospheric O₃ and
22 large amounts of Asian pollution above Las Vegas during Week 4. Unfortunately, there were no
23 ozonesondes or aircraft sorties during the first part of the week to help establish the source of
24 the high O₃.
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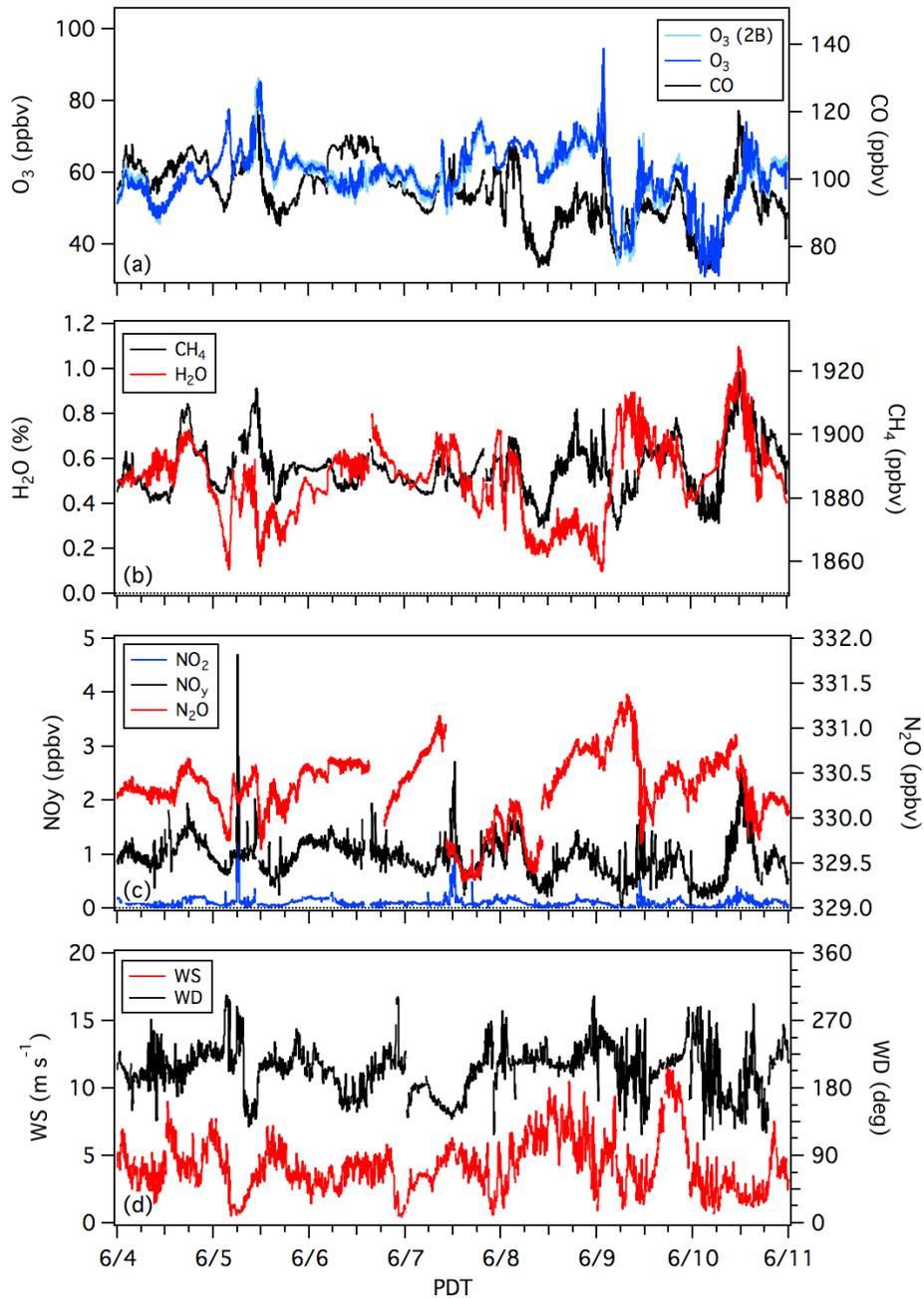
Figure 9-22. Time-height curtain plots of the TOPAZ (a) ozone, and (b) backscatter, measured during Week 4. The ascending Joe Neal ozonesonde profiles from IOP3 are superimposed. The colored horizontal bands in (a) show the in-situ measurements from the NLVA and AP. The black line shows the local mixed-layer height from the micro-Doppler lidar.



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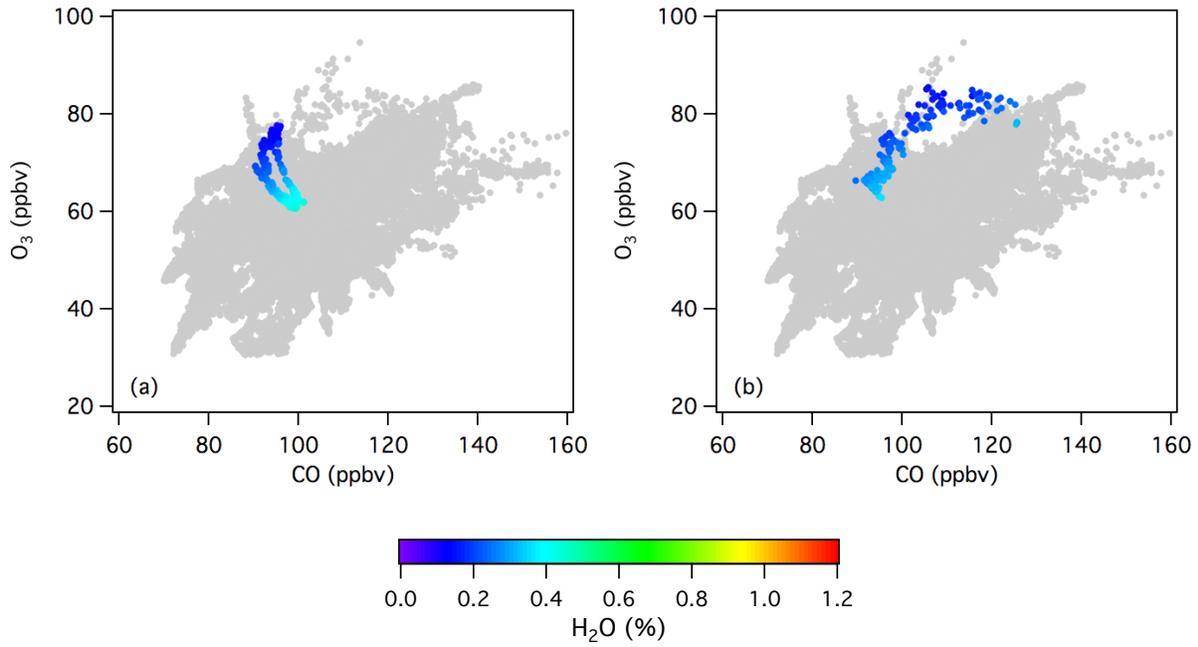
Figure 9-23. Time-height curtain plots of the FLEXPART (a) STO3, (b) ASCO, and (c) BBCO tracer distributions above the LVV during Week 4 of FAST-LVOS.

1 **Figures 9-21 and 9-22** show that some of the O₃ aloft reached the summit of Angel Peak on the
 2 morning of June 5, and the time series in **Figure 9-24** show that the first peak to 77 ppbv at
 3 about 0400 PDT was accompanied by decreases in CO and CH₄, but the second peak of 85 ppbv
 4 around local noon was accompanied first by increased, and then decreased CO and CH₄.
 5



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 8 **Figure 9-24.** Time series of (a) O₃ and CO, (b) H₂O and CH₄, (c) NO_x, NO_y and N₂O, and (d) wind
 9 speed and direction measured by the mobile laboratory on Angel Peak during Week 4 of FAST-
 10 LVOS.

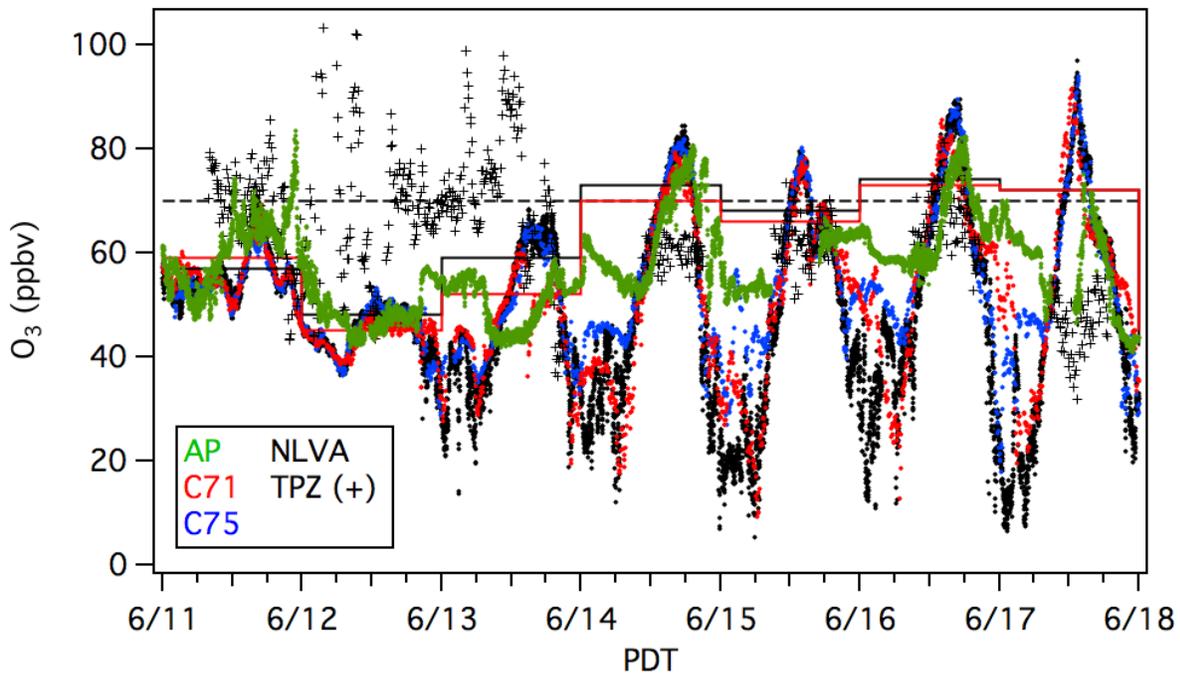
1 These correlations (**Figure 9-25**) show that the descending air was composed of intermingled
2 stratospheric air and Asian pollution. **Figure 9-24a** also shows a narrow O_3 spike around 0200
3 PDT on June 9 that was accompanied by increased CO and CH_4 consistent with Asian pollution.
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8
9 **Figure 9-25.** Scatter plots showing the correlations between in-situ O_3 and CO on Angel Peak.
10 The measurements from (a) 0000 to 0445 PDT, and (b) 1030-1400 PDT on June 5 are colored
11 by the water vapor measurements.
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1 **9.5 Week 5: June 11-17 (IOP3)**

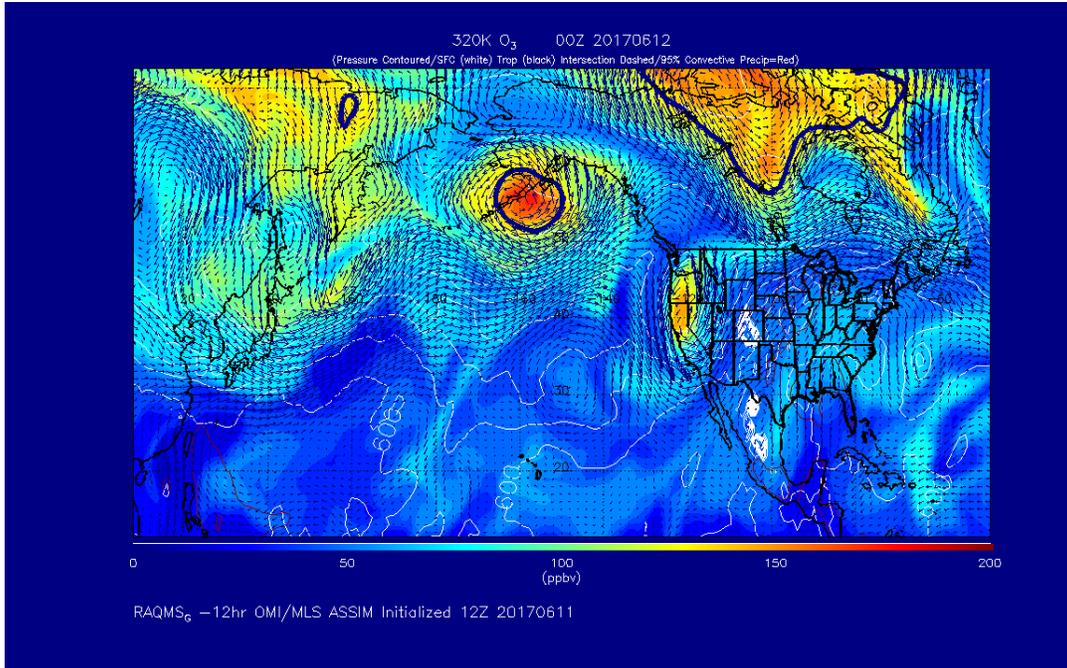
2
3 The arrival of the cold front ahead of the deep upper-level trough (cf. **Figure 6-3**) brought
4 welcome relief from the hot weather of the previous week. The temperatures dropped below
5 freezing at Angel Peak on the morning of on June 12 (cf. **Figure 6-2**) and the daily high
6 temperature reached only 26.7 at the KVG T station with a low of 20.6°C (80 and 69°F,
7 respectively). Sustained southerly winds of up to 15 m s⁻¹ kept the nocturnal boundary layer
8 from forming so that the concentrations did not fall during the night of June 11-12, and **Figure**
9 **9-26** shows that surface ozone fell below 50 ppbv at both the NLVA and AP on June 12 as free
10 tropospheric air descended ahead of the tropopause fold. Ozone remained high at 4 km
11 through the evening of June 13, but then decreased as the trough receded eastward and was
12 followed by a strong ridge and a return to hot conditions. The peak MDA8 concentrations in the
13 LVV increased to values between 68 and 75 ppbv over the next 4 days as the temperatures
14 climbed to a high of 42.8°C (109°F).
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16



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19 **Figure 9-26.** Time series of the TOPAZ O₃ mixing ratios at 4.0 km agl (black +) during Week 5
20 plotted with the 1-min O₃ measurements from the NLVA (black) and AP (green) and the 5-min
21 measurements from the Walter Johnson (C71, red) and Joe Neal (C75, blue) monitors. The red
22 and blue steps show the corresponding C71 and C75 MDA8 measurements. The dashed black
23 line indicates the 70 ppbv NAAQS.
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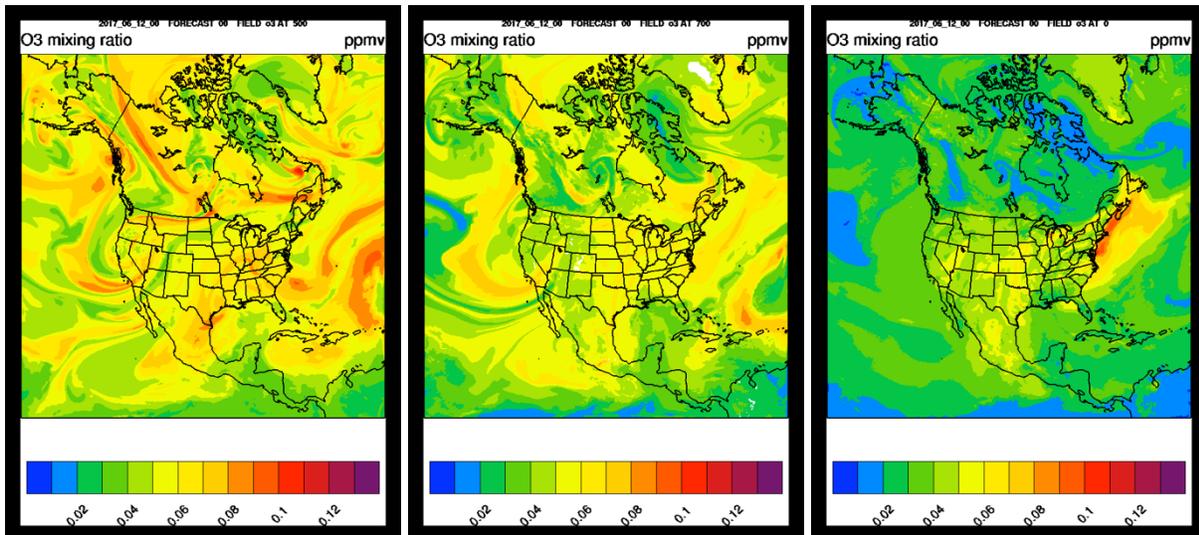
1 The tropopause fold appeared in both the RAQS and RAP-Chem forecasts (**Figures 9-27 and 9-**
2 **28**) and the third IOP was scheduled for June 10-14.

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7 **Figure 9-27.** NOAA NESDIS RAQMS model 320 K O₃ forecast for June 12 00UT. Forecast
8 initialized June 11 12UT.

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13 **Figure 9-28.** RAP-Chem total ozone analyses for 00UT on June 12 at (left to right) 500 hPa, 700
14 hPa, and the surface. Concentrations are given in ppmv.

15

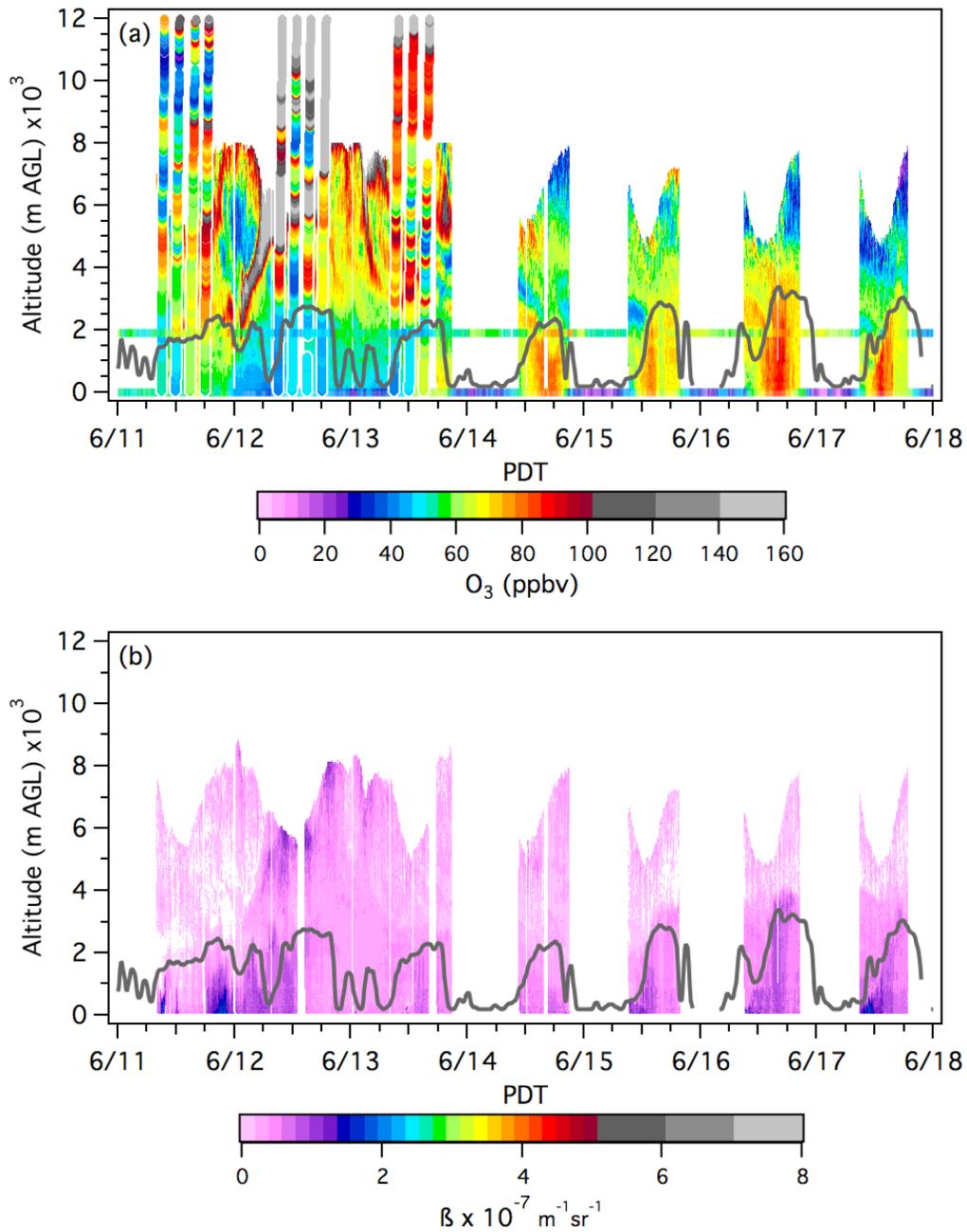
1 The TOPAZ measurements and June 11-13 ozonesondes in **Figure 9-29** show a great deal of fine
2 structure in the ozone measurements from the continuous 60-hour lidar run lasting from the
3 morning of June 11 through the evening of June 13. Both the lidar and one of the Joe Neal
4 ozonesondes measured O₃ mixing ratios in excess of 200 ppbv around 6 km agl at 0900 PDT on
5 June 12. The TOPAZ O₃ curtains also appear to show entrainment from aloft on June 14 and 16.
6

7 The FLEXPART STO3 curtain in **Figure 9-30** shows the descent of more than 30 ppbv of
8 stratospheric ozone to the surface on the morning of June 12 (cf. **Figure 8-2**) with smaller fluxes
9 through the night of the 13th. **Figure 9-30** also shows significant enhancements in the ASCO
10 tracer throughout the week.
11

12 The Angel Peak measurements in **Figures 9-31** and **9-32** show a short-lived spike in O₃ to more
13 than 80 ppbv around midnight on June 11-12, with corresponding dips in CO, CH₄, N₂O, and
14 H₂O. This appearance of this stratospheric filament coincided with the lowest descent of the
15 gray high O₃ tongue in the **Figure 9-29** curtain plot. Except for this brief occurrence, the surface
16 O₃ at and in the valley remained low during the passage of the tropopause fold. This seemingly
17 paradoxical occurrence reflects the fact that the deep descent of tropopause fold can cause a
18 local *decrease* in surface O₃ if the previously polluted boundary layer is fumigated by
19 descending free tropospheric air before it is replaced by a mixture of upper tropospheric and
20 lower stratospheric air with 45-50 ppbv of O₃ [*Langford et al., 2012*]. The stratospheric air does
21 not simply add more ozone to that which is already present. The convective entrainment of
22 quasi-horizontal filaments originating from tropopause folds that form farther upstream is
23 much more likely to cause an O₃ exceedance since the entrained O₃ *does* add to the local
24 pollution. The tropopause fold may have significantly contributed to the widespread
25 exceedances reported in Arizona on June 13-14, but probably had a relatively small impact on
26 Clark County.
27

28 The Angel Peak measurements from June 14, and to a lesser extent those from the following 3
29 days, show large afternoon increases in O₃, CO, CH₄, and NO_y that appear to be caused by
30 upslope flow from the LVV. Scientific Aviation conducted an additional flight on June 14, and
31 the airborne O₃ measurements from the flights on June 13 and 14 are compared in **Figures 9-33**
32 and **9-34**. The measurements from June 13 show highly structured filaments of O₃ trailing
33 behind the receding fold between 4 and 6 km that are consistent with the lidar and ozonesonde
34 measurements. These filaments appear not only above the LVV, but also above the Mojave
35 Desert between the LVV and Big Bear. The measurements from June 14 appear quite different,
36 however, with high O₃ in the mixed layer above both the NLVA and Barstow, but much lower
37 concentrations above Jean and the Mojave Desert that lies between them. These observations,
38 together with those from Angel Peak, imply that 20-25 ppbv of the surface O₃ in the LVV on
39 June 14 was formed by local photochemistry. High O₃ spread across much of the Southwest on
40 June 16 and 17 as stagnation developed, and the MDA8 concentrations measured by the Joe
41 Neal monitor on those days (74 and 72 ppbv) were only ≈10-15 ppbv higher than those
42 measured at Jean (63 and 63 ppbv) and the Grand Canyon National Park (59 and 60 ppbv).
43

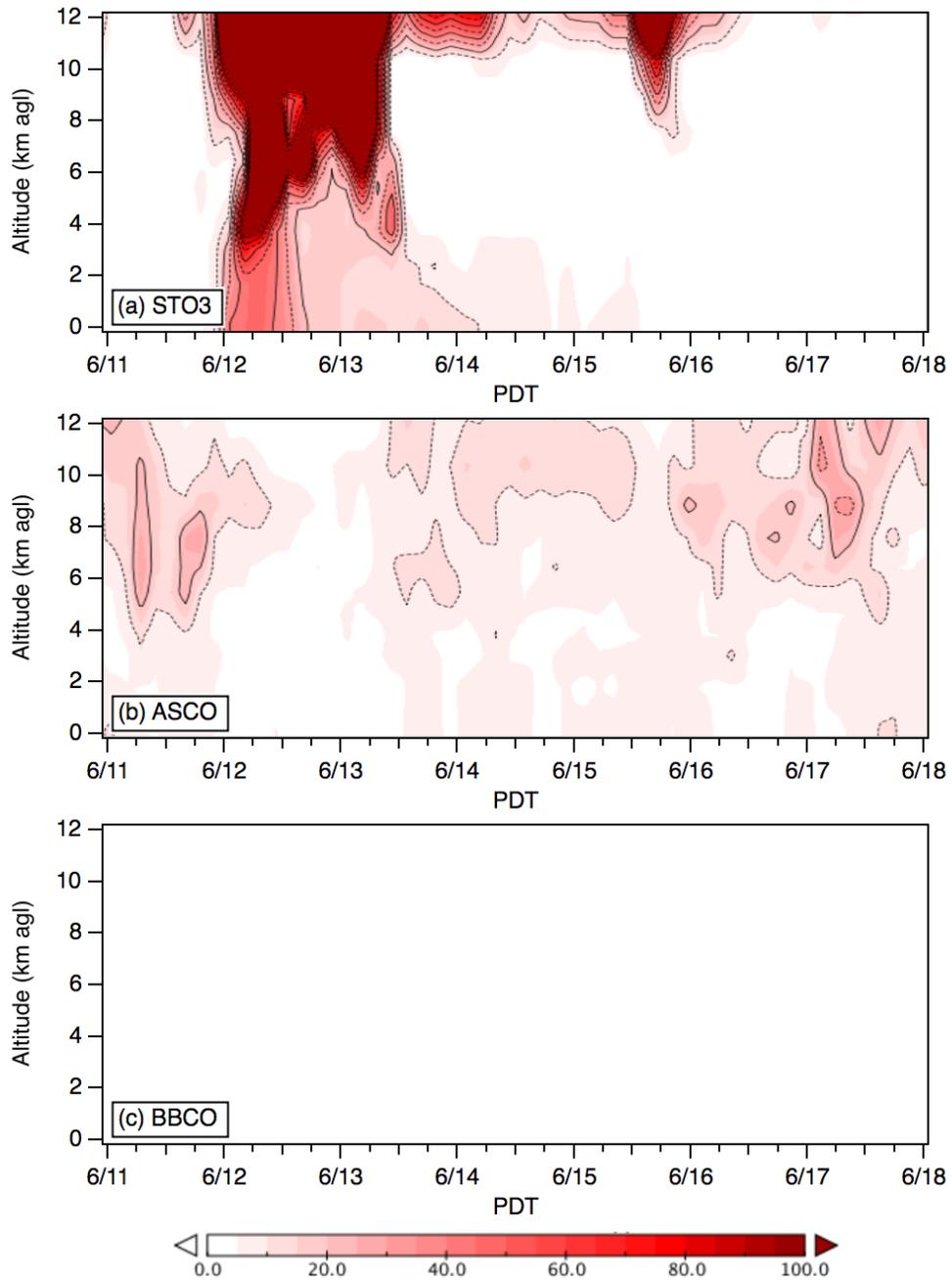
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Figure 9-29. Time-height curtain plots of the TOPAZ (a) ozone, and (b) backscatter, measured during Week 5 of FAST-LVOS. The ascending Joe Neal ozonesonde profiles from IOP3 are superimposed. The colored horizontal bands in (a) show the in-situ measurements from the NLVA and AP. The black line shows the local mixed-layer height from the micro-Doppler lidar. Note that strong winds kept the boundary layer well-mixed throughout the night of June 11-12.

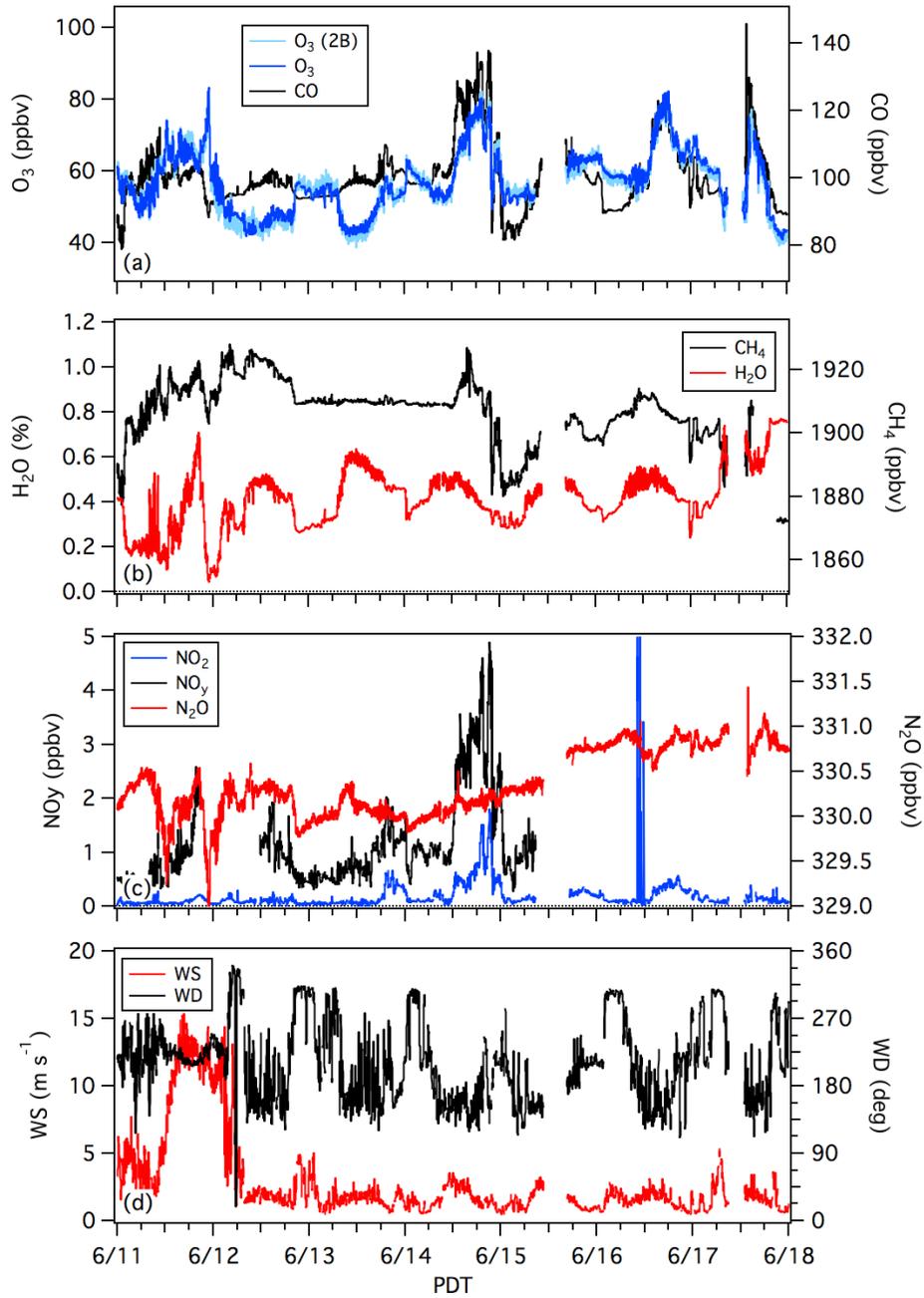
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Figure 9-30. Time-height curtain plots of the FLEXPART (a) STO3, (b) ASCO, and (c) BBCO tracer distributions above the LVV during Week 5 of FAST-LVOS.

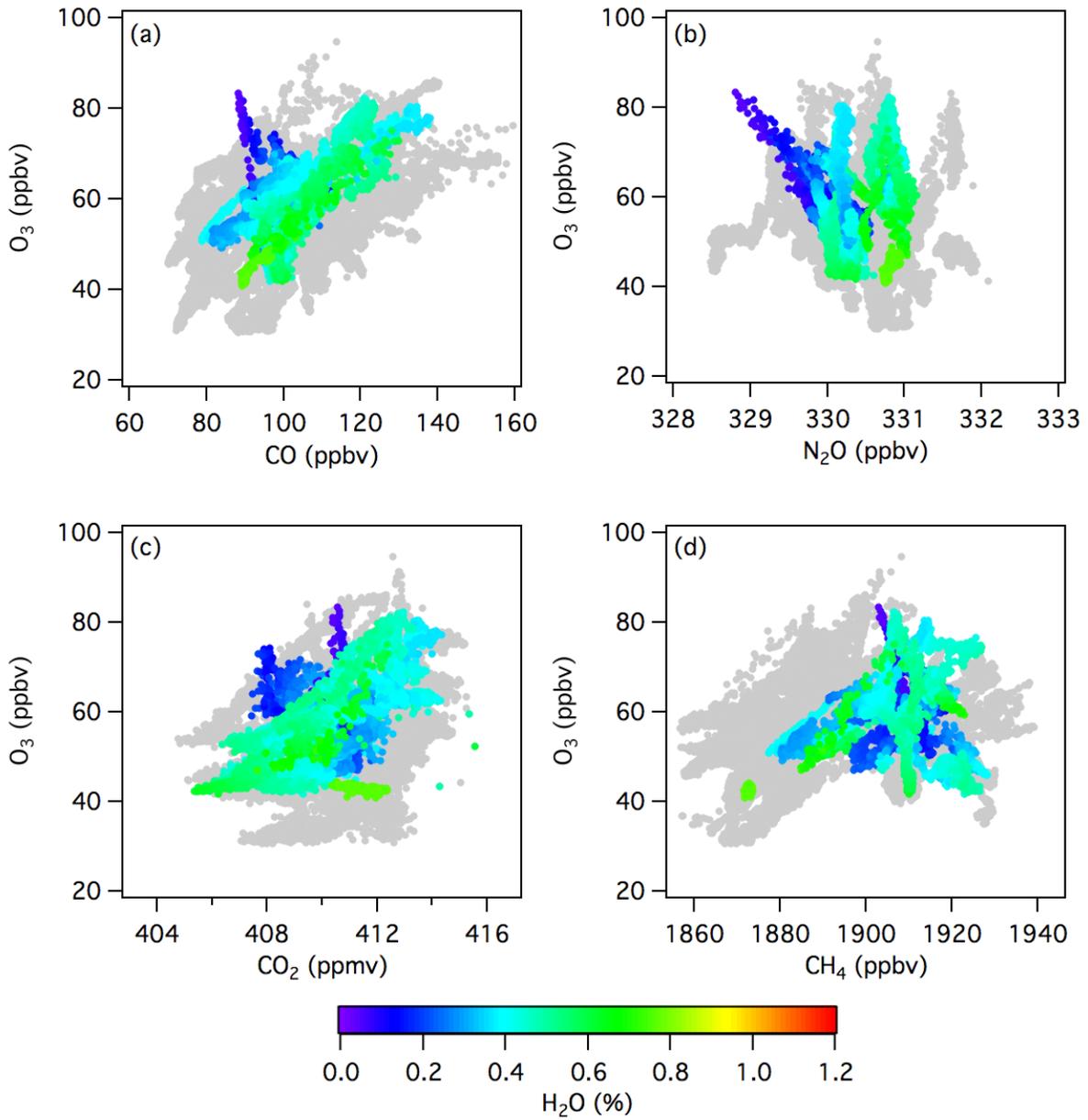
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Figure 9-31. Time series of (a) O₃ and CO, (b) H₂O and CH₄, (c) NO_x, NO_y and N₂O, and (d) wind speed and direction measured by the mobile laboratory on Angel Peak during Week 5 of FAST-LVOS.

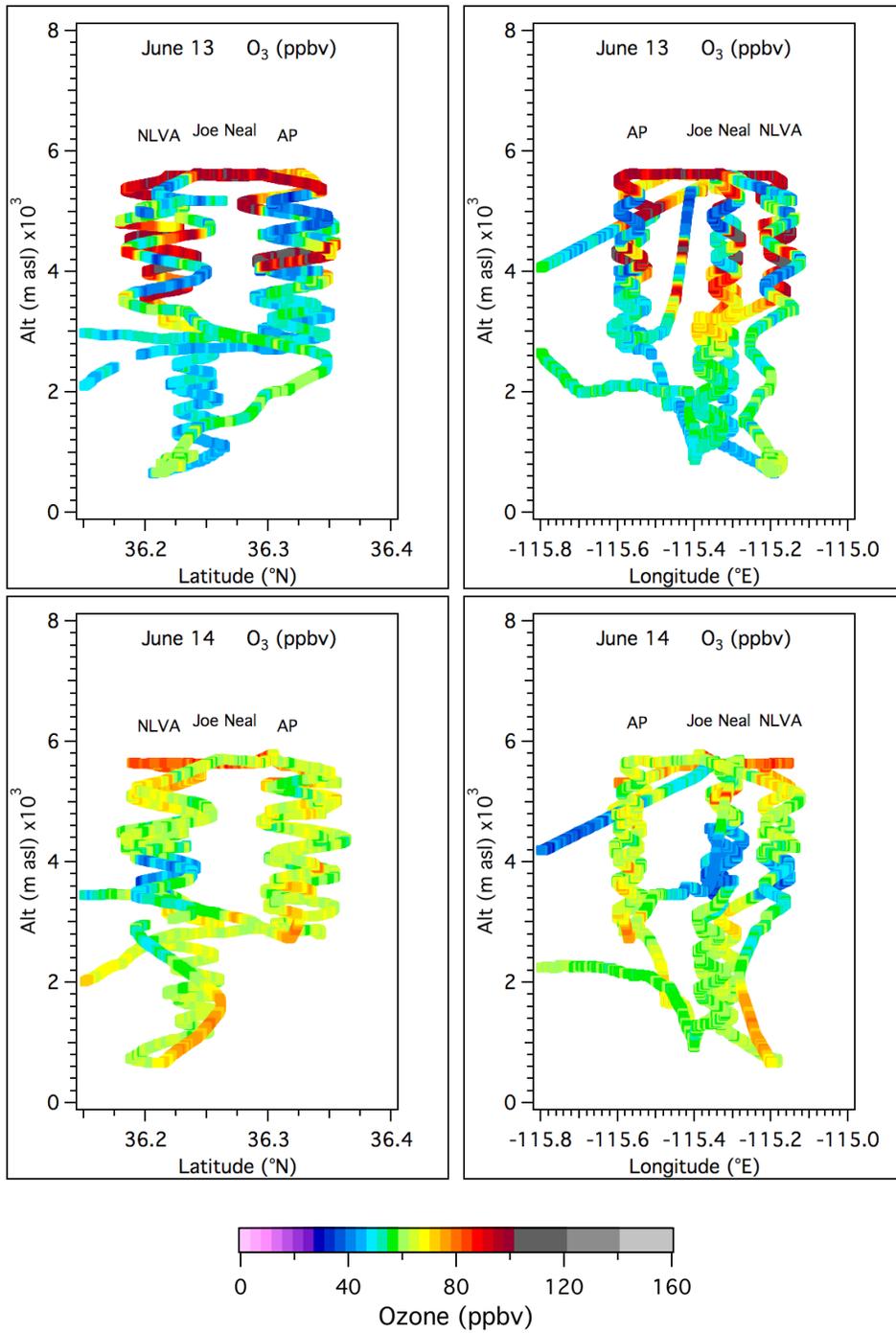
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Figure 9-32. Scatter plots showing the correlations between in-situ O_3 and (a) CO, (b) N_2O , (c) CO_2 , and (d) CH_4 measurements from the mobile lab on Angel Peak. The measurements from Week 5 are colorized by the water vapor measurements.

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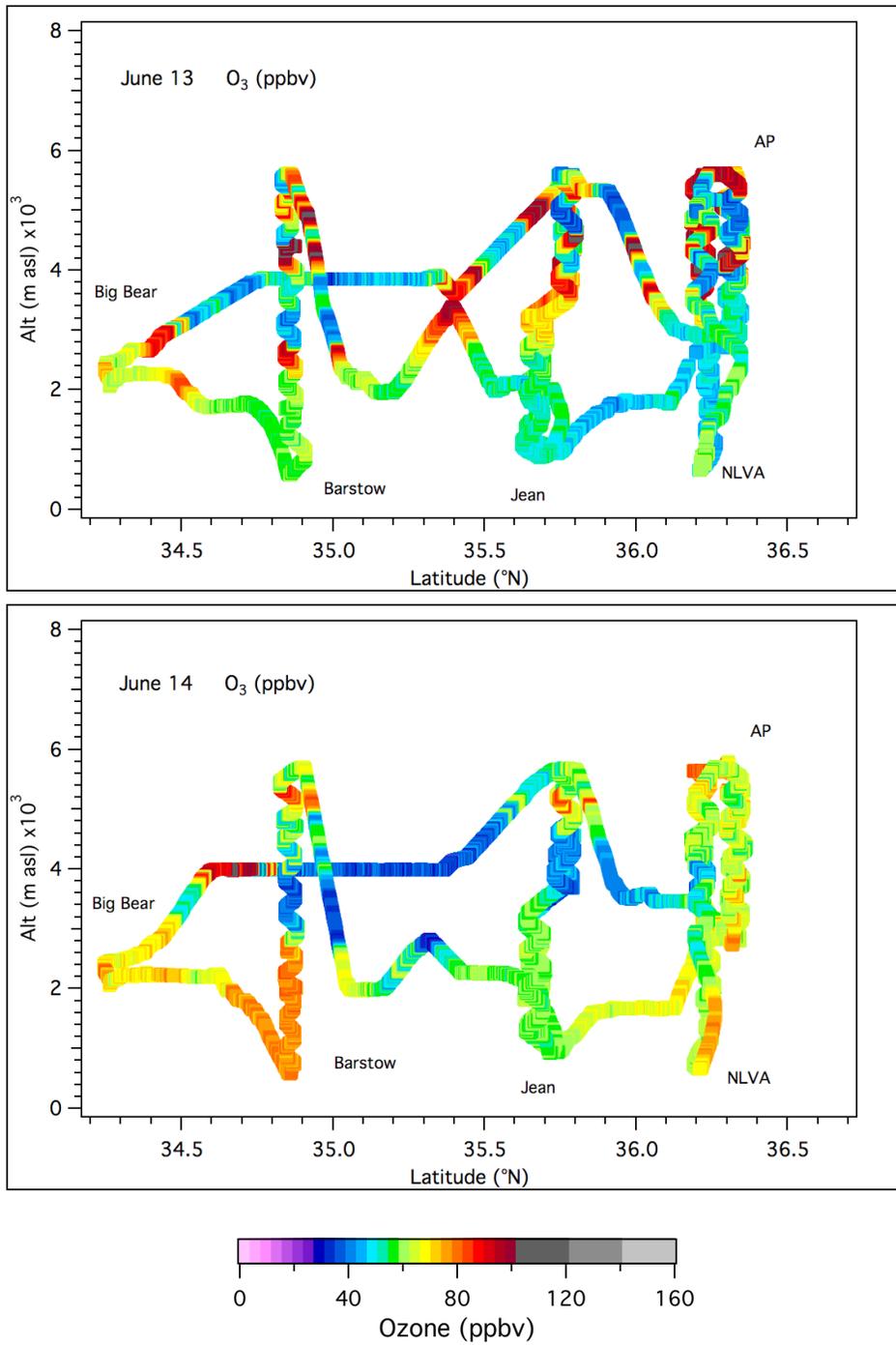
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Figure 9-33. Scientific Aviation O₃ profiles above the NLVA, Joe Neal, and Angel Peak on June 13 and 14.

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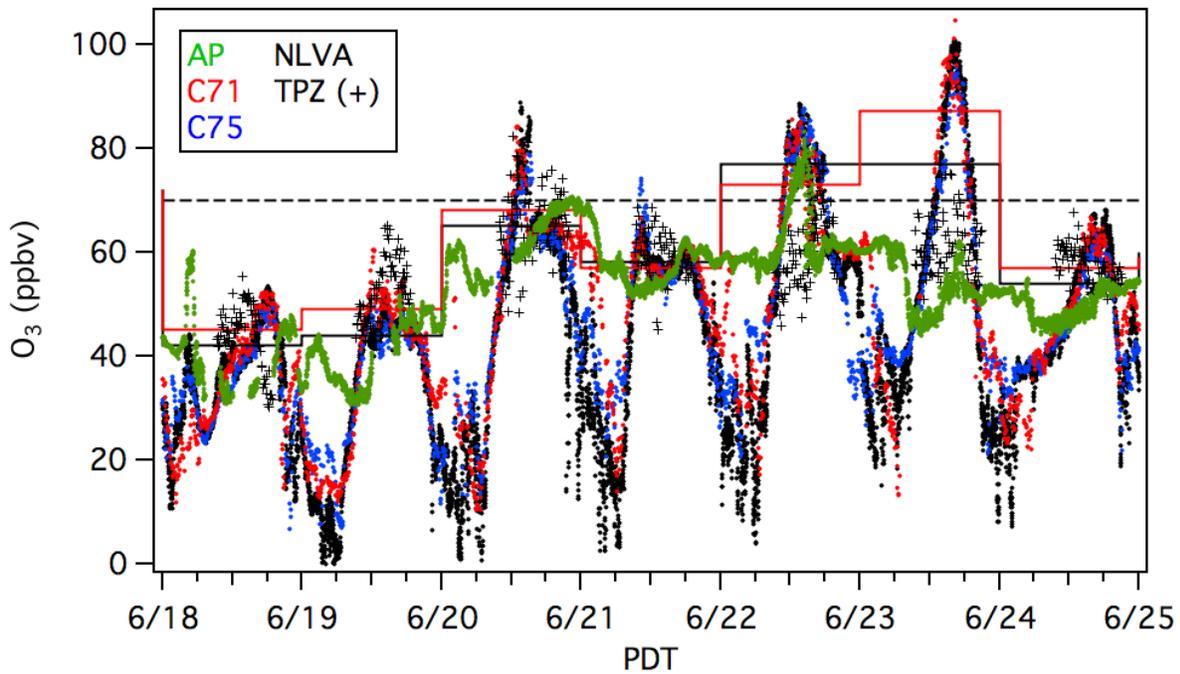
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Figure 9-34. Scientific Aviation O₃ measurements from the flights on June 13 and 14.

1 **9.6 Week 6: June 18-24**

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3 The week of June 18-24 was extremely hot with daily maximum temperatures ranging from
4 43.3 to 46.7°C (110 to 116°F) at the nearby NWS KVG T station. The official KVEF station at
5 McCarran International Airport reached 117°F on June 20, tying the all-time Las Vegas record
6 set on July 24, 1942 (this record was also tied on June 30, 2013, the final day of the LVOS field
7 campaign). The week began with the lowest measured surface O₃ concentrations of the entire
8 campaign and ended with the highest. The low concentrations at the beginning of the week
9 followed the advection of clean marine air into the Desert Southwest from the Eastern Pacific,
10 and the MDA8 O₃ did not exceed 49 ppbv at any of the monitoring sites in the LVV on June 18.
11 Multiple O₃ exceedances occurred on both June 22 (4 monitors) and 23 (5 monitors), however.
12 The highest concentrations were measured on June 23 when the MDA8 O₃ at the Joe Neal
13 monitor reached 77 ppbv. The Walter Johnson monitor officially reported 87 ppbv for June 23,
14 but this value was based on a 6-h average and would have been closer to 77 ppbv had the
15 monitor not been offline for much of the morning. **Figure 9-35** shows that the O₃ mixing ratios
16 at 4 km were less than or equal to the surface concentrations all week.
17



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20 **Figure 9-35.** Time series of the TOPAZ O₃ mixing ratios at 4.0 km agl (black +) during Week 6
21 plotted with the 1-min O₃ measurements from the NLVA (black) and AP (green) and the 5-min
22 measurements from the Walter Johnson (C71, red) and Joe Neal (C75, blue) monitors. The red
23 and blue steps show the corresponding C71 and C75 MDA8 measurements. The dashed black
24 line indicates the 70 ppbv NAAQS.
25
26

1 The FLEXPART BBCO tracer plots in **Figures 8-1 and 8-4** show a large increase in fire activity
2 across the Southwestern U.S. and Mexico during the last two weeks of the study, and the
3 TOPAZ backscatter curtain plot in **Figure 9-36** shows high backscatter appearing around 4 km
4 on the evening of June 18 and filling the lower troposphere on June 20. The backscatter
5 remained high at 4 km on June 21, and was also elevated near the surface on June 23 and at 4
6 km on June 24. The O₃ curtain shows low O₃ extending all the way from the surface to at least 8
7 km on June 18 and remaining low near the surface on June 19. Ozone increased along with
8 backscatter on June 20, and was very high in the mixed layer on June 22 and 23 when the
9 NAAQS exceedances occurred.

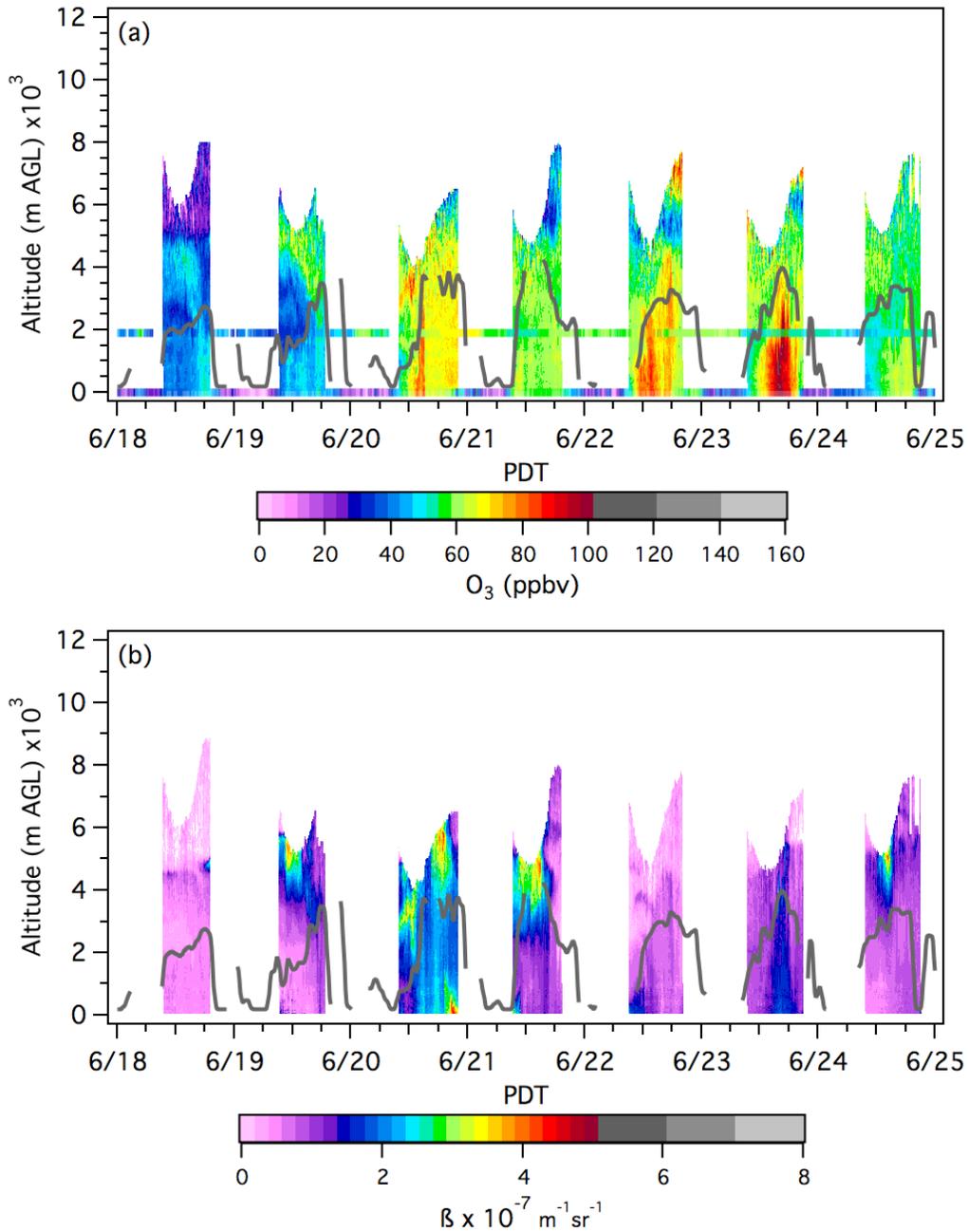
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11 The FLEXPART STO3 tracer curtain in **Figure 9-37** shows tongues of stratospheric O₃ and Asian
12 CO descending into the upper troposphere over the course of the week. The ASCO tracer also
13 shows some transported pollution filling the tropospheric column on June 18-19 when the low
14 O₃ was observed. The BBCO tracer appears quite similar to the backscatter distribution in
15 **Figure 9-36** with the highest values on the night of June 20-21 and a lesser peak on June 23. The
16 BBCO maps in **Figure 8-4** suggest that the tracer appearing above the LVV on June 20 was
17 advected anticyclonically from agricultural fires in Mexico. This conclusion is also supported by
18 the HRRR-smoke model; **Figure 9-38** shows near surface and vertically integrated smoke
19 distribution forecasts for the afternoons (22 UT or 15 PDT) of June 20 and 22. The Mexican
20 smoke was much reduced on June 22, and HRRR-smoke indicates that the tenuous smoke
21 reaching the LVV on that day arrived in a narrow plume from the small (1503 acres) Holcomb
22 Fire near Big Bear. HRRR-smoke also shows the plume from the much larger and closer Brian
23 Head Fire (68407 acres) that started on June 17 in southwestern Utah is streaming eastward
24 and away from Clark County on June 20 and 22, but shows some of the Brian Head smoke
25 reaching the LVV on the morning of June 23 when the highest O₃ was measured.

26
27 The in-situ measurements from Angel Peak (**Figure 9-39 and 9-40**) show broad CO, NO_y, and O₃
28 peaks on June 20 when the TOPAZ backscatter was highest. The O₃ mixing ratios in the aged
29 Mexican smoke plume reached about 70 ppbv. The measurements also show a series of narrow
30 peaks with up to 450 ppbv of CO between 1000 and 1600 PDT on June 22. Matching peaks are
31 also seen in the CH₄, H₂O, and N₂O measurements, and in both the NO₂ and NO_y
32 measurements. The presence of NO₂ shows that the plume is fresh, and the O₃ peak of about
33 80 ppbv shows an enhancement of about 20 ppbv compared to the background. This
34 enhancement is also seen in the 5-min measurements at Walter Johnson and Joe Neal, which
35 peaked at more than 85 ppbv, with the MDA8 O₃ concentrations reaching 73 and 77 ppbv,
36 respectively.

37
38 The HRRR-smoke forecast for June 23 in **Figures 9-38** indicates that most of the smoke from the
39 Brian Head Fire smoke was transported into the LVV near the surface in the early morning, and
40 **Figures 9-35 and 39** show only a small increase in CO and O₃ at Angel Peak created by upslope
41 transport from the LVV. Nevertheless, it seems likely that the Brian Head plume contributed to
42 the very high surface O₃ (5-min peak of 104 ppbv at Walter Johnson) on June 23.

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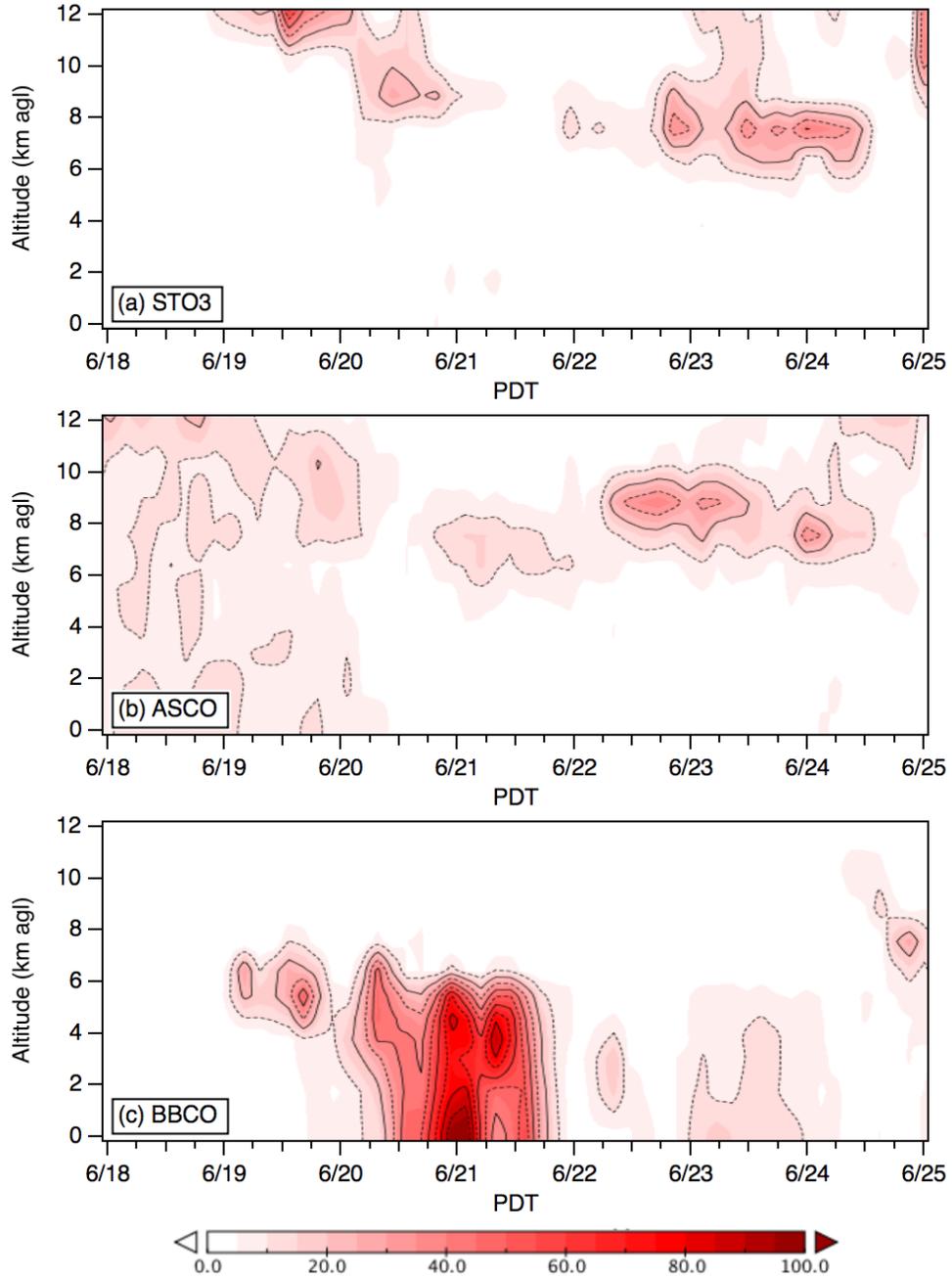
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Figure 9-36. Time-height curtain plots of the TOPAZ (a) ozone, and (b) backscatter, measured during Week 6 of FAST-LVOS. The colored horizontal bands in (a) show the in-situ measurements from the NLVA and AP. The black line shows the local mixed-layer height from the micro-Doppler lidar.

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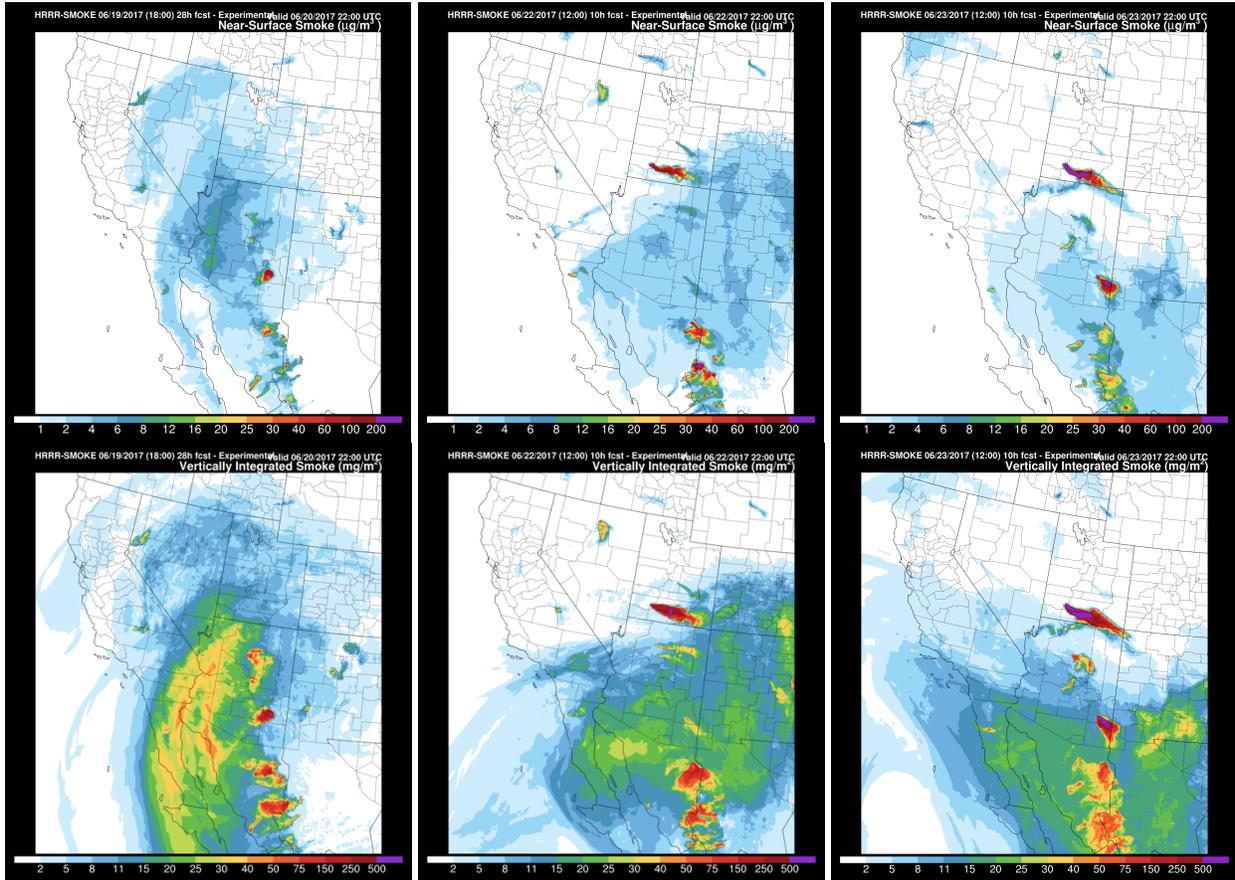


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Figure 9-37. Time-height curtain plots of the FLEXPART (a) STO3, (b) ASCO, and (c) BBCO tracer distributions above the LVV during Week 6 of FAST-LVOS.

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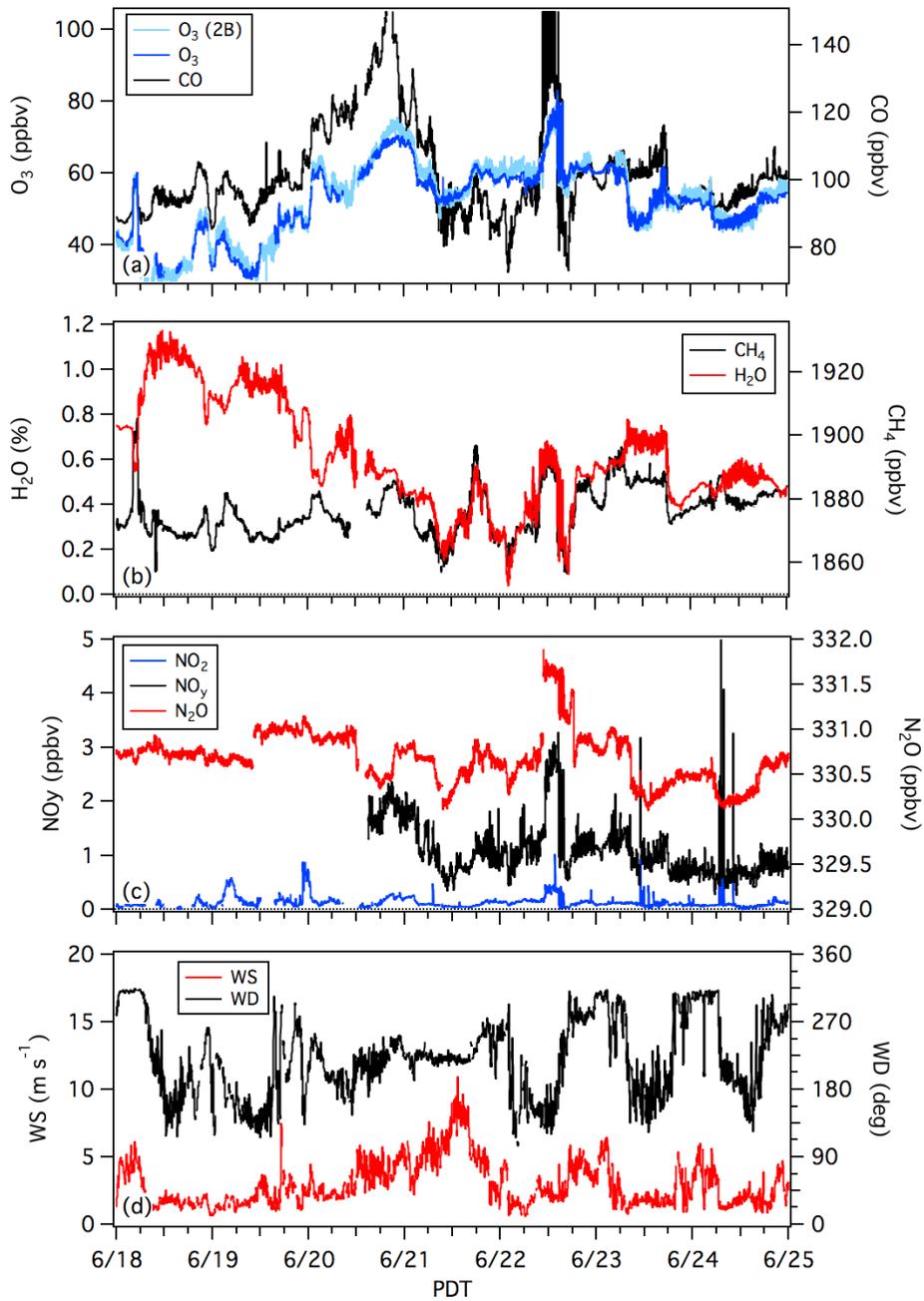
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Figure 9-38. NOAA HRRR-smoke model near surface (top) and vertically integrated (bottom) smoke forecasts for: (left to right) June 20 15 PDT, June 22 15 PDT, and June 23 11 PDT.

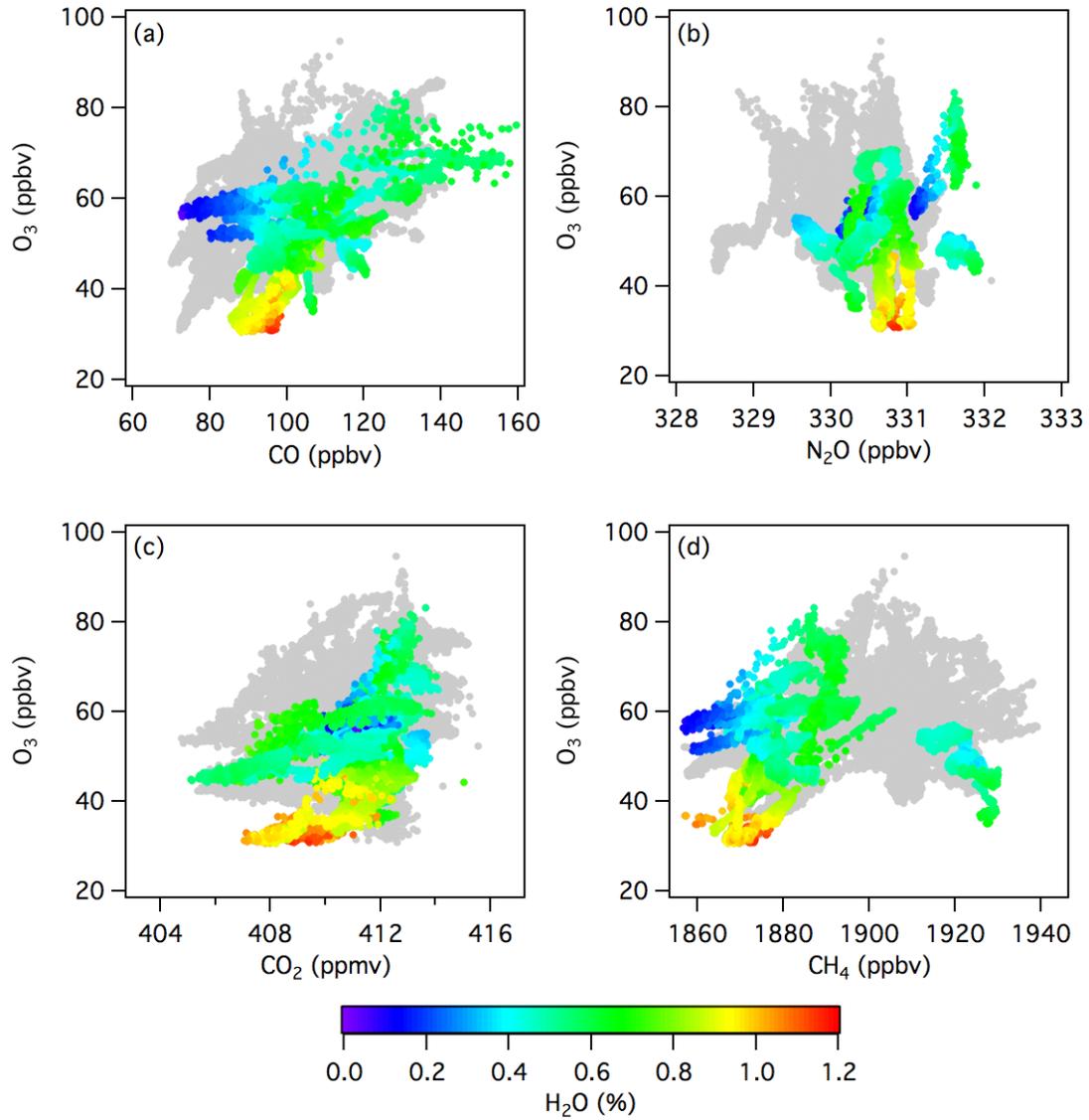
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Figure 9-39. Time series of (a) O_3 and CO, (b) H_2O and CH_4 , (c) NO_2 , NO_y and N_2O , and (d) wind speed and direction measured by the mobile laboratory on Angel Peak during Week 6 of FAST-LVOS. The CO measurements in the Holcomb Fire plume on June 22 reached over 450 ppbv.

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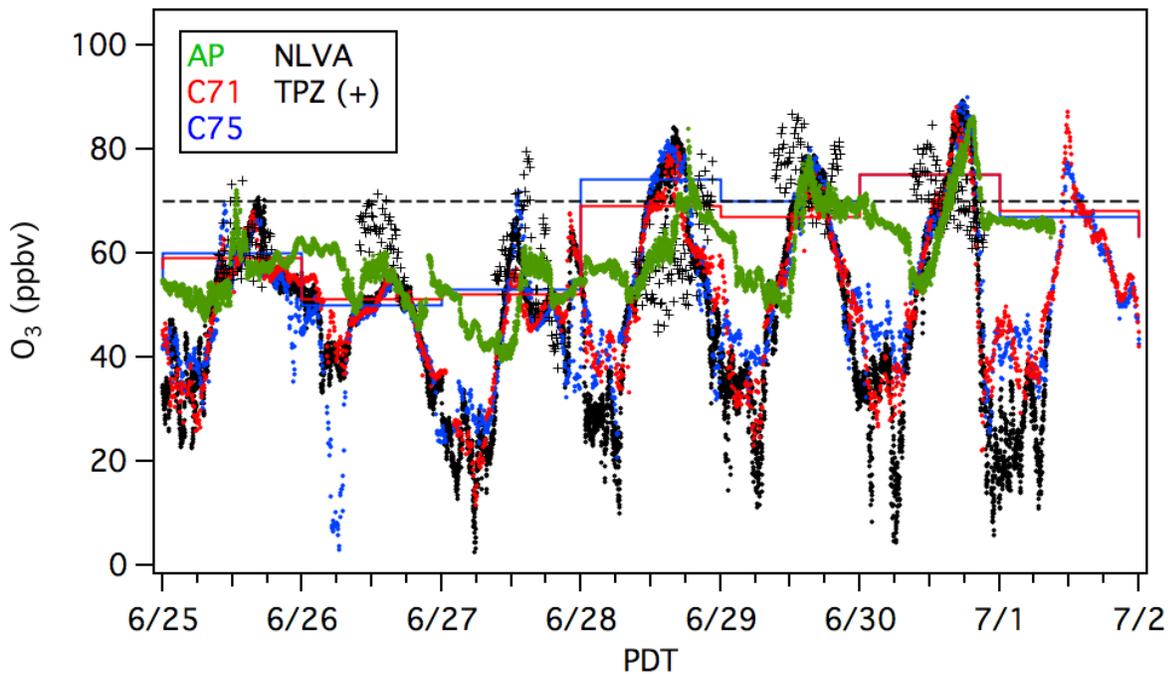
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Figure 9-40. Scatter plots showing the correlations between in-situ O₃ and (a) CO, (b) N₂O, (c) CO₂, and (d) CH₄ measurements from the mobile lab on Angel Peak. The measurements from Week 6 are colorized by the water vapor measurements. The CO measurements in the Holcomb Fire plume are off scale.

1 **9.7 Week 7: June 25-30 (IOP4)**

2
3 The last week of FAST-LVOS was also very hot with the daily high temperatures at KVGT ranging
4 between 44.4 (112°F) on June 25 and 41.1°C (106°F) on June 30. The GFS model showed
5 another developing trough and the final three ozonesondes were launched at local noon on
6 June 28, 29, and 30. The final Scientific Aviation flight hours were expended in flights on June
7 27, 28, 29, and 30. The campaign ended with O₃ exceedances at Walter Johnson and/or Joe
8 Neal on June 28 and 30, but the **Figure 9-41** shows that the O₃ measured by TOPAZ at 4 km was
9 similar to the surface O₃ in the LVV and at AP.

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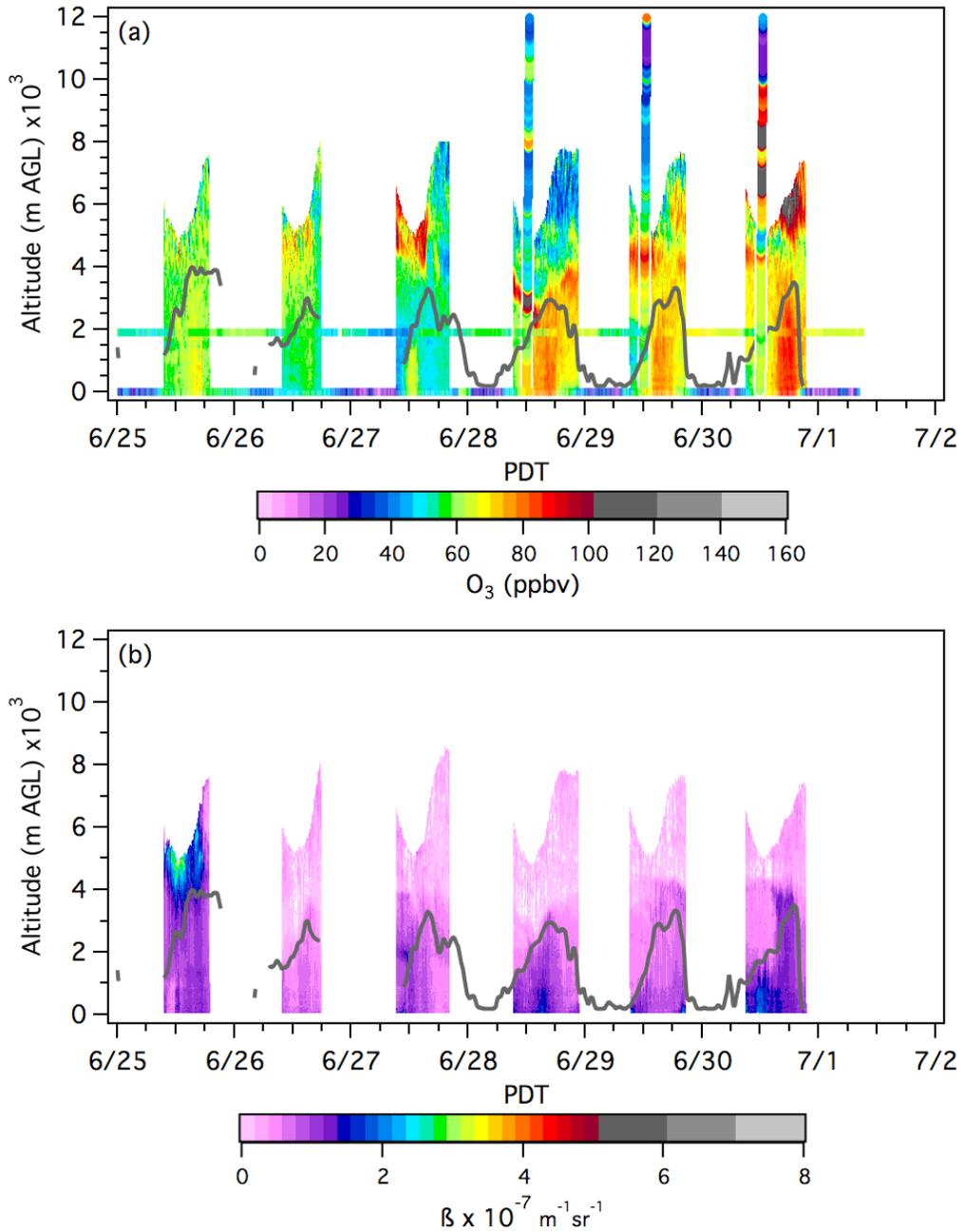


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12 **Figure 9-41.** Time series of the TOPAZ O₃ mixing ratios at 4.0 km agl (black +) during Week 7
13 plotted with the 1-min O₃ measurements from the NLVA (black) and AP (green) and the 5-min
14 measurements from the Walter Johnson (C71, red) and Joe Neal (C75, blue) monitors. The red
15 and blue steps show the corresponding C71 and C75 MDA8 measurements. The dashed black
16 line indicates the 70 ppbv NAAQS.

17
18 The ozone curtain plot in **Figure 9-42** shows high O₃ above 4 km on July 27,29, and 30, with very
19 high concentrations around 6 km on June 30. The backscatter curtain plot shows smoke above
20 4 km on June 25, and possibly in the boundary layer on June 28 and 30. The FLEXPART STO3
21 curtain in **Figure 9-43** show a weak stratospheric intrusion descending into the lower
22 troposphere on June 25. It also shows a small enhancement around 6 km in the early morning
23 of June 30, but the simulations ended in the morning and did not capture the apparent
24 stratospheric intrusion. The ASCO tracer shows Asian pollution in the upper troposphere on all
25 6 days, and the BBCO tracer shows wildfire influence in the lower troposphere on June 25-28.

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4 **Figure 9-42.** Time-height curtain plots of the TOPAZ (a) ozone, and (b) backscatter, measured

5 during Week 7 of FAST-LVOS. The ascending Joe Neal ozonesonde profiles from IOP4 are

6 superimposed. The colored horizontal bands in (a) show the in-situ measurements from the

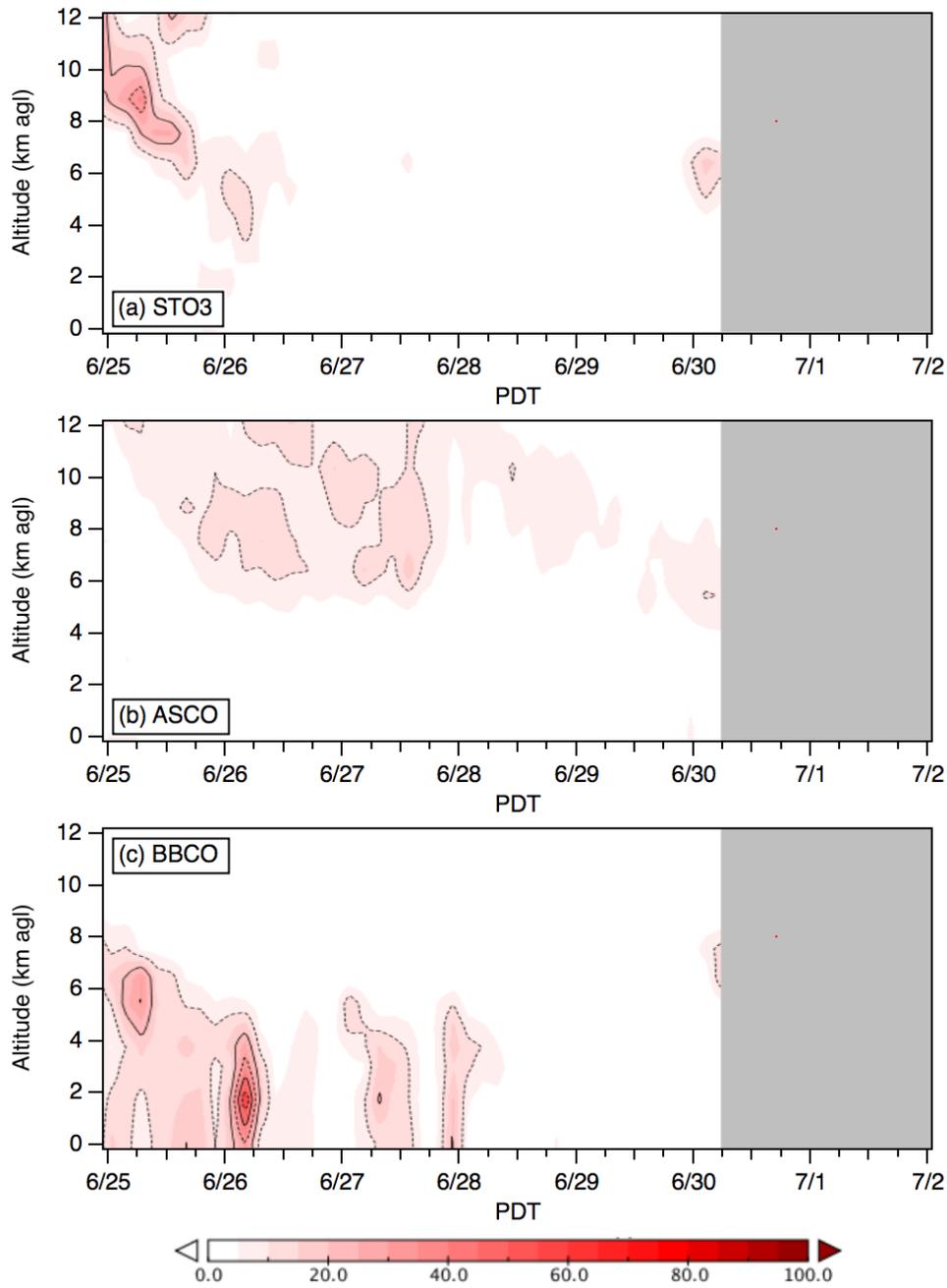
7 NLVA and AP. The black line shows the local mixed-layer height from the micro-Doppler lidar.

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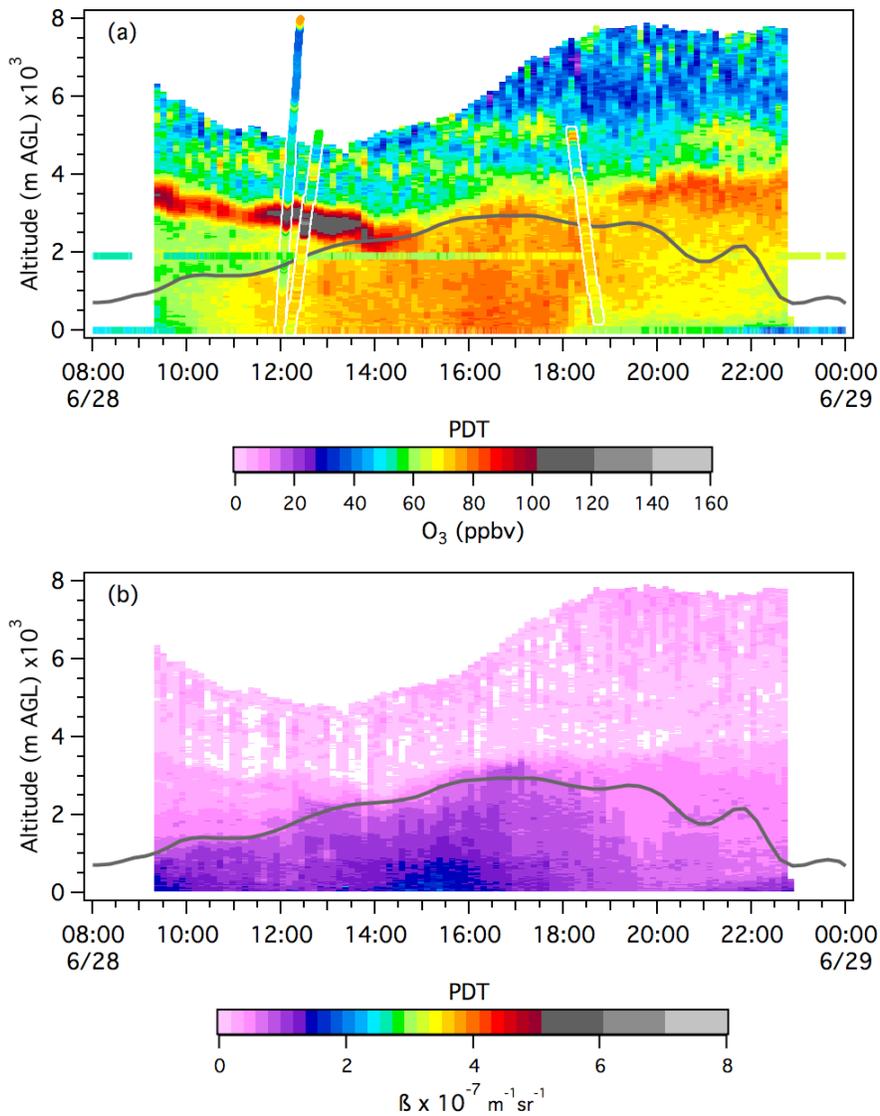
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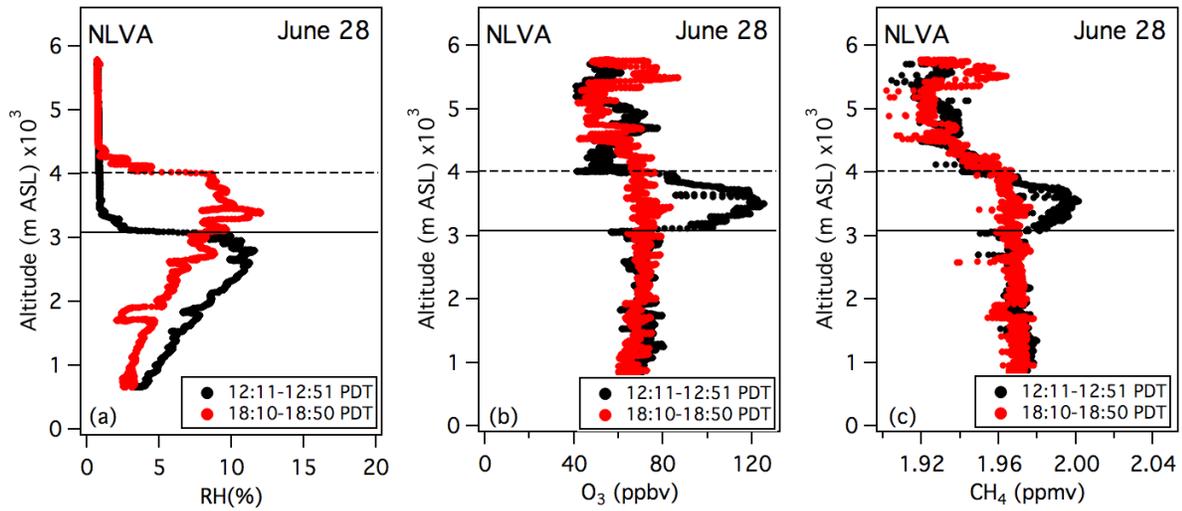
Figure 9-43. Time-height curtain plots of the FLEXPART (a) STO3, (b) ASCO, and (c) BBCO tracer distributions above the LVV during Week 7 of FAST-LVOS.

1 The measurements from June 28 give the clearest example of an elevated O_3 plume being
 2 entrained by the mixed layer. The expanded curtain plots in **Figure 9-44** show a thin plume with
 3 more than 100 ppbv being entrained by the growing mixed layer around 1400 PDT. The
 4 superimposed profiles show that the plume was also sampled above the LVV by the Joe Neal
 5 ozonesonde and the Scientific Aviation Mooney. The Mooney profiles (**Figure 9-45**) show that
 6 the layer also had high CH_4 showing that it didn't originate in the stratosphere.
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 10 **Figure 9-44.** Time-height curtain plots of the TOPAZ (a) ozone, and (b) backscatter, measured on
 11 June 28, 2017. The colored horizontal bands in (a) show the in-situ measurements from the
 12 NLVA and AP, and the nearly vertical colored line shows the profile from the ascending Joe Neal
 13 ozonesonde. The colored tilted lines show the outbound and inbound profiles from the Scientific
 14 Aviation Mooney. The black line represents the local mixed-layer height from the micro-Doppler
 15 lidar.

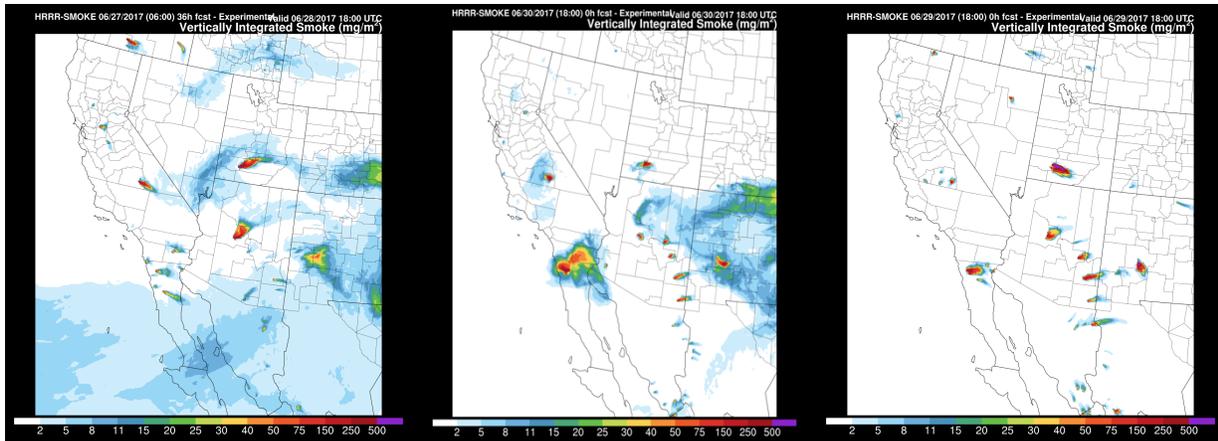
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Figure 9-45. Profiles of (a) RH, (b) O₃, and (c) CH₄ above the NLVA/Joe Neal area from the Scientific Aviation flight on June 28. The black profiles are from the outbound leg near midday and the red profiles from the inbound leg in the late afternoon. The solid and dashed lines show the corresponding mixed layer heights.

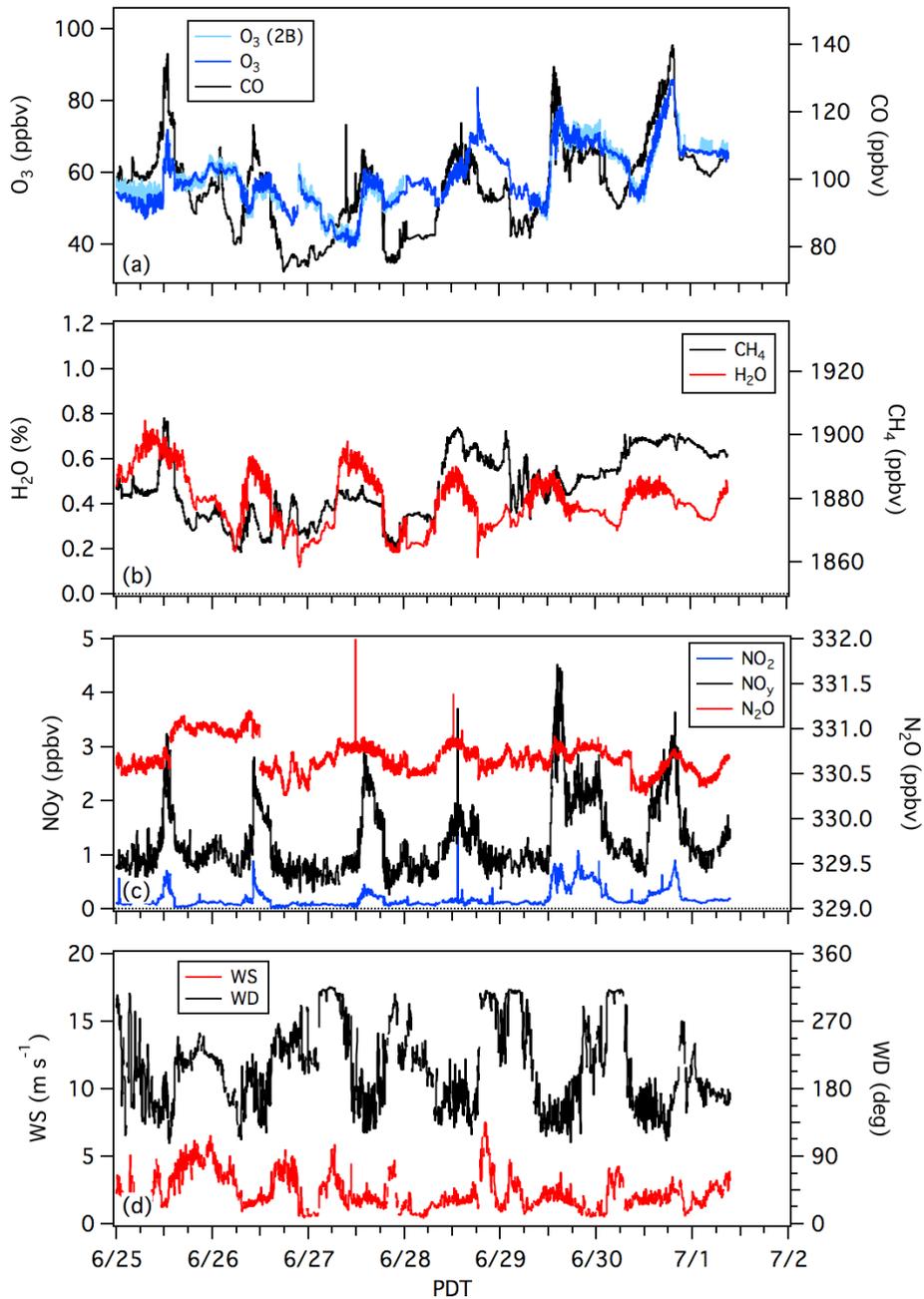
The HRRR-smoke model (**Figure 9-46**) suggests that entrained plume originated from the 16,031-acre Schaeffer Fire in Sequoia National Forest that was started by lightning on June 24.



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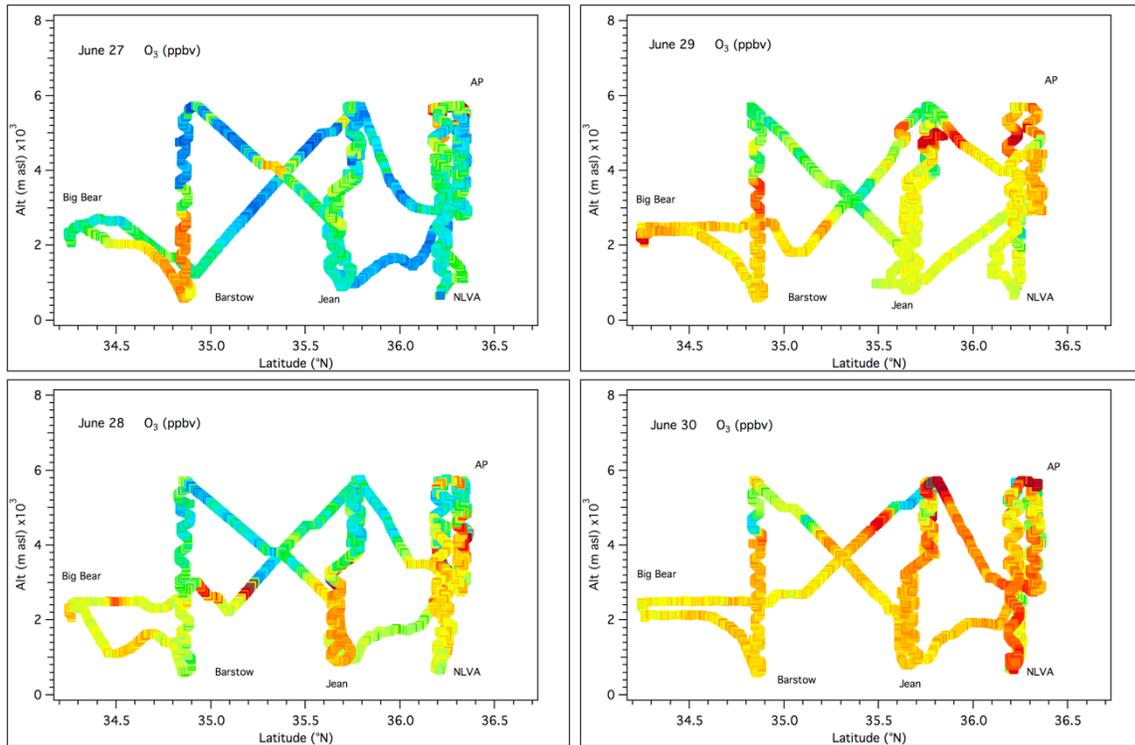
Figure 9-46. NOAA HRRR-smoke model vertically integrated smoke forecasts for: (left to right) June 28, June 29, and June 30 at 18 UT (11 PDT).

1 **Figure 9-44** suggests that the smoke plume did not directly impact Angel Peak, and the
 2 afternoon peaks in O₃, CO, NO_y, and H₂O in the in-situ measurements from the mobile
 3 laboratory plotted in **Figure 9-47** show that the summit was strongly impacted by upslope
 4 transport from the LVV on June 28-30.
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 8 **Figure 9-47.** Time series of (a) O₃ and CO, (b) H₂O and CH₄, (c) NO_x, NO_y and N₂O, and (d) wind
 9 speed and direction measured by the mobile laboratory on Angel Peak during Week 7 of FAST-
 10 LVOS.
 11

1 The Scientific Aviation flights from IOP4 show the gradual buildup of O₃ across the Mojave
2 Desert during the last 4 days of June. The MDA8 O₃ concentrations on June 30 were highest in
3 the western LVV with 75 ppbv at both Walter Johnson and Joe Neal, but the MDA8 also reached
4 69 ppbv at Jean and 66 ppbv at Grand Canyon National Park.
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9 **Figure 9-48.** Ozone measurements from the Scientific Aviation flights to Big Bear, CA on June 27-
10 30.
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10. Comparison with the NOAA GFDL AM4 model

The preceding analyses can be compared to the global model results of Li Zhang and Meiyun Lin of NOAA GFDL and Princeton University that were conducted under a separate contract (CBE-605334-19) with Clark County. NOAA GFDL used the AM4 and GEOS-Chem global models to examine ozone in the southwestern U.S. during the FAST-LVOS study period. The following discussion is based primarily on a summary of the AM4 results recently submitted for publication by Zhang et al. [Zhang et al., 2019]. Details of the model calculations are given in that publication and in the final report for Contract No. CBE605334-19.

The GFDL analyses focused on the deep stratospheric intrusion event of June 11–14 (Section 9-5), the wildfire event of June 22 (Section 9-6), the regional pollution episode of June 16 (Section 9-5), the Asian transport event of May 24 (Section 9-2), and the “mysterious” event of June 28 (Section 9-7), which we attribute to the Schaeffer Fire. Figures 10-1 to 10-3, which are adapted from Zhang et al. [Zhang et al., 2019], show curtain plots comparing the AM4 and GEOS-Chem model results to the TOPAZ measurements.

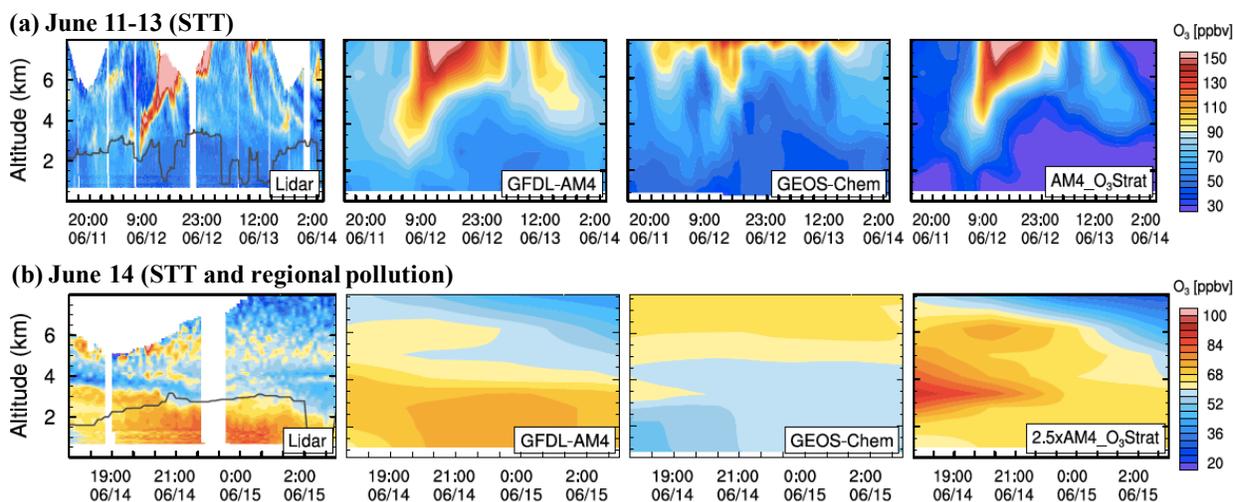
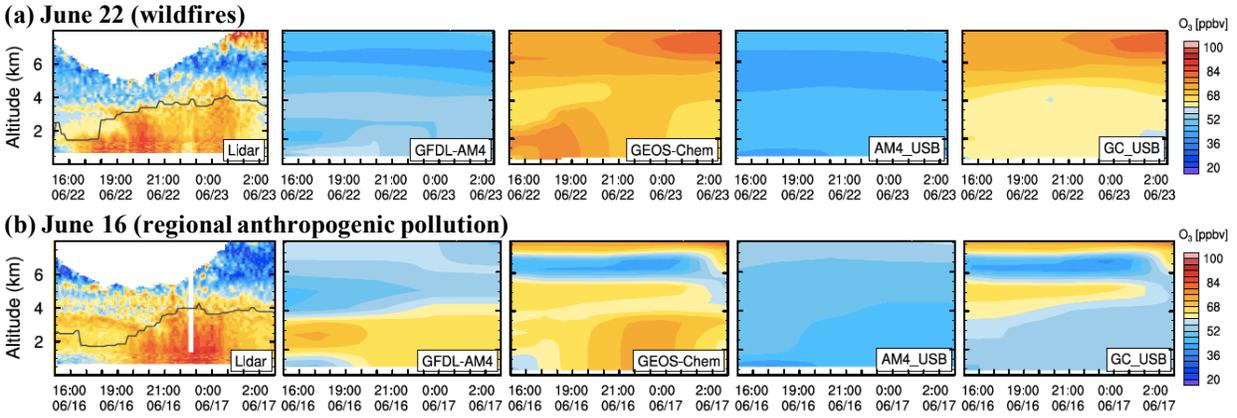
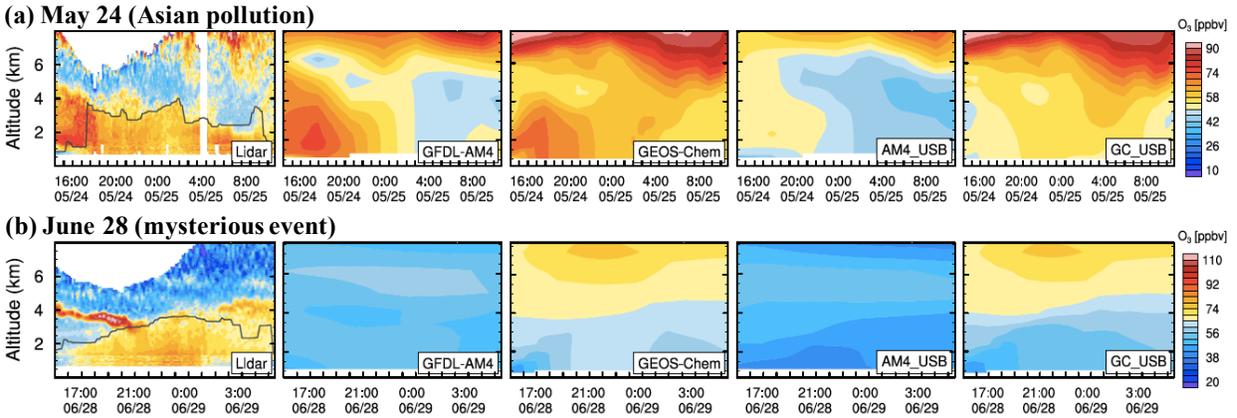


Figure 10-1. Time-height curtain plots of O_3 above NLVA as observed with TOPAZ lidar and simulated with GFDL-AM4 ($\approx 50 \text{ km} \times 50 \text{ km}$; interpolated from 3-hourly data) and GEOS-Chem ($0.25^\circ \times 0.3125^\circ$; interpolated from hourly data) during the STT event on (a) June 11–13 and (b) June 14, 2017 (UTC). The rightmost panel shows AM4 stratospheric O_3 tracer (AM4_ O_3 Strat). Note that AM4_ O_3 Strat for June 14 is scaled by a factor of 2.5 for clarity. Here and in other figures, the solid black lines in the O_3 lidar plots represent boundary layer height inferred from the micro-Doppler lidar measurements.



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Figure 10-2. Same as Figure 10-1, but for (a) the wildfire event on June 22 and (b) the regional anthropogenic pollution event on June 16, 2017 (UTC). The right panels compare USB O₃ from the two models.



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Figure 10-3. Same as Figure 10-1, but for (a) the Asian pollution event on May 24 and (b) the unattributed pollution event on June 28, 2017 (UTC).

9

10 **Figures 10-1 through 10-3** show that AM4 does a good job of capturing the stratospheric
 11 intrusion (June 11-13), regional pollution (June 14 and 16), and Asian pollution (May 24)
 12 events, but is unable to resolve the two wildfire events (June 22 and 28) due to the relatively large grid
 13 size of the global model. The GEOS-Chem model fails to reproduce the stratospheric intrusion,
 14 but does a better job with the regional and Asian pollution events. The higher resolution GEOS-
 15 Chem model also captures the first wildfire event, but misses the June 28 fire plume.

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1 The global model simulations should, in principle, be better able to resolve the contributions of
2 intermingled sources as when stratospheric intrusions mix with local or regional pollution.
3 **Figure 10-4** separates the U.S. Background (USB), which includes the stratospheric contribution,
4 from the AM4 total model MDA8 O₃. The modeled surface USB in Clark County peaks at about
5 60 ppbv on June 11 and decreases to about 50 ppbv on June 12. These results are in good
6 agreement with the Clark County surface measurements from June 11 and 12 plotted in **Figure**
7 9-26. The USB continues to decrease to 45-50 ppbv on June 14 as the intrusion tracks south and
8 east. This suggests that the local and regional contribution to the total MDA8 of ≈70 ppbv on
9 June 14 is about 20-25 ppbv.

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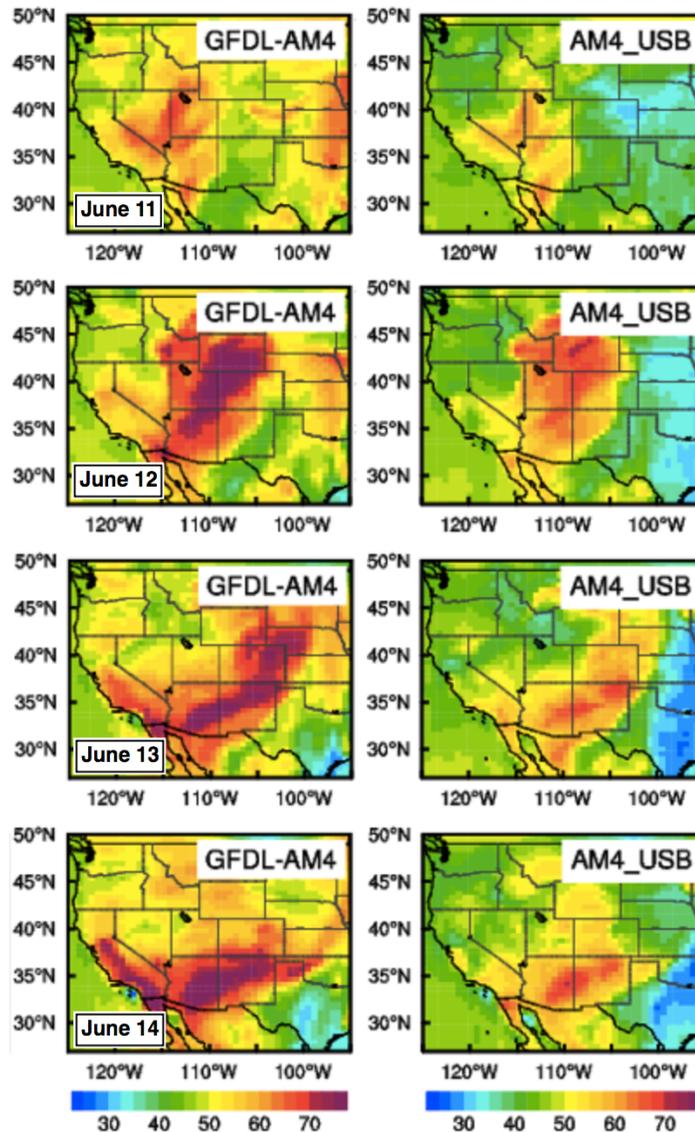


Figure 10-4. Maps of total MDA8 O₃ as observed and simulated with GFDL-AM4 along with the model estimated USB level, during the STT event of June 11–14, 2017. (adapted from Zhang et al., 2019).

1 **11. Contrasts between LVOS and FAST-LVOS**

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3 The FAST-LVOS measurements directly confirm the hypothesis that entrainment of high O₃
4 layers aloft by the deep mixed layers of the Mojave Desert can and does influence surface O₃
5 concentrations in the LVV. In contrast to 2013, however, this entrainment did not directly lead
6 to exceedances of the O₃ NAAQS in Clark County during FAST-LVOS. In the following section, we
7 examine the differences between the 2013 and 2017 measurements and consider possible
8 explanations and implications.
9

10 **Table 11-1** lists the 4 highest MDA8 O₃ values measured by each of the active Clark County
11 regulatory monitors during 2013 (http://airquality.clarkcountynv.gov/cgi-bin/8hr_4highest.pl).
12 The MDA8 values are highlighted to reflect the revised Air Quality Index (AQI) with values
13 exceeding the 2015 NAAQS shown in red or orange. The 2008 NAAQS (75 ppbv) then in effect
14 was exceeded by one or more of the Clark County regulatory O₃ monitors on 6 days in 2013
15 with 18 reported exceedances. The 2015 NAAQS was exceeded 82 times on 25 days
16 (http://airquality.clarkcountynv.gov/cgi-bin/8hr_exceed.pl).
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19 **Table 11-1. Four highest O₃ days at the Clark County regulatory monitors in 2013**

Area	Monitoring Site	POC	Highest			Second Highest			Third Highest			Fourth Highest		
			Date	Time	Value	Date	Time	Value	Date	Time	Value	Date	Time	Value
Clark County														
	Apex	1	06/21/2013	1000	78	04/30/2013	0800	74	05/05/2013	0600	73	05/04/2013	1400	73
	Mesquite	1	05/22/2013	1100	68	05/16/2013	1100	68	06/21/2013	1100	67	06/18/2013	1100	67
	Paul Meyer	1	07/03/2013	1100	87	05/04/2013	1400	80	05/25/2013	1000	76	06/21/2013	1000	75
	Walter Johnson	1	07/03/2013	1100	87	05/04/2013	1400	80	05/25/2013	1200	75	07/19/2013	0900	74
	Palo Verde	1	07/03/2013	1000	83	05/04/2013	1400	82	05/25/2013	1200	76	07/19/2013	1000	74
	Joe Neal	1	07/03/2013	1100	81	06/21/2013	1000	77	05/04/2013	1400	77	07/20/2013	1000	76
	Winterwood	2	05/04/2013	1600	76	06/21/2013	1000	75	05/25/2013	0900	73	05/21/2013	1500	71
	Jerome Mack	1	05/04/2013	1400	74	05/25/2013	1100	73	06/21/2013	1000	72	05/21/2013	1400	69
	Boulder City	1	06/21/2013	1100	74	05/22/2013	1400	72	05/21/2013	1600	72	06/22/2013	1000	71
	Jean	1	05/04/2013	1300	84	05/21/2013	1400	78	05/25/2013	1100	76	06/21/2013	1100	75
POC (Parameter Occurrence Code): a code used to correctly separate data from multiple instruments at one site.														

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22
23 Three of the six 75 ppbv exceedance days occurred during the LVOS field campaign: May 21,
24 May 25, and June 21. The LVOS measurements and modeling convincingly showed that all of
25 these exceedances were caused, at least in part, by non-controllable ozone sources (NCOS)
26 [Jaffe et al., 2018] including stratospheric intrusions and Asian pollution. There is strong
27 evidence that the other 3 exceedance days in 2013 were also influenced by NCOS, with the
28 highest O₃ concentrations (July 3) linked to the nearby 28,000-acre Carpenter 1 Fire on Mt.
29 Charleston and the next highest (May 4) with the stratospheric intrusion and wind driven
30 24,251-acre Springs Fire in Ventura County, CA [Langford et al., 2015a]. The high O₃
31 concentrations measured on the one remaining exceedance day (July 19) were likely influenced

1 by the 27,531-acre Mountain Fire in Riverside County, CA. These regional and synoptic-scale
 2 events affected hundreds of km² and raised the O₃ concentrations across most of Clark County.

3
 4 **Table 11-2** shows the corresponding 4 highest O₃ days for 2017. The total number of NAAQS
 5 exceedances was much smaller in 2017 with 10 reported exceedances of the 2008 NAAQS and
 6 43 exceedances of the 2015 NAAQS ([http://airquality.clarkcountynv.gov/cgi-](http://airquality.clarkcountynv.gov/cgi-bin/8hr_exceed.pl)
 7 [bin/8hr_exceed.pl](http://airquality.clarkcountynv.gov/cgi-bin/8hr_exceed.pl)). The corresponding number of days with exceedances of the 2008 and 2015
 8 NAAQS were 6 and 18, respectively. The highest MDA8 of 87 ppbv recorded in both 2013 and
 9 2017 was measured by the Walter Johnson monitor².

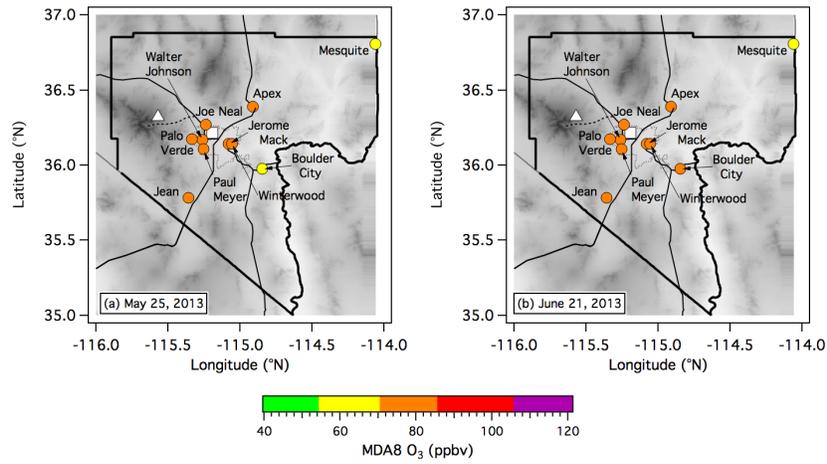
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 12 **Table 11-2. Four highest O₃ days at the Clark County regulatory monitors in 2017**

Area	Monitoring Site	POC	Highest			Second Highest			Third Highest			Fourth Highest		
			Date	Time	Value	Date	Time	Value	Date	Time	Value	Date	Time	Value
Clark County														
	Apex	1	07/11/2017	1000	73	06/28/2017	1000	71	06/22/2017	1000	70	07/04/2017	0900	69
	Mesquite	1	05/14/2017	1100	63	05/12/2017	1400	63	04/23/2017	1100	63	07/02/2017	1000	62
	Paul Meyer	1	08/31/2017	0900	76	07/12/2017	0900	71	06/23/2017	1000	71	07/15/2017	0900	70
	Walter Johnson	1	06/23/2017	1000	87	07/07/2017	0900	79	07/04/2017	0900	76	06/30/2017	1100	75
	Palo Verde	1	07/07/2017	0900	76	06/23/2017	1000	76	06/16/2017	0900	75	07/12/2017	0900	74
	Joe Neal	1	07/04/2017	0900	80	06/23/2017	1100	77	06/22/2017	0900	77	07/15/2017	1000	76
	Green Valley	1	07/11/2017	0900	73	06/17/2017	0900	72	08/31/2017	1000	71	07/08/2017	0900	70
	Jerome Mack	1	08/31/2017	1000	70	07/04/2017	0900	67	06/17/2017	0900	66	07/11/2017	1100	65
	Boulder City	1	06/28/2017	1200	69	05/13/2017	1000	68	07/11/2017	0900	67	04/23/2017	1200	67
	Jean	1	06/30/2017	1000	69	04/11/2017	1200	69	07/03/2017	1100	67	10/04/2017	1100	66
	Indian Springs	1	07/03/2017	0900	69	07/04/2017	0900	66	07/01/2017	0900	66	05/11/2017	1000	66
POC (Parameter Occurrence Code): a code used to correctly separate data from multiple instruments at one site.														

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 16 **Tables 11-1** and **11-2** appear qualitatively different with fewer monitors exceeding the 2015
 17 NAAQS in 2017 than in 2013. Furthermore, in contrast to 2013 when nearly all of the monitors
 18 exceeded the NAAQS on the highest O₃ days (**Figure 11-1**) no more than 4 monitors exceeded
 19 the NAAQS on the 2017 highest O₃ days in 2017 (**Figure 11-2**). Note that while several
 20 monitoring stations were not operational during both 2013 and 2017, the core monitors were
 21 the same in both years. The more limited spatial extent of the high surface O₃ in 2017
 22 compared to 2013 suggests that stratospheric intrusions did not contribute as much to the
 23 2017 exceedances. The O₃ scale in both figures reflects the EPA AQI colors, with yellow colors
 24 corresponding to “Moderate” concentrations and orange colors “Unhealthy for Sensitive
 25 Groups”.

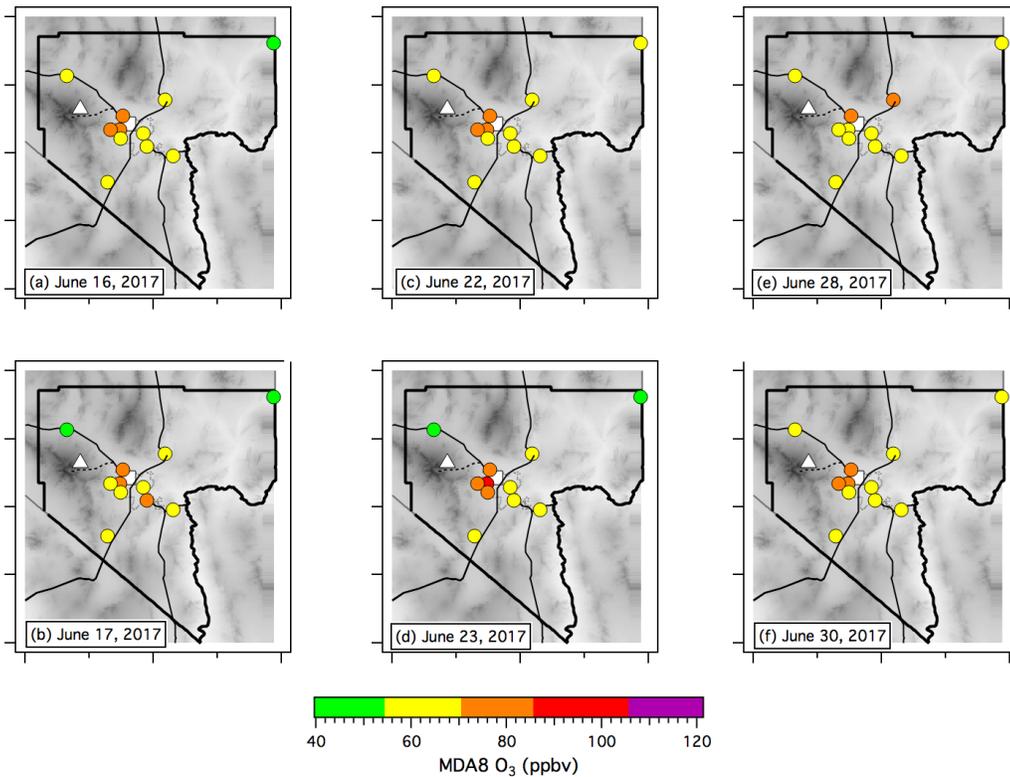
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² The MDA8 of 87 ppbv recorded at Walter Johnson on June 23, 2017 is only a 6-h average since
 the monitor was offline for calibration in the morning. Comparisons with the nearby Joe Neal
 and Palo Verde monitors suggest that the full 8-h MDA8 would have been closer to 77 ppbv.

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Figure 11-1. MDA8 O₃ measured by the Clark County regulatory monitors on two of the three NAAQS exceedance days that occurred during the 2013 LVOS field campaign: (a) May 25, and (b) June 21. The white square and triangle show the locations of NLVA and Angel Peak, respectively.



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Figure 11-2. MDA8 O₃ measured by the Clark County regulatory monitors on days with multiple exceedances during FAST-LVOS. The white squares and triangles show the locations of the NOAA measurement sites at NLVA and AP, respectively.

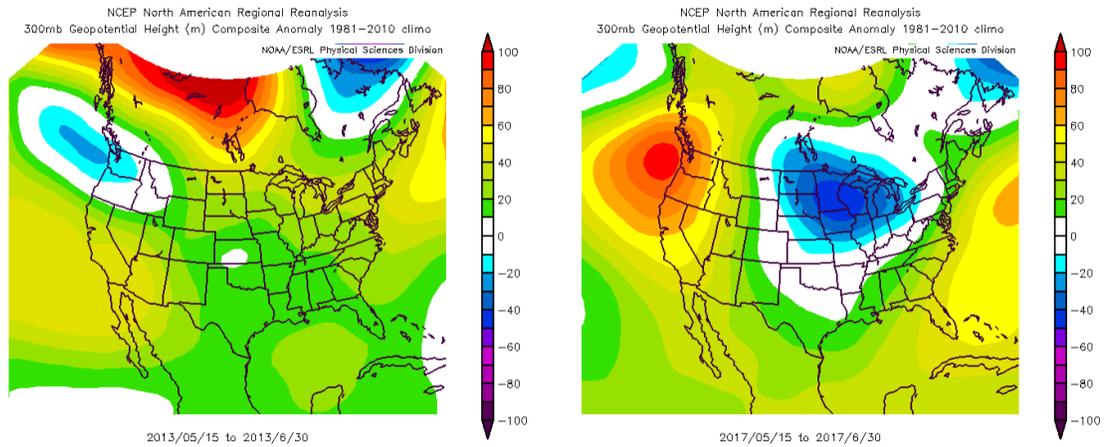
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3 The contributions of STT to tropospheric O₃ in the western U.S. have been linked to climatic
4 oscillations including ENSO [Lin et al., 2015] and the Northern Annular Mode (NAM) [Albers et
5 al., 2018]. ENSO modulates STT by changing the mean positions of the subtropical [Langford et
6 al., 1998] and polar front jets [Lin et al., 2015]. Under El Niño conditions, the PFJ migrates
7 northward over Canada and the subtropical jet moves eastward toward the coast of southern
8 California. This leads to increased STT of O₃ into the middle and upper troposphere by the STJ
9 [Langford, 1999], but less STT of O₃ into the lower troposphere. Under La Niña conditions, the
10 STJ shifts westward and the PFJ meanders southward over the western U.S. where it creates
11 more deep tropopause folds.

12
13 The equatorial Pacific remained in an ENSO-neutral state during the winter (November 2012-to
14 January 2013) preceding the LVOS campaign, but the precipitation patterns in the western U.S.
15 (<https://www.ncdc.noaa.gov/sotc/synoptic/201301>) more closely resembled La Niña
16 conditions. The late spring of 2013 was accordingly characterized by a very active jet stream
17 that brought frequent upper-level troughs through the Pacific Northwest and across the
18 contiguous U.S. during both May and June (<https://www.ncdc.noaa.gov/sotc/synoptic/201306>).
19 These upper-level lows spawned several large tropopause folds and were sufficiently deep and
20 frequent to create a low-pressure anomaly on the 300 hPa geopotential surface compared to
21 the 30-year (1981-2010) climatology (**Figure 11-3a**).

22
23 The equatorial Pacific was characterized by La Niña conditions in November of 2016, but
24 became ENSO-neutral by January 2017 (<https://www.ncdc.noaa.gov/sotc/synoptic/201701>).
25 The precipitation patterns above the western U.S. exhibited none of the La Niña of 2013, and
26 the upper-level troughs that moved through the western U.S. during FAST-LVOS did not dip as
27 far south or penetrate as deeply into the troposphere as those that occurred during LVOS.
28 **Figure 11-3b** shows a persistent high-pressure ridge parked over the U.S. West Coast that
29 pushed the jet stream northward into Canada during FAST-LVOS, bringing warmer and drier-
30 than-normal weather and an active wildfire season to the Southwestern U.S.
31 (<https://www.ncdc.noaa.gov/sotc/synoptic/201705>). This greatly diminished the influence of
32 stratospheric intrusions and Asian pollution on surface O₃ in Clark County compared to 2013.

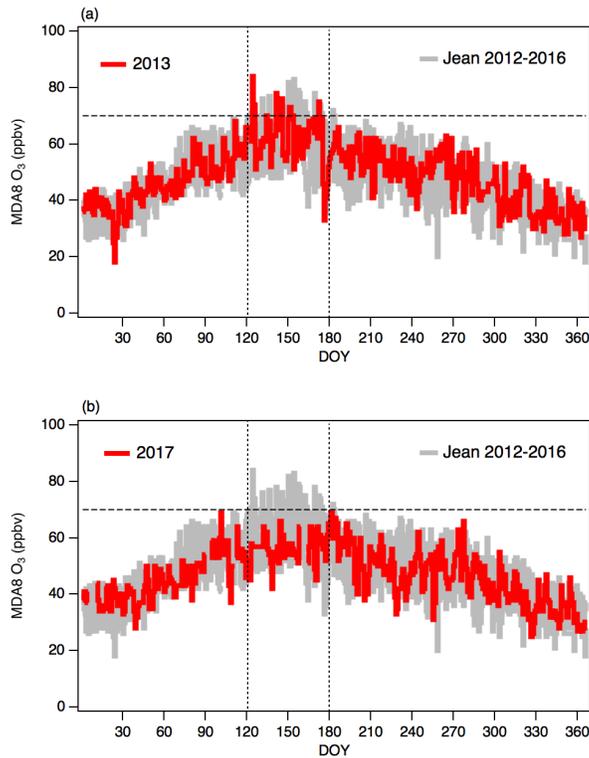
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34 These differences are reflected in the MDA8 O₃ measurements from Jean which are shown in
35 **Figure 11-4**. The MDA8 O₃ measured during May and June of 2017 (57.6±5.7 ppbv) was
36 generally lower and with fewer short-lived excursions than in the previous 5 years including
37 2013 (62.2±9 ppbv).

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 2 **Figure 11-3.** NCEP North American Reanalysis (NARR) 300 hPa geopotential anomalies during
 3 (a) LVOS, and (b) FAST-LVOS field campaigns. Plots provided by the NOAA/ESRL Physical
 4 Sciences Division (PSD).

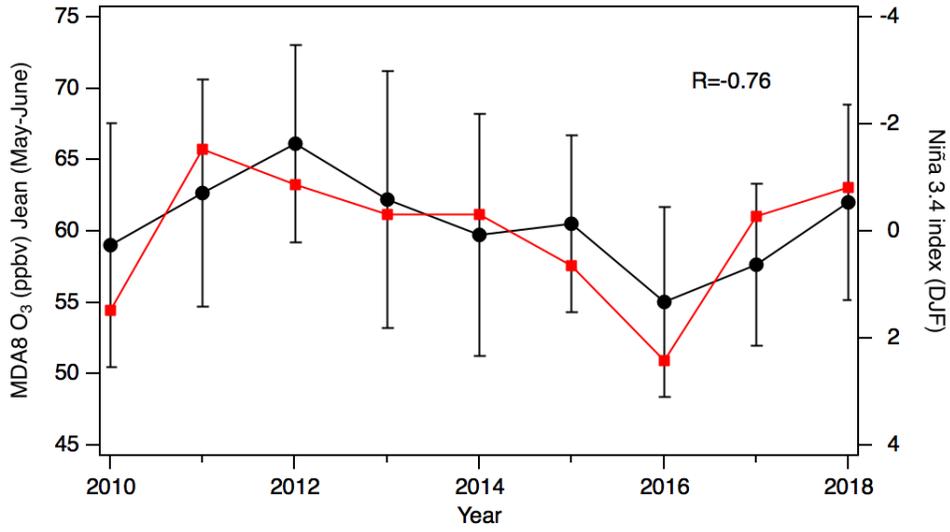
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 10 **Figure 11-4.** MDA8 O₃ measured by the Clark County DAQ monitor at Jean for the years (a)
 11 2013 and (b) 2017. The measurements from 2012-2016 shown in Figure 2-2 are plotted in gray.
 12 The horizontal dashed line marks the 70 ppbv NAAQS. May and June lie between the two
 13 vertical dotted lines.

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Figure 11-5 shows that these differences are part of a larger interannual variability (IAV) with the ENSO influence noted above. The correlation with the Niña 3.4 anomaly index [Rayner *et al.*, 2003] is robust and the IAV significant, with the mean May-June MDA8 O₃ ranging by more than 10 ppbv from 55.0±6.6 ppbv in 2016 to 66.2±6.9 ppbv in 2012. Note that the 2012 mean value lies within 4 ppbv of the 2015 NAAQS.



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Figure 11-5. Mean May-June MDA8 O₃ measured at Jean plotted with the Niña 3.4 index (https://www.esrl.noaa.gov/psd/gcos_wgsp/Timeseries/Nino34/).

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15

1 **Appendix A: Model Descriptions**

2 ***NOAA NESDIS RAQMS model***

3 The NOAA/NESDIS RAQMS (Realtime Air Quality Modeling System) model is a unified
4 (stratosphere-troposphere) online global chemical and aerosol assimilation/forecasting system
5 developed specifically to support airborne field missions [*Pierce et al., 2003; Pierce et al., 2007*].
6 The meteorological forecasts are conducted using the University of Wisconsin Hybrid Model
7 [*Schaack et al., 2004*], and the chemical forecasts are initialized daily at 12 UT with real-time
8 assimilation of Ozone Monitoring Instrument (OMI) cloud-cleared total column ozone and
9 Microwave Limb Sounder (MLS) ozone profiles from the NASA *Aura* satellite, and MODIS
10 aerosol optical depth from the NASA *Terra* and *Aqua* satellites. RAQMS has been run routinely
11 in a forecast mode since 2010 with 2° x 2° resolution analyses and forecasts prior to 2012, and
12 1° x 1° analyses since 2012. The model predicts global O₃, CO, SO₄, and black organic carbon
13 distributions at 6-hour intervals for the next 4 days. The real-time analyses are archived online
14 (<http://raqms-ops.ssec.wisc.edu>).
15

16 ***NOAA ESRL RAP-Chem model***

17 The NOAA ESRL Global Systems Division (GSD) RAP-Chem forecast system
18 (<https://rapidrefresh.noaa.gov/RAPchem/Welcome.cgi>) uses the Weather Research and
19 Forecasting model coupled to Chemistry (WRF-Chem) model [*Grell et al., 2005; Pagowski et al.,*
20 2010] driven by meteorology from the NCEP's GFS analysis at 00 UTC and 6-hourly forecasts
21 thereafter. This 13-km resolution North American model uses the Regional Atmospheric
22 Chemistry Mechanism (RACM) [*Goliff et al., 2013*] and aerosol scheme by *Ahmadov et*
23 *al.* [*Ahmadov et al., 2012*]. It incorporates MEGAN biogenic emissions [*Guenther et al., 2006*],
24 and the U.S EPA NEI-11 (CONUS) version-1 anthropogenic emission inventories with oil/gas
25 sector emissions updated to version-2. Biomass burning emissions are derived from MODIS and
26 GOES-West satellite observations.

27 The model includes surface deposition, photolysis, and convective and turbulent chemical
28 transport, and wet removal of aerosols by resolved and convective precipitation with advective
29 chemical transport performed simultaneously with the meteorology. The lateral boundary
30 conditions for chemistry are obtained from the RAQM model [*Pierce et al., 2003*]. Operational
31 forecasts are run at 00 UTC at 3-h forecast intervals for up to 48 hours, and plots of the CO and
32 O₃ distributions on the 500, 700, and 850 hPa and surface levels were archived for the entire
33 CABOTS period by M. Pagowski of the NOAA ESRL Global Systems Division and the Cooperative
34 Institute for Research in Environmental Sciences (CIRES) at the University of Colorado.
35

36 ***FLEXPART particle dispersion model***

37
38 FLEXPART (FLEXible PARTicle) [*Stohl et al., 2005*] is a passive tracer model (no chemistry) that
39 calculates the evolving distribution of a multitude of “particles” released from a specified
40 source region and transported forward in time for up to 20 days. FLEXPART calculations played

1 a key role in earlier TOPAZ studies of long-range transport and stratosphere-to-troposphere
2 transport (STT) in the western U.S. including LVOS [Langford et al., 2012; Langford et al., 2015].
3 For FAST-LVOS, evolving particle distributions representing stratospheric ozone (STO3), Asian
4 pollution (ASCO) and biomass burning (BBCO) were calculated with 0.25° resolution over an
5 output domain extending from 130 to 70°W and from 20 to 50°N. The particles were
6 transported by winds from the European Centre for Medium-Range Weather Forecasts
7 (ECMWF) operational (0.5° x 0.5°) model forecasts and by sub grid motions including
8 convection and turbulence based on the parameterization schemes of Emanuel and Zivkovic-
9 Rothman [Emanuel and Zivkovic-Rothman, 1999] and Hanna [Hanna, 1982] respectively. The
10 tracer distributions were calculated at the surface (1 km above ground level) and on the
11 700hPa, 500hPa, 400hPa levels at 3-h intervals from May 15 to June 30, 2017.

12
13 The stratospheric O₃ tracer was represented by particles released into the stratosphere (>2
14 potential vorticity units or PVU) and converted to O₃ mixing ratios using a linear relationship
15 between O₃ and potential vorticity (60 ppbv/PVU) at the particle origin. The Asian CO tracer is
16 based on the amount of CO released into the boundary layer from anthropogenic sources in
17 East Asia using the EDGAR 3.2 fast track inventory [Olivier et al., 2005]. Biomass burning CO
18 emissions from North American sources were calculated using the ECMWF Copernicus
19 Atmosphere Monitoring System (CAMS) near-real time biomass burning emission estimates
20 based on the Global Fire Assimilation System (GFAS), which converts Fire Radiative Power (FRP)
21 observations from MODIS satellites into smoke constituents [Di Giuseppe et al., 2016]. Since CO
22 is chemically inert over the timescale of the FLEXPART calculations, these two tracers show the
23 relative importance and temporal evolution of these potential O₃ sources, but cannot be
24 quantitatively compared to the stratospheric O₃ flux since FLEXPART does not include
25 chemistry. The FLEXPART calculations used here were generously provided by S. Evan and J.
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1 **Appendix B: FAST-LVOS Daily Summaries**

2

3 Wednesday, May 17

4 TOPAZ arrived at NLVA around 13:00 PDT, the system was set up, the laser was realigned, and
5 data were taken from 17:37 - 20:57 PDT. Rather cool temperatures and strong winds out of the
6 NW. High ozone (80 - 120 ppbv) was observed above 4.5 km AGL. Boundary layer and surface
7 ozone peaked around 19:00 PDT at around 70 ppbv.

8 Thursday, May 18

9 TOPAZ data were taken from 09:37 - 21:13 PDT. The data gap from about 11:00 - 14:00 PDT was
10 due to an extensive laser system realignment. Cool temperatures throughout the day, low
11 clouds and strong winds from a NE/NW direction in the afternoon and early evening. Due to the
12 clouds and strong winds, surface and boundary layer ozone dropped to below 40 ppbv in the
13 afternoon. High ozone (80 - 120 ppbv) aloft descended during the day to about 2 km AGL, but
14 did not reach the surface.

15 Friday, May 19

16 TOPAZ data were taken from 10:39 - 18:23 PDT. Still below normal temperatures, light winds.
17 The high-ozone layer (75 - 95 ppbv, 4 - 6 km AGL) aloft persisted and thinned over the course of
18 the day. Surface and BL ozone did not vary much during the day and stayed below 60 ppbv.

19

20 Saturday, May 20

21 TOPAZ data were taken from 10:25 - 18:09 PDT. Seasonal temperatures with a high around 33
22 C, light winds. The high-ozone layer that had been observed the previous three days thinned
23 further and contained less ozone (~ 80 ppbv). It lowered to about 2.5 km AGL by 14:00 PDT and
24 then most likely moved away from the Las Vegas area. The mixed layer was not deep enough to
25 entrain this ozone layer. Surface and BL ozone peaked in the low 60s ppbv in the afternoon

26

27 Sunday, May 21

28 TOPAZ data were taken from 11:08 - 19:08 PDT. Warm temperatures with a high around 38 C,
29 light winds, and a period of Cu clouds in the late afternoon. Fairly uniform surface and BL ozone
30 throughout the day with a maximum of 70 ppbv around 17:00 PDT. Enhanced ozone around 6
31 km AGL.

32

33 Monday, May 22

34 TOPAZ data were taken from 09:14 - 19:14 PDT. Warm temperatures with a high around 38 C,
35 light winds. Surface and BL ozone peaked in the low 60s ppbv in the early evening. A moderate
36 ozone layer containing enhanced aerosols was observed between 2 and 4 km AGL until midday.
37 Enhanced ozone was measured again around 6 km AGL.

38 Tuesday, May 23

39 TOPAZ data were taken from 08:04 - 22:04 PDT. Hot temperatures with a high around 40 C,
40 light winds. Locally produced ozone built up in the afternoon reaching a maximum

1 concentration of about 80 ppbv at the surface and throughout the 2-km deep BL. The Clark
2 County DAQ site Walter Johnson recorded a MDA8 ozone value of 71 ppbv, which exceeded the
3 NAAQS. A high-ozone layer (80 - 90 ppbv) was observed above 6 km AGL.

4 Wednesday, May 24

5 TOPAZ data were taken from 07:45 - 03:53 PDT the next day. Hot temperatures with a high
6 around 39 C, light winds until midday, followed by a sharp increase in wind speeds to about 15
7 m/s. An ozone layer aloft (70 - 80 ppbv, 500 m - 4 km AGL) was entrained into the growing BL
8 around 11:00 PDT, resulting in a sharp increase in surface ozone to just above 70 ppbv.
9 However, this increase was short-lived as the high winds in the afternoon diluted ozone
10 concentrations. Surface and BL ozone stayed in the mid 60s ppbv throughout most of the day. A
11 secondary ozone maximum of approx. 70 ppbv was observed around midnight. High ozone
12 concentrations up to about 100 ppbv were measured above 6.5 km AGL.

13

14 Thursday, May 25

15 TOPAZ data were taken from 08:03 - 17:55 PDT. Warm temperatures with a high around 36 C,
16 light winds from the early morning hours until midday, followed by an increase in wind speeds
17 to about 20 m/s. Moderate ozone in the residual layer aloft was entrained into the growing BL
18 around 10:30 PDT, resulting in an increase in surface ozone to about 65 ppbv. The very strong
19 winds in the afternoon and evening diluted ozone concentrations, causing surface ozone to stay
20 in the 50s ppbv during the remainder of the day. A high-ozone layer (70 - 85 ppbv) was
21 observed above 4 km AGL after 11:00 PDT.

22

23 Friday, May 26

24 TOPAZ data were taken from 09:23 - 18:51 PDT. The TOPAZ laser needed to be realigned, hence
25 the data gap between 11:00 and 14:40 PDT. Seasonal temps (33 C max). Winds from the SW.
26 Breezy overnight and in the early evening. Surface ozone stayed elevated (40-55 ppb)
27 overnight, increased to about 65 ppb around midday, and then stayed level at about 60 ppb
28 into the night. In the afternoon, ozone was well mixed up to 3 km AGL. High ozone (80-90 ppb)
29 was observed above 6 km AGL.

30

31 Saturday, May 27

32 TOPAZ data were taken from 07:40 - 18:20 PDT. High pressure was moving in from the west
33 with a shortwave trough to the northeast. Seasonal temps (34 C), light winds, a few Cu clouds
34 over the nearby mountains. MicroDop lidar observed intermittent turbulence at night up to
35 about 3 km AGL. The convective daytime BL rose to about 2.5 km AGL in the afternoon. Surface
36 ozone stayed elevated (35-55 ppb) overnight and increased to about 60 ppb in the afternoon.
37 TOPAZ observed an ozone layer aloft (\approx 65 ppb) between 2.5 and 3.5 km AGL. That layer
38 thinned during the day and was most likely entrained into the BL in the early afternoon. There
39 was high ozone (\approx 80 ppb) above 4.5 km AGL associated with the shortwave trough.

40

41 Sunday, May 28

42 After TOPAZ maintenance in the morning, data were taken from 12:27 - 18:51 PDT. High
43 pressure in control, hot (36°C max), light to moderate winds, isolated Cu clouds over the nearby

1 mountains. The MicroDop lidar observed intermittent turbulence at night. The convective
2 daytime BL rose to about 2.5 km AGL in the afternoon. Surface ozone stayed somewhat
3 elevated (25-45 ppb) overnight and increased to the low 60s ppb in the afternoon. TOPAZ
4 observed high ozone above 4.5 km AGL (75 - 90 ppb).

5
6 Monday, May 29

7 TOPAZ data were taken from 10:40 - 22:08 PDT. Hot (39 C max), light winds, isolated Cu clouds
8 over the nearby mountains. Cirrus shield associated with the West Coast trough was moving in
9 from the SW. The BL rose to just over 4 km AGL in the afternoon, as evidenced by the MicroDop
10 lidar and the afternoon sounding at McCarren Airport. The deep BL entrained the bottom part
11 of the high ozone layer (80-90 ppb), that had been observed by TOPAZ between 4 and 7 km AGL
12 for several days in a row. The air in this layer was very dry, so the high ozone likely originated in
13 the UTLS. The entrained ozone mixed down to surface resulting in a several-ppb jump in ozone
14 concentration.

15
16 Tuesday, May 30

17 TOPAZ data were taken from 11:28 - 19:20 PDT. Laser maintenance in the morning and around
18 midday. TOPAZ performance was not optimal during the late AM data segment. Hot (38 C max),
19 light winds early, increasing in the afternoon. Thick Cirrus shield associated with the West Coast
20 trough. The BL rose to about 4 km AGL in the afternoon, just to the bottom of the high ozone
21 layer (80-90 ppb). Contrary to the day before, the high ozone aloft was not entrained into the
22 BL. Cirrus clouds at 7.5 km AGL caused artifacts in the ozone and aerosol retrievals.

23
24 Wednesday, May 31

25 TOPAZ data were recorded from 7:34 PDT through the night until 7:02 PDT on June 1 (data
26 collection was continued the next day without interruption). Slight cooldown compared to the
27 day before (35 C max), light winds early, increasing to 15 m/s in the afternoon and becoming
28 light again after sunset. Mid- and high-level clouds associated with the West Coast trough,
29 clearing out in the early morning hours the next day. Between 11 and 12 PDT, a sudden
30 increase in ozone by about 10 ppb at the surface and throughout the BL (up to 1.5 km AGL) was
31 observed. This bump up in ozone was likely associated with the advection of air in southerly
32 flow aloft that was influenced by fires burning in western Mexico. This was corroborated by the
33 in-situ observations at Angel Peak (enhanced NO_x, NO_y, CO, CO₂, and CH₄, but no increase in
34 O₃) and back trajectory calculations. The BL rose to about 3.5 km AGL in the late afternoon.
35 TOPAZ observed a persistent ozone layer (75 - 85 ppb) above 4 km AGL. Clouds above 6 km AGL
36 caused artifacts in the ozone and aerosol retrievals at times.

37
38 Thursday, June 1

39 TOPAZ data were recorded from 7:14 - 20:42 PDT. Seasonal temperatures (34°C max) after
40 trough passage, light winds early, increasing to about 15 m/s around midday, and becoming
41 light again after sunset. Isolated Cu clouds over the nearby mountain ranges, otherwise clear.
42 Ozone increased to the mid 60s ppb after the growing BL had entrained the residual layer. In
43 the late morning, the BL grew rapidly from 700 to 2700 m AGL and clean air above 1.5 km AGL
44 was entrained into the BL, which likely caused the ≈10 ppbv drop in ozone concentrations. The

1 strong southerly winds kept the area well ventilated and surface ozone stayed in the low to mid
2 50s ppb. The BL rose to about 3.5 km AGL in the late afternoon and entrained part of the
3 persistent ozone layer that had been observed near that altitude for several days. This was
4 likely the cause for the slight increase in ozone around 1700 PDT. A high-ozone layer (\approx 100
5 ppb) associated with a stratospheric intrusion on the back side of the departing trough was
6 observed above 5 km AGL, well above the top of the BL.

7
8 Friday, June 2

9 TOPAZ data were recorded from 7:02 - 23:10 PDT. First day of a long warming trend (36.5°C
10 max) with high pressure building from the west. Light winds throughout the day. Isolated Cu
11 clouds over the nearby mountain ranges, otherwise clear. The mixed layer was suppressed by a
12 strong 1.2 m deep inversion during the morning, but eventually reached about 3 km around
13 1500 PDT. The growing ML entrained the RL around noon, and entrained the lowest (3-3.5 km
14 agl) of two elevated ozone layers from yesterday's intrusion around 1500 PDT. The aerosol
15 measurements suggest that some of the higher layer may also have been mixed with Asian
16 pollution or fire plumes. A pronounced tongue of ozone subsided after the ML collapsed and
17 nearly reached the surface in the late evening. Surface ozone hovered around 70 ppbv at NLVA
18 throughout the afternoon and reached a maximum of nearly 80 ppbv at Angel Peak around
19 1900 PDT. The Walter Johnson monitor reached an hourly maximum of 74 ppbv and an MDA8
20 of 68 ppbv.

21
22 Saturday, June 3

23 TOPAZ data were recorded from 9:47 - 19:55 PDT. Very hot with an official max of 102 F at the
24 NLVA NWS station (42C or 107 F at TOPAZ). Very light southeasterly winds throughout the day
25 with broken cirrus and persistent contrails aloft. Mixing was initially suppressed by another
26 surface inversion, but surface ozone rose rapidly when the RL was entrained between 1200 and
27 1300 PDT. The monitor in TOPAZ peaked around 80 ppbv at this time, and the hourly mean
28 surface concentration at the Walter Johnson (WJ) monitor hit 78 ppbv. The MDA8 at WJ
29 reached 70.5 ppbv and thus did not quite exceed the standard. The winds at AP remained
30 southerly throughout the day and the mobile van did not see the high ozone measured over the
31 LVV. The ML at NLVA eventually reached 4 to 5 km deep in the afternoon, first entraining
32 cleaner air from between 2 and 4 km agl, and then a thin high ozone (and aerosol) layer of
33 possible Asian origin present around 4.5 km agl, and finally very clean air from above 5 km agl
34 in the late afternoon.

35
36 Sunday, June 4

37 TOPAZ data were recorded from 9:38 - 22:18 PDT. Very hot again with an official max of 102°F
38 at the NLVA NWS station. Clear skies with a few scattered Cu over the mountains. Very light
39 northwesterly winds in the early morning that shifted abruptly to strong (15 m/s mean)
40 southwesterly winds associated with a trough moving into the PNW after 1030 PDT. Surface
41 ozone at the NLVA rose quickly to about 65 ppbv before the wind shift, then fell to less than 50
42 ppbv, after which it slowly ramped back up to 65 ppbv. The strong wind kept a constant dust
43 layer in motion near the surface, and contributed to very deep mixing to more than 4 km. The
44 ozone concentrations were 50-60 ppbv up to about 4 km agl in the morning, with less than 40

1 ppbv between 4.5 and 5.5 km agl. A high ozone layer corresponding to a dry air streamer in the
2 GOES WV imagery became prominent around 4 km asl at 1400 PDT, and was slowly nibbled
3 from below by the deep mixed layer. A second layer with more than 100 ppbv of ozone (and
4 more aerosol) descended and merged with the first layer around 1900 PDT. Both the surface
5 winds and ozone subsided briefly between 1900 and 2000 PDT, but then shifted to the west
6 and increased again. The turbulent mixing aloft continued until after 2100 PDT.

7
8 Monday, June 5

9 TOPAZ data were recorded from 9:46 - 18:10 PDT. Hot once again with a max of 101°F at the
10 NLVA (VGT) NWS station. Clear skies with a few scattered Cu over the mountains. Strong
11 southwesterly winds persisted throughout the night (there was an apparent sensor malfunction
12 around 0300 PDT) before calming between 0800 and 1200 PDT. The winds prevented surface
13 deposition and in situ ozone remained relatively high except for a period of local titration
14 around 0600 PDT. TOPAZ showed the high ozone associated with the streamers observed
15 yesterday to still be present. Some of this ozone was mixed down to the surface when the ML
16 rapidly grew around 1100 PDT, causing a short-lived increase in the surface concentrations to
17 more than 70 ppbv. High O3/Low CO UT/LS air also descended to Angel Peak around 0300 PDT
18 and later in the morning. The MLH was limited to 3.2 km by the stable cap of the streamer, and
19 the deep SW flow (from the surface to the stratosphere) advected relatively clean air (50-55
20 ppbv) to Las Vegas beneath the streamer. This air offset any ozone increase caused by
21 entrainment of the overlying streamer, and the in-situ concentrations at NLVA dropped to 55
22 ppbv when the 10-15 m/s SW winds resumed at noon, and remained below 65 ppbv for the rest
23 of the day.

24
25 Tuesday, June 6

26 TOPAZ data were recorded from 9:49 - 20:21 PDT. Another hot day with a max of 103°F at the
27 nearby NLVA (VGT) NWS station. Clear skies in the morning followed by widely dispersed
28 contrail cirrus. The surface winds calmed during the night and ozone decreased to a locally
29 titrated minimum between 5 and 10 ppbv between 0600 and 0700 PDT. The southwesterly
30 winds gradually increased during the late morning with a mean speed of 12 m/s during the
31 afternoon. TOPAZ showed the Pacific streamer to still be above Las Vegas between 3.5 and 4
32 km. The streamer was lifted up and disrupted by the mixed layer, which reached a height of
33 about 3.5 km during the afternoon, but recovered most of its ozone density as the ML subsided
34 in the early evening. A substantial amount of ozone was entrained into the ML, but not enough
35 to raise the surface concentrations above 70 ppbv.

36
37 Wednesday, June 7

38 TOPAZ data were recorded from 9:52 - 22:24 PDT. Another hot day with a max of 101°F at the
39 nearby (500 m) NLVA (VGT) NWS station. Hazy with widely dispersed contrail cirrus all day. The
40 winds were from the S and SE for most of the day before rotating to the SSW around 1500 PDT.
41 Winds between 5 and 10 m/s prevented the accumulation of locally produced ozone. A layer of
42 high ozone was present just above 2 km when the measurements commenced in the morning,
43 with the bottom of the layer rising to 3.5 km agl by the late afternoon. The concentrations at 3
44 km agl exceeded 100 ppbv from 1400 to 1600 PDT, when some of the ozone was entrained by

1 the mixed layer, causing the surface concentrations to rapidly increase from 65 to 75 ppbv. The
2 TOPAZ in situ concentrations remained above 70 ppbv until about 2000 PDT, causing the MDA8
3 to reach about 68 ppbv. The MDA8 at 5 of the CCDAQ monitors reached 65 ppbv, and the
4 monitor at Death Valley NP, 150 km WNW of Joe Neal and 120 km W of AP, reached 70 ppbv.
5 High ozone was also measured at Angel Peak. The measurements there indicate that the
6 composition of the elevated layer was complex. There was no visible smoke, but the aerosol
7 backscatter was somewhat enhanced. The O₃ at Angel Peak was sometimes correlated with
8 NO_y and CO, but not at other times. It was usually correlated with CH₄, however. Stratospheric
9 and (AZ?) fire contributions are suspected.

10
11 Thursday, June 8

12 TOPAZ data were recorded from 9:57 - 22:45 PDT. Slightly cooler today with a max of 99°F at
13 the NLVA(VGT) NWS station. There were thick high clouds and winds ahead of the advancing
14 trough. Surface ozone remained above 60 ppbv from 0900 to 2100 PDT. Surface ozone
15 concentrations at AP exceeded 70 ppbv for much of the day with different degrees of
16 correlation with CO and other tracers. The surface winds were from the S to SSW at 10-15 m/s
17 throughout this period (the CCDAQ wind profiler showed similar winds to at least 3 km agl).
18 This implies that most of the ozone was advected into the LVV, but photochemical production
19 and accumulation was also depressed in the SoCAB. TOPAZ once again showed a high (80-90
20 ppbv) ozone layer between 3 and 4 km agl with relatively high concentrations (65-75 ppbv)
21 between this layer and the surface. The microDop showed the ML reached about 2500 m agl.
22 Rapids periods of deep mixing around 1200 and 1900 PDT increased the surface concentrations
23 by 5 to 10 ppbv. Similar increases were mirrored by the Apex CCDAQ monitor and several other
24 CCDAQ monitors peaked during the evening episode. The MDA8 at most of the CCDAQ
25 monitors reached at least 60 ppbv. As was the case yesterday, a combination of aged UT/LS air
26 and other sources is suspected

27
28 Friday, June 9

29 TOPAZ data were recorded from 9:50 - 18:54 PDT. Cooler today with a max of 95°F at the NLVA
30 (VGT) NWS station. Mostly clear with widely dispersed contrails. Strong SW to SSW winds
31 greater than 20 m/s with gusts exceeding 25 m/s swept away locally produced O₃ and created a
32 wind-driven dust cloud over the Strip. The surface O₃ concentrations were about 40 ppbv in the
33 morning, and the remnants of the Pacific streamer were present when TOPAZ measurements
34 commenced at 10 PDT. The first lidar measurements showed more than 100 ppbv of O₃
35 between 2.5 and 3 km agl, but the streamer was highly fragmented with much lower
36 concentrations during most of the day. A modest amount of entrainment and mixing occurred
37 after 1300 PDT, but the surface concentrations remained below 55 ppbv, and only one of the
38 CCDAQ monitors (Mesquite) reached 60 ppbv for the MDA8. As noted before, a combination of
39 aged UT/LS air and transported biomass burning or pollution sources is suspected. Despite the
40 location of the streamer near the elevation of Angel Peak, only traces of UT/LS air were
41 detected there.

42
43 Saturday, June 10

44 TOPAZ data were recorded from 08:32 to 21:36 PDT. Cooldown prior to the arrival of the deep

1 trough continued with a max temperature of 32 C. After a clear morning, a thick Cirrus cloud
2 deck moved in from the SW around midday and cleared out of the area after sunset. Strong SW
3 to SSW winds from 15 to 20 m/s with higher gusts swept away locally produced O3. The Pacific
4 streamer, with ozone concentrations exceeding 100 ppb at times, reappeared at 10 PDT and
5 persisted into the evening. Compared to the previous days, the streamer descended to lower
6 altitudes (1.5 - 3 km AGL) but stayed mostly above the mixed layer with only minimal
7 entrainment in the late afternoon and evening, which caused a slight increase in surface ozone
8 concentrations. The high-ozone streamer was also sampled by the ozonesondes launched at
9 Joe Neal at 9 and 12 PDT, the Mooney aircraft, and the mobile van at Angel Peak. The air above
10 the streamer up to about 5 km AGL was very clean with ozone concentrations around 30 ppb
11 and very little aerosol. The Cirrus clouds above 7 km AGL caused a few artifacts in the TOPAZ
12 ozone and aerosol data between 16 and 20 PDT.

13

14 Sunday, June 11

15 TOPAZ data were recorded from 8:03 PDT through the night until 7:39 PDT on June 12 (data
16 collection was continued the next day without interruption). Cooldown prior to the arrival of
17 the deep trough continued with a max temperature around 30 C. Clear throughout the day.
18 Strong S to SW winds from 15 to 25 m/s with higher gusts swept away any locally produced O3.
19 Several layers of high ozone aloft were observed, including a tongue of very high ozone (peak
20 values around 280 ppb near 5.5 km AGL) associated with a strong stratospheric intrusion as the
21 trough axis passed over Las Vegas in the early morning hours on June 12. The high ozone did
22 not get transported down to the surface, except in the afternoon on June 11 when entrainment
23 by the mixed layer pushed surface ozone up to about 70 ppb.

24

25 Monday, June 12

26 TOPAZ data were recorded from 7:44 PDT through the night until 6:32 PDT on June 13 (data
27 collection was continued the next day without interruption). Cool temperatures (max. of 29°C)
28 and clear conditions. Decreasing winds in the morning and then light winds veering from SW to
29 NW during the remainder of the day in the wake of the departing trough. Very high ozone was
30 observed above 5 km AGL (max. ozone values near 280 ppb in the morning with several ozone
31 filaments descending to about 3 km AGL. The ML was not deep enough to entrain any of this
32 ozone. Surface ozone stayed below 50 ppb in the cool post-frontal conditions.

33

34 Tuesday, June 13

35 TOPAZ data were recorded from 6:44 - 21:00 PDT. These measurements were a continuation of
36 the long-time series that started at 0800 PDT on June 11. Warming temperatures (max. of 34
37 C), clear conditions, and light to moderate winds from the NW. Highly structured ozone aloft
38 with multiple thin layers associated with yesterday's intrusion. High ozone descended to below
39 3 km agl around midday, but the the mixed layer only reached about 2200 m agl and was too
40 shallow to entrain much of it. The surface concentrations reached about 65 ppbv during the
41 afternoon with similar concentrations at the CC monitors.

42

43 Wednesday, June 14

44 TOPAZ data were recorded from 10:46 - 21:02 PDT. Warm temperatures (max. of 39°C), clear

1 conditions, and light winds that followed the diurnal flow pattern. The mixed layer only reached
2 about 1800 m agl. It entrained a thin ozone layer aloft containing moderate ozone
3 concentrations (≈ 75 ppb) in the early afternoon and then slowly declined in depth through the
4 afternoon. The warm temperatures, light winds, and the shallow boundary layer caused a
5 buildup of ozone that reached 83 ppb at the surface in the late afternoon. MDA8 ozone
6 reached 74 ppb.

7 Thursday, June 15

8 TOPAZ data were recorded from 09:19 - 19:51 PDT. Hot (max. of 41°C), clear conditions, except
9 some high clouds to the north in the evening. Light winds that followed the diurnal flow pattern
10 thru mid-afternoon, then a wind shift to the S and then W with wind speeds increasing to ≈ 8
11 m/s. The mixed layer was very shallow until midday and then grew quite rapidly to a depth of
12 2800 m agl. Surface and boundary layer ozone increased to almost 80 ppb around 14:00 PDT,
13 dropped down to around 60 ppb in the middle of the afternoon, and then increased again at
14 16:30 to around 70 ppb, before dropping below 60 ppb at sunset.

15 Friday, June 16

16 TOPAZ data were recorded from 09:05 - 20:33 PDT. Very hot (max. of 43°C), clear conditions,
17 and light winds that followed the diurnal flow pattern. The mixed layer gradually grew to 3300
18 m agl. It entrained a thin ozone layer aloft containing moderate ozone concentrations (≈ 70 ppb)
19 in the early afternoon. The hot temperatures and light winds caused a buildup of ozone that
20 reached 87 ppb at the surface in the afternoon. MDA8 ozone reached 76 ppb.

21 Saturday, June 17

22 TOPAZ data were recorded from 08:56 - 18:48 PDT. Very hot (max. of 44°C). Clear conditions,
23 except for a few Cirrus clouds to the north in the afternoon. Generally light winds that followed
24 the diurnal flow pattern. Wind speeds increased to 5 - 10 m/s in the afternoon and early
25 evening. The mixed layer stayed shallow through 1200 PDT and then grew to about 3000 m agl
26 by 1500 PDT. Surface ozone quickly increased to 95 ppb by midday. The growing mixed layer
27 entrained cleaner air from aloft, which caused surface ozone to drop in a series of steps down
28 to the mid 60s ppb by 1600 PDT. Despite this afternoon decrease, MDA8 ozone at three
29 monitoring sites in the Las Vegas Valley exceeded the 70 ppbv standard.

30 Sunday, June 18

31 TOPAZ data were recorded from 09:31 - 19:07 PDT. Extremely hot (max. of 46°C at the lidar
32 site). Shallow Cumulus over the nearby mountain ranges in the afternoon. Light winds that
33 followed the diurnal flow pattern. Overnight, clean marine air had been advected into southern
34 Nevada, which decreased ozone concentrations to below 40 ppb in the lower troposphere.
35 TOPAZ observed a layer with moderate aerosol concentrations and 45-50 ppb ozone between 3
36 and 5 km AGL. Local production caused a late increase of surface ozone, which peaked at 53
37 ppb around 1800 PDT.

38

39 Monday, June 19

40 TOPAZ data were recorded from 09:23 - 18:43 PDT. Extremely hot (max. of 47°C at the lidar

1 site). Generally light winds that followed the diurnal flow pattern. Cumulus clouds and isolated
2 t-storms over the nearby mountain ranges. A few clouds drifted over the Las Vegas Valley in the
3 late afternoon. The clean marine air from the previous day with lower tropospheric ozone
4 concentrations of less than 40 ppb was still in place. TOPAZ observed a layer with moderate to
5 high aerosol concentrations and 55-60 ppb ozone above 4 km AGL. This layer descended to 2.5
6 km AGL by late afternoon and was partially entrained into the boundary layer. Back trajectory
7 calculations suggest that this layer may be associated with fires in Arizona. Surface ozone
8 peaked at 55 ppb around 1730 PDT.
9

10 Tuesday, June 20

11 TOPAZ data were recorded from 09:54 - 22:02 PDT. Extremely hot again (max. of 48°C at the
12 lidar site) with the all-time Las Vegas record of 117 F tied at McCarran Intl Airport. Scattered
13 fair weather cumulus clouds formed over the nearby mountain ranges with a few drifting over
14 the Las Vegas Valley in the late afternoon. Very light NW winds in the early morning that
15 rotated CCW to the east between 0900 and 1500 PDT. Surface ozone at NLVA increased rapidly
16 during this time and reaching 84 ppb around 1300 PDT; the 1 h mean at Walter Johnson peaked
17 at 78 ppbv between 1300 and 1400 PDT. The mixed layer grew explosively to more than 5 km
18 deep around 1500 PDT and the surface winds abruptly increased to nearly 15 m/s, rotating first
19 to the NW and then to the SW. This caused the surface ozone concentrations to drop to about
20 65 ppbv where they remained with very little variation into the early evening. The SW winds
21 blew the ozone towards Apex, which measured the hourly concentrations of 75 to 76 ppbv
22 during the late afternoon. The initial TOPAZ measurements showed the high aerosol/ozone
23 layer attributed to the AZ and MEX fires and seen the day before near 3.5 km asl, but with
24 much high aerosol loadings and 70-80 ppb of ozone. This layer was apparently entrained as the
25 ML deepened after 1400 PDT, but the resulting dilution, mixing, and venting decreased and
26 homogenized the concentrations to an average value of around 64 ppbv. The 00 UT (17 PDT)
27 VEF sounding put the top of the ML at more than 6 km asl (about 5.5 km agl), in good
28 agreement with the afternoon TOPAZ ozone and aerosol measurements. The ML did not begin
29 subsiding until after 2200 PDT.
30

31 Wednesday, June 21

32 TOPAZ data were recorded from 09:20 - 19:20 PDT. Another extremely hot day (max. of 45°C at
33 the lidar site) and 113 F at the NLVA NWS tower. Very clear with only a few FWC over the
34 mountains. The elevated smoke layer from AZ and Mexico was visible to the eye and prominent
35 in the TOPAZ aerosol backscatter between 3 and 5 km agl. There was very little ozone
36 enhancement in the smoke layer, however, and modest entrainment of the smoke by the
37 mixed layer, which reached a depth of 'only' about 4 km. The ML subsided around 1700 PDT.
38 The winds were weak and NW in the early morning, but rotated to the S and increased to 10-15
39 m/s around 1000 PDT. Surface ozone also increased to about 65 ppbv, but then decreased to
40 55-60 ppbv as the winds rotated to the SW. TOPAZ measured similar concentrations up to
41 about 5 km agl. There was no sign of smoke from the Holcomb Fire near San Bernardino.
42

43 Thursday, June 22

44 TOPAZ data were recorded from 09:10 - 20:06 PDT. Another extremely hot day (max. of 45°C at

1 the lidar site) and 113 F at the NLVA NWS tower. Very clear with only a very few clouds over the
2 mountains. The winds were light during most of the day, shifting from northwesterly
3 downslope flow to very weak southerly flow around 0900 PDT. The elevated smoke layer seen
4 the previous two days was mostly gone, with only two thin and weak aerosol layers (but no
5 ozone) remaining near 4 and 5 km agl. A thin, but much more concentrated aerosol layer was
6 present at the surface in the morning. Ozone began to increase early in the lowest 1 km, rising
7 from about 50 to more than 80 ppbv between 0900 and 1100 PDT. Another layer of enhanced
8 ozone and aerosol was present between 2.5 and 3 km agl and intensified around 1100 PDT.
9 Surface O₃, CO, and CH₄ began increasing at AP around the same time, and HYSPLIT back
10 trajectories suggest that this layer originated from the Holcomb Fire near Big Bear Lake
11 although there was no obvious smoke associated with the fire plume. This layer was entrained
12 by the growing mixed layer between 1300 and 1400 PDT, contributing to the high surface
13 ozone. The surface and ML column O₃ decreased around 1500 PDT as the growing ML
14 entrained lower O₃ air from above 3 km agl. A brief shift in the surface winds to the north
15 caused a short-lived drop in the O₃ and aerosol concentrations below 1 km around 1600 PDT,
16 but the concentrations recovered and the ML grew to more than 4 km agl. Ozone fell to about
17 60 ppbv throughout the column after the winds shifted to the west and increased to more than
18 10 m/s around 1800 PDT. The MDA₈ was greater than 75 ppbv at the TOPAZ truck, and reached
19 values of 77,73,73, and 70 at the Joe Neal, Palo Verde, Walter Johnson, and Apex monitoring
20 sites, respectively.

21
22 Friday, June 23

23 TOPAZ data were recorded from 09:29 - 20:57 PDT. Another extremely hot day (max. of 45°C at
24 the lidar site) and 112 F at the NLVA NWS tower. Very clear with a few scattered FWC in the
25 late afternoon. Calm for most of the day with very light (5 m/s) NW winds before 0900 and W
26 after 1800 PDT. There was no evidence of elevated aerosol or ozone layers from either the aged
27 AZ and Mexican Fires, or from the Holcomb Fire aloft, and the in situ ozone concentrations
28 remained below 65 ppbv at Angel Peak. Surface ozone at NLVA was about 40 ppbv at 0900 PDT,
29 but increased monotonically to a peak of more than 100 ppbv around 1600 PDT. TOPAZ found
30 similar concentrations throughout the mixed layer, which remained less than 1500 m deep until
31 after 1400 PDT. The ML then developed a series of short-lived 4 km deep plumes that persisted
32 until nearly 1900 PDT, and modulated the aerosol and ozone concentrations above the surface.
33 The ML abruptly collapsed around 1900 PDT, leading to a rapid drop in ozone from the surface
34 to 2 km agl as the surface winds shifted to the west and increased to 5 m/s. The MDA₈
35 calculated from the 1-min running average ozone was greater than 80 ppbv at the TOPAZ truck,
36 and the MDA₈ reached values of 77,76,78, and 75 at the Joe Neal, Palo Verde, Walter Johnson,
37 and JD Smith monitoring sites, respectively. Note: The strange ozone and aerosol profiles
38 around 1300 PDT are spurious.

39
40 Saturday, June 24

41 TOPAZ data were recorded from 09:45 - 20:13 PDT. Another extremely hot day (max. of 45°C at
42 the lidar site) and 111 F at the NLVA NWS tower. Mostly clear with scattered high clouds and
43 some cumulus development to the SW and over the mountains. Calm or very light winds all
44 day. The elevated aerosol and ozone layer from the AZ and Mexican fires was present once

1 again between 4 and 6 km asl, but was only weakly entrained by the relatively shallow mixed
2 layer. Ozone production was modest with peak concentrations of 65 ppbv, despite very similar
3 conditions as yesterday when surface ozone reached 100 ppbv. This Friday-Saturday
4 combination could be a good candidate for a modeling study.

5

6 Sunday, June 25

7 TOPAZ data were recorded from 09:36 - 18:56 PDT. Another extremely hot day (max. of 45°C at
8 the lidar site) and 111 F at the NLVA NWS tower. Scattered high clouds in the morning, followed
9 by mostly clear skies in the afternoon. Calm or very light winds all day. Very similar to yesterday
10 with modest ozone production and peak concentrations of about 70 ppbv. The mixed layer
11 grew to more than 4 km around 1500 PDT, which helped to dilute the ozone concentrations.
12 This Friday-Saturday-Sunday combination will be a good candidate for a modeling study.

13

14 Monday, June 26

15 TOPAZ data were recorded from 09:49 - 17:49 PDT. Slightly cooler today (max. of 42°C at the
16 lidar site) and 108 F at the NLVA NWS tower. Clear skies all day, but a blowing dust cloud to the
17 south driven by the 10-15 m/s (in the afternoon) SW winds associated with the approaching
18 trough. Some entrainment of ozone from above 4 km by the mixed layer, but little
19 accumulation near the surface where the peak mixing ratios remained less than 60 ppbv.

20

21 Tuesday, June 27

22 TOPAZ data were recorded from 09:17 - 20:13 PDT. Normal temperatures today (max. of 42°C
23 at the lidar site) and 106 F at the NLVA NWS tower. Clear skies all day, with some suspended
24 dust in the lowest 500 m during the morning. A thin aerosol layer (with no ozone enhancement)
25 of possible fire origin was also present around 2 km agl in the morning. The normal morning
26 NW winds were light, but the afternoon SW winds increased to about 10 m/s. A layer of high
27 ozone with very low aerosol of probable UT/LS origin was present above 4.5 km agl for much of
28 the day, but disappeared around 1600 PDT with no obvious entrainment by the ML, which only
29 reached to about 3.6 km agl. Surface ozone peaked at about 65 ppbv around noon, but
30 decreased as the ML expanded and the surface winds increased.

31

32 Wednesday, June 28

33 TOPAZ data were recorded from 09:19 - 22:47 PDT. Normal temperatures today (max. of 42°C
34 at the lidar site) and 106 F at the NLVA NWS tower. Severe clear skies all day, with haze/smoke
35 below 2 km that was transported from the SJV during the night. The winds were generally light,
36 but followed the usual pattern of morning northwesterly winds, weak mid-day southwesterly
37 winds, and stronger (10 m/s) evening westerlies. The nocturnal low-level transport from the SJV
38 also led to relatively high ozone at Jean and several other CC sites during the night. A residual
39 layer containing this transported ozone was entrained by 1000 PDT, increasing the surface
40 concentrations at NLVA to about 65 ppbv. A narrow layer of predominantly UT/LS origin with
41 high (100-120 ppbv) ozone, and some aerosol consistent with mixing of pollution or fire plumes
42 from Asia, descended from 4 km asl at 0920 PDT to 3 km asl at 1400 PDT, when it was entrained
43 by the ML, which grew to about 3.6 km agl. This helped increase the surface concentrations at
44 NLVA to a peak of about 80 ppbv. Peak hourly concentrations in excess of 70 ppbv were

1 measured at 8 of the CC regulatory monitors, with the 2015 NAAQS of 70 ppbv exceeded at Joe
2 Neal (74 ppbv) and Apex (71 ppbv).

3
4 Thursday, June 29

5 TOPAZ data were recorded from 09:21 - 20:41 PDT. High-normal temperatures today (max. of
6 42°C at the lidar site) and 106 F at the NLVA NWS tower. Clear skies all day. The winds followed
7 the usual pattern of morning northwesterly winds, and weak mid-day southwesterly winds, but
8 with 10-15 m/s ESE winds in the late afternoon and evening. Thick haze/smoke near the surface
9 in the morning. HYSPLIT suggests this haze may have come from fires in northwestern Nevada
10 or Northern California. Possible smoke from Arizona or the Brian Head fire in Utah after ESE
11 winds of 10-15 m/s kicked up at 1800 PDT. The elevated UT/LS layer advected from the Pacific
12 Ocean was still present above 4 km asl, but was thicker and more diffuse with 80-90 ppbv of
13 ozone. A probable Asian pollution layer with lower ozone, but higher aerosol lay just beneath it.
14 This lower layer was mostly entrained and the upper layer partially entrained by the mixed
15 layer, which grew to about 3.2 km. The surface concentrations at NLVA peaked at 75-80 ppbv,
16 and the nearby Joe Neal monitor recorded peak hourly concentrations of 77 and MDA8 of 70
17 ppbv.

18
19 Friday, June 30

20 TOPAZ data were recorded from 09:07 - 20:31 PDT. The final day of FAST-LVOS measurements.
21 High-normal temperatures again today (max. of 42°C at the lidar site) and 106 F at the NLVA
22 NWS tower. Clear skies all day. The winds followed the usual pattern of morning northwesterly
23 winds, and weak mid-day southwesterly winds, but were very light all day, allowing ozone to
24 accumulate near the surface. Thick haze/smoke near the surface in the morning once again. As
25 before, HYSPLIT suggests this haze may have come from fires in northwestern Nevada or
26 Northern California. The elevated UT/LS layer advected from the Pacific Ocean was still present
27 above 5 km asl in the morning, but was thicker and more diffuse with 80-90 ppbv of ozone.
28 Another UT/LS layer with higher concentrations lay above it and a diffuse Asian pollution layer
29 with lower ozone, but higher aerosol still lay beneath it. This lower layer was mostly entrained
30 by the ML, which grew to about 3 km. The surface concentrations at NLVA peaked at 90 ppbv,
31 and the CCDAQ monitors at Joe Neal, Walter Johnson, and Palo Verde recorded MDA8
32 concentrations in excess of 70 ppbv.

33
34
35

1 **Appendix C: Angel Peak Summary Plots**

2
3 This first section of this Appendix shows time series and scatter plots of the O₃, CO, CH₄, H₂O,
4 NO_y, N₂O, wind speed and direction measurements from the mobile laboratory on Angel Peak
5 on or before high O₃ days in Clark County. The second section shows weekly summary plots.
6

7 The complete mobile laboratory data set is available at:

8 [https://esrl.noaa.gov/csd/groups/csd7/measurements/csd_mobilelab/MobileLabLVOS/DataDo](https://esrl.noaa.gov/csd/groups/csd7/measurements/csd_mobilelab/MobileLabLVOS/DataDownload/)
9 [wnload/](https://esrl.noaa.gov/csd/groups/csd7/measurements/csd_mobilelab/MobileLabLVOS/DataDownload/)

15 **1. Angel Peak measurements on high ozone days**

16
17 **Table 7-1** shows the 2015 NAAQS of 70 ppbv was equaled or exceeded by one or more of the
18 Clark County regulatory monitors on 10 of the 45 FAST-LVOS measurement days: May 23, June
19 3, June 14, June 16-17, June 22-23, and June 28-30. The 2B O₃ measurements from the TOPAZ
20 truck at the NLVA also exceeded the NAAQS on all 10 of these days, and the CRDS
21 measurements from the NOAA/ESRL mobile laboratory on Angel Peak exceeded the NAAQS on
22 June 14, 16, 29, and 30. The correlation among the AP measurements during these high O₃
23 episodes provides some insight into the origin of the high O₃.
24

25 The first figure in each section shows the MDA8 ozone distribution on the high O₃ day(s). The
26 monitors are colorized using the U.S. EPA AQI scale. The NOAA monitors at Angel Peak and the
27 NLVA are represented by triangles.
28

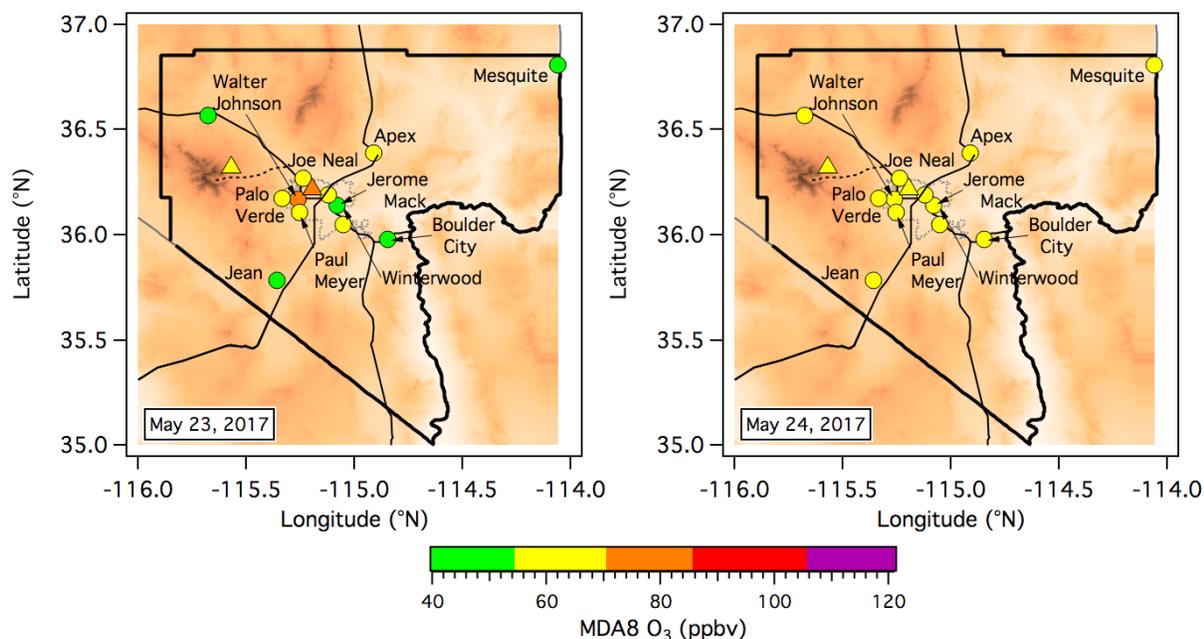
29 The second set of figures shows a 2 or 3-day time series plot(s) of the mobile laboratory in-situ
30 measurements of (a) O₃ and CO, (b) CO₂ and CH₄, (c) NO_y and N₂O, (d) T and RH, and (e) wind
31 speed and direction from Angel Peak. The horizontal dashed and dotted lines in the O₃ plots
32 correspond to 60 and 70 ppbv, respectively. The O₃ measurements made between 1200 and
33 2300 PDT are highlighted by filled circles.
34

35 The third set of figures shows scatter plots of the Angel Peak (a) zonal and meridional winds, (b)
36 O₃-CO, (c) O₃-CH₄, (d) O₃-N₂O, (e) O₃-NO₂, and (f) O₃-NO_y. The O₃-CO₂ relationship is
37 complicated by the drawdown of CO₂ by local vegetation and are not shown. The gray points
38 show all of the FAST-LVOS measurements, and the points colorized by the H₂O concentrations
39 correspond to the 1200-2300 PDT measurements highlighted in the O₃ time series plots.

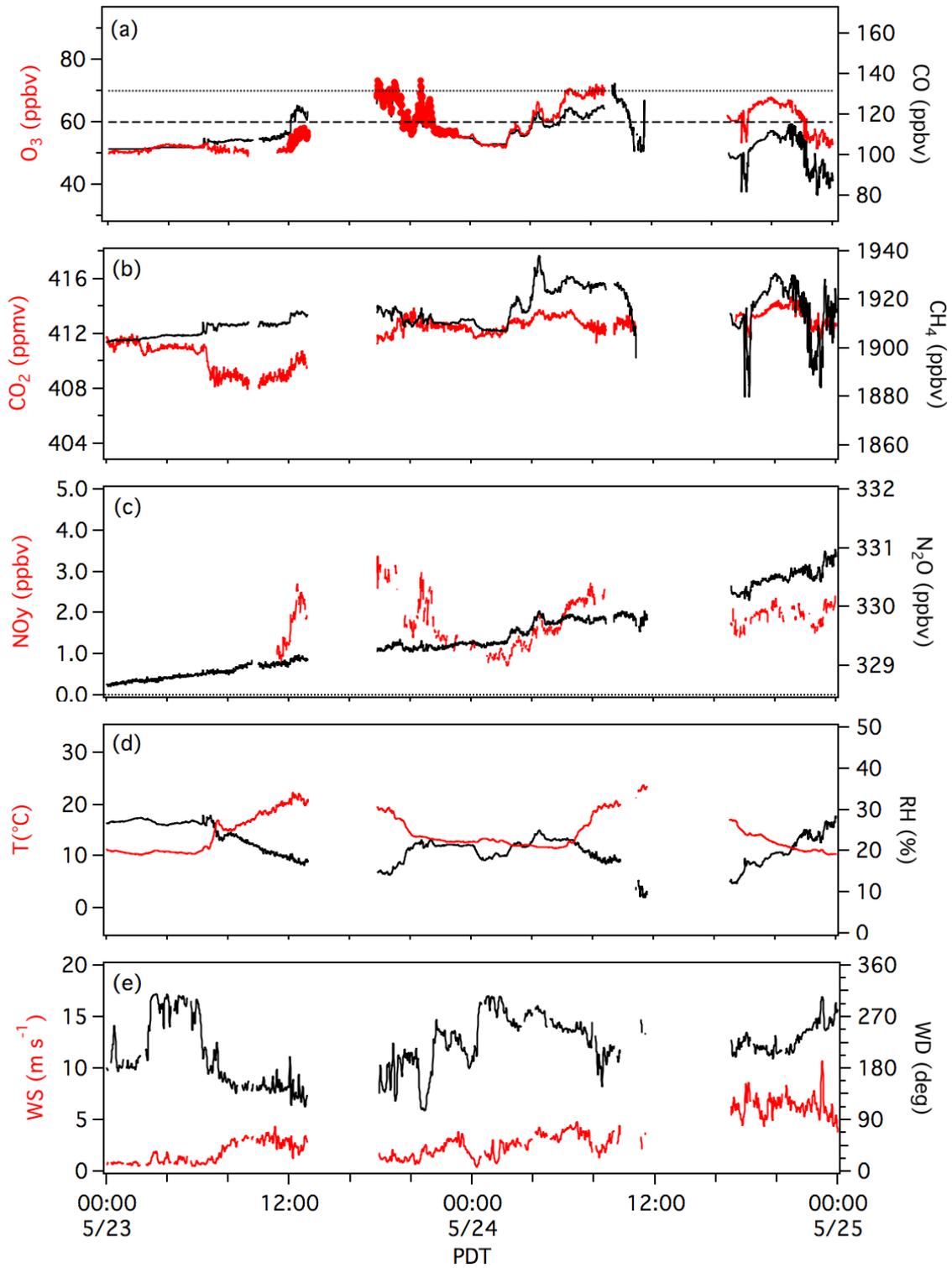
1
2 **May 23-24**
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4 The first ozone NAAQS exceedance in Clark County during 2017 occurred on May 23 when the
5 Walter Johnson monitor (C71) recorded an MDA8 concentration of 71 ppbv. MDA8 O₃ of 65
6 ppbv or more was also measured at 3 other nearby monitors: Joe Neal, Palo Verde, and Paul
7 Meyer. The TOPAZ monitor at the NLVA recorded an MDA8 of 72 ppbv. **Figure C-1** shows that
8 the high O₃ on the 23rd was confined to the northwest Las Vegas area. Ozone was moderate
9 (60-65 ppbv) at all of the LVV monitors on the 24th.

10
11 The NOAA mobile laboratory spent the early afternoon driving between Angel Peak and the
12 LVV, but measured an 8-h average of 64 ppbv after it returned to Angel Peak around 1745 PDT
13 (**Figure C-2**). The TOPAZ measurements showed a layer of high O₃ above 5 km asl on both days,
14 but this O₃ remained well above the elevation (2.7 km asl) of Angel Peak. The mobile lab
15 measurements show that the high O₃ measured there on the 23rd was transported up to the
16 summit from the LVV by moist southeasterly upslope flow and strongly correlated with CO, CH₄,
17 and NO_y (**Figure C-3**).
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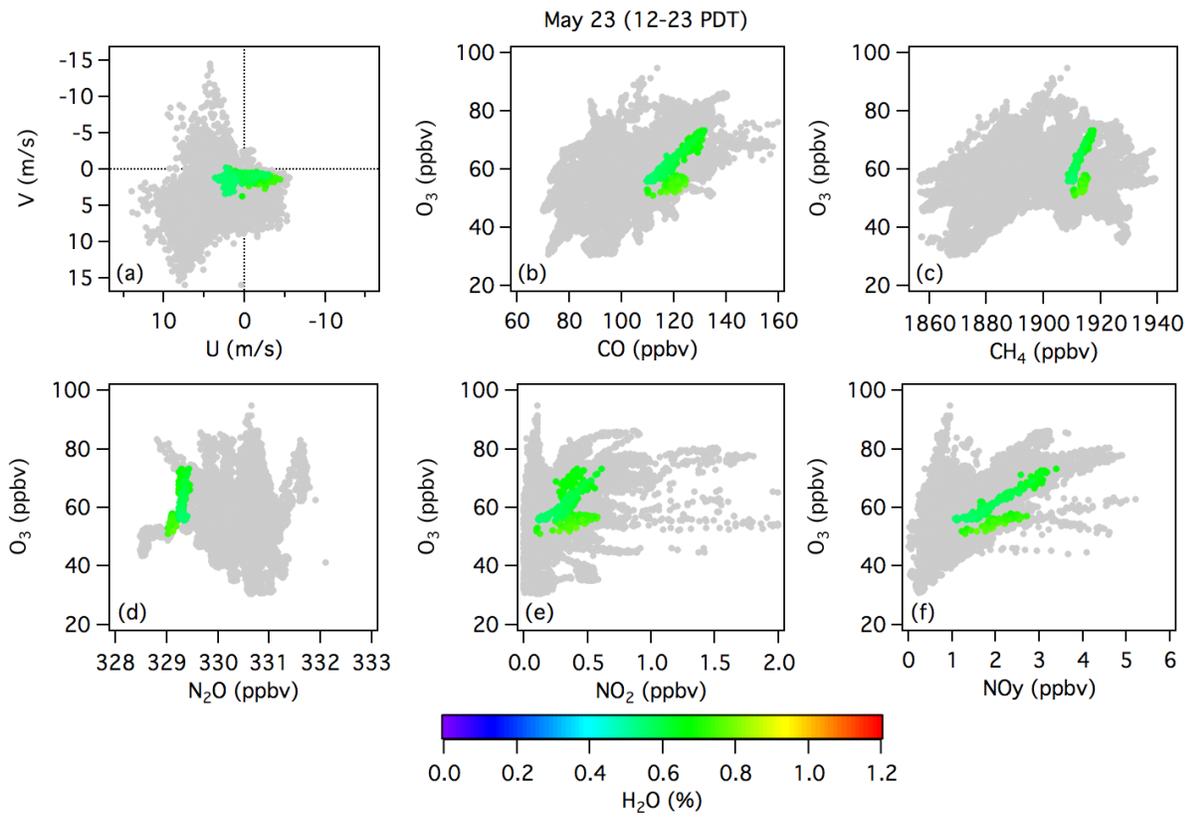


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23 **Figure C-1.** Topographic maps of Clark County showing the May 23-24 MDA8 O₃ distributions.
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2 **Figure C-2.** Time series of the Angel Peak mobile laboratory measurements on May 23-24.
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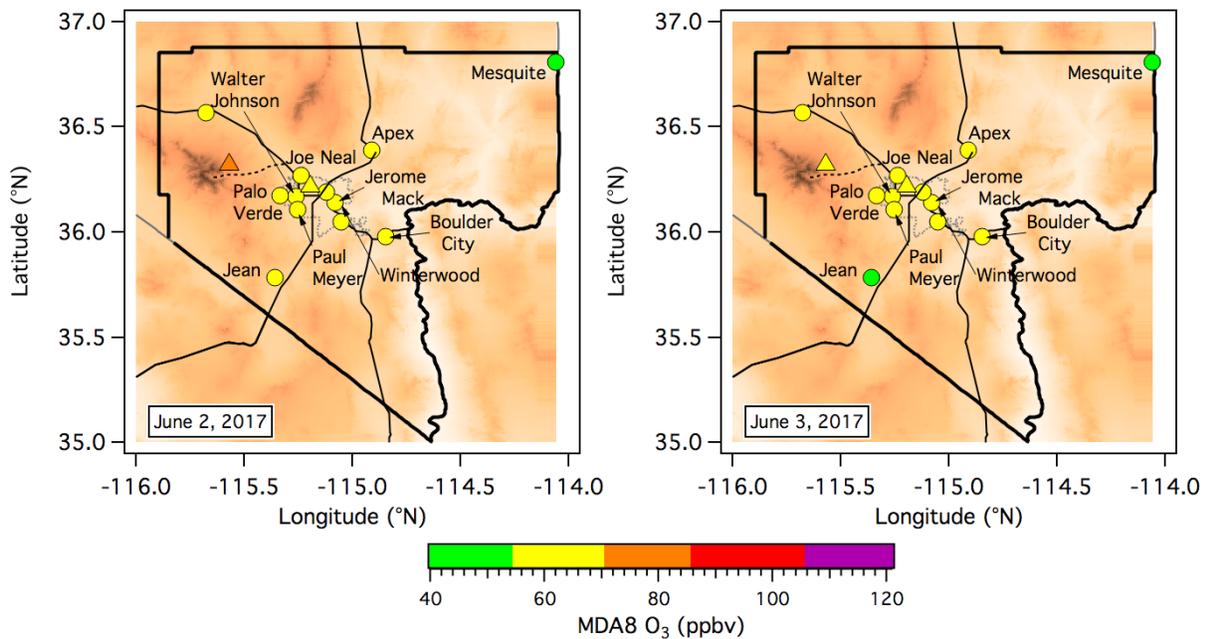
Figure C-3. Scatter plots showing the relationships between O₃ and the other parameters measured by the mobile laboratory on May 23.

1 **June 2-3**

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3 Moderate to high O₃ was measured by most of the CCDAQ monitors on June 2-3 with MDA8 O₃
4 concentrations in excess of 65 ppbv recorded by 3 monitors on June 2, and by 5 monitors on
5 June 3 (**Figure C-4**). The highest O₃ was measured by the Joe Neal and Walter Johnson
6 monitors, which reported 69 and 70 ppbv, respectively, on June 3. The NOAA monitors at the
7 NLVA and AP recorded 67 and 71 ppbv, respectively, on June 2, and 70 and 60 ppbv on June 3.
8

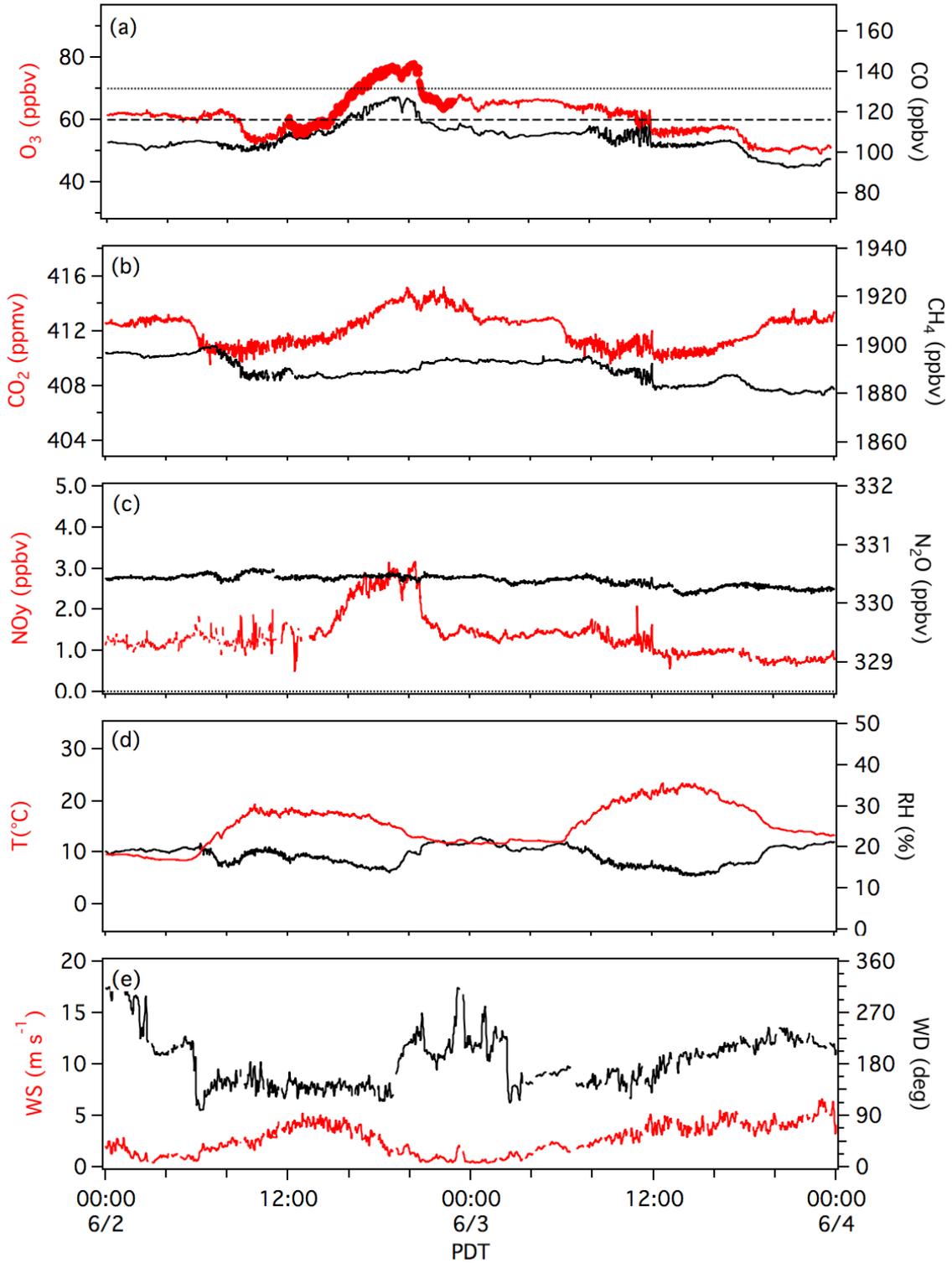
9 **Figure C-7a** shows that the high O₃ on June 2 was measured during the afternoon when
10 relatively moist air was transported to AP from the LVV by southeasterly upslope winds. The
11 high O₃ was positively correlated with CO and NO_y, but not with CH₄ or N₂O. This suggests that
12 pollution from the LVV was the primary source of the high O₃ at Angel Peak on June 2.
13

14 The MDA8 O₃ at AP was much lower (60 ppbv) on June 3, and the time series plots show that
15 the O₃ concentrations at AP dropped from a high of 64 at 0700 PDT to 53 ppbv at 1900 PDT as
16 the winds rotated from S to SSW. Ozone remained correlated with CO, however, suggesting
17 that these measurements reflected regional photochemistry.
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19



20 **Figure C-4.** Topographic maps of Clark County showing the June 2-3 MDA8 O₃ distributions.
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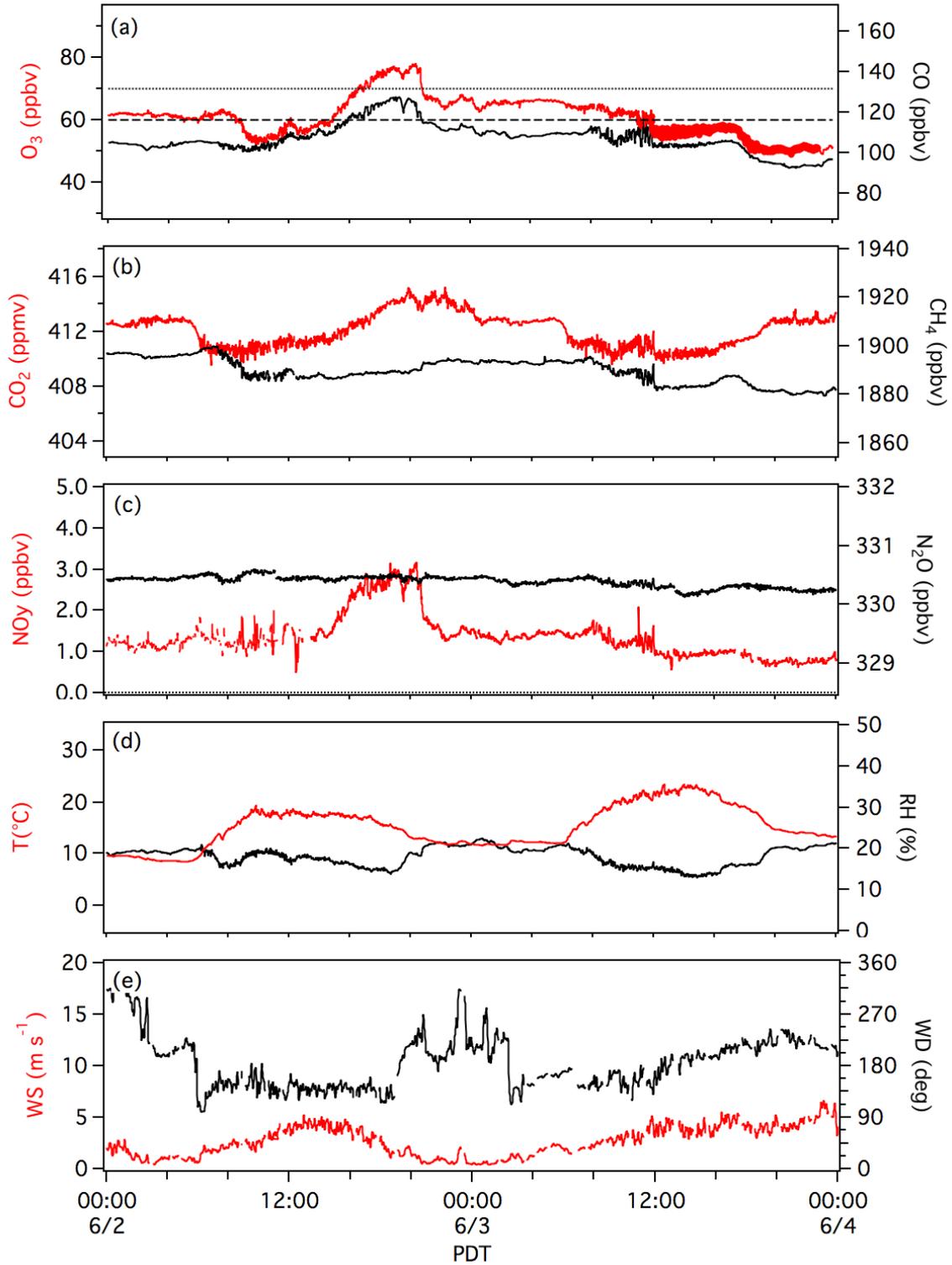
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Figure C-5. Time series of the Angel Peak mobile laboratory measurements on June 2.

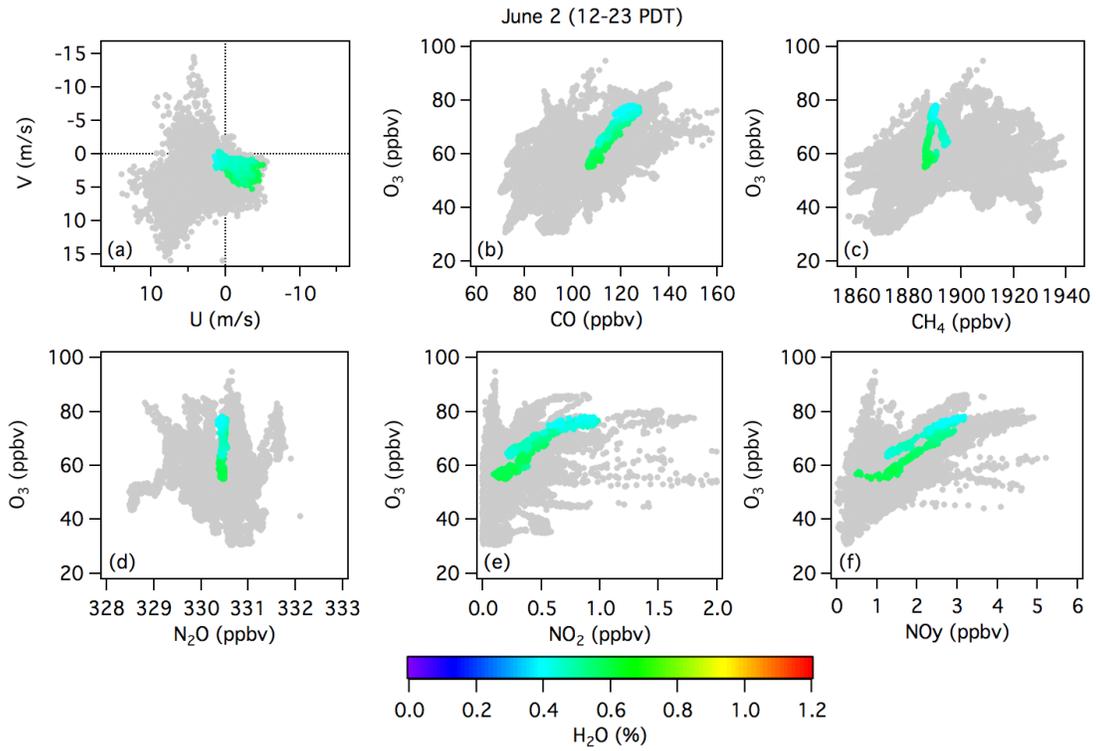
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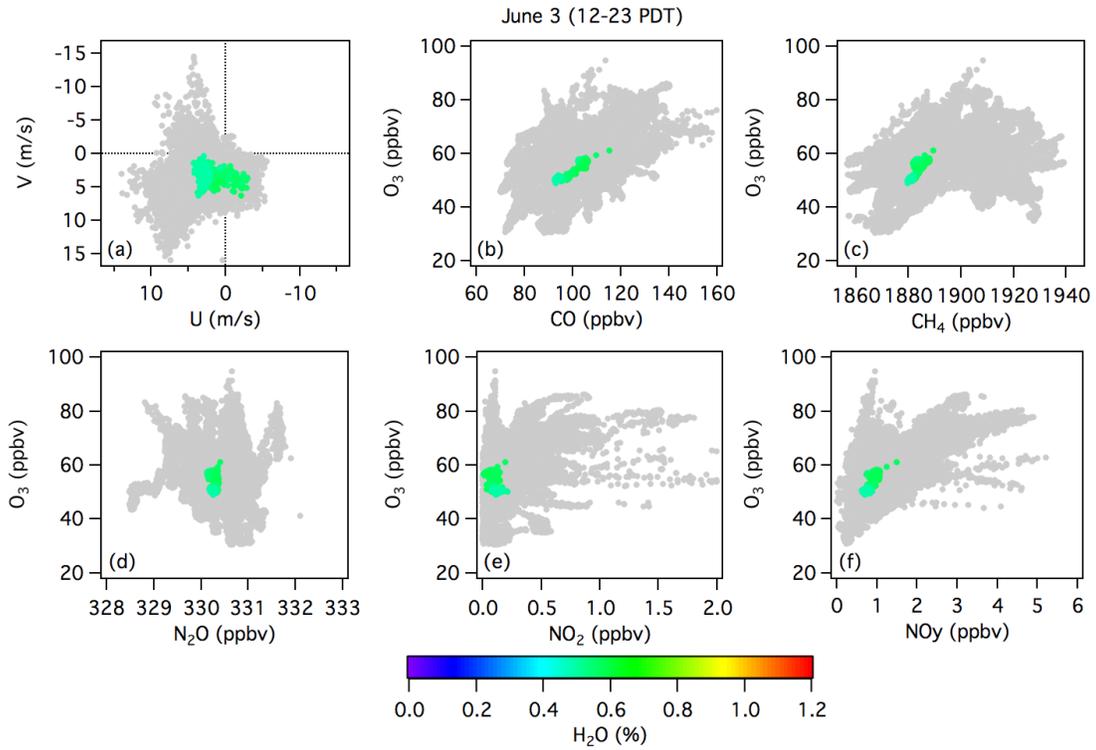
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Figure C-6. Time series of the Angel Peak mobile laboratory measurements on June 3.

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5 **Figure C-7.** Scatter plots showing the relationships between O_3 and the other parameters

6 measured by the mobile laboratory on June 2 and 3.

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2 **June 11-12**
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4 The cold front associated with the deep upper level low that passed through the SWUS on June
5 10-12 brought cool temperatures to Clark County and the temperature on Angel Peak fell below
6 freezing on the night of June 11-12. Ozone also decreased across the LVV as O₃-poor mid-
7 tropospheric air was advected inland from the Pacific Ocean. This wholesale replacement of the
8 continental airmass more than offset the impact of the O₃-rich air transported downward from
9 the lower stratosphere by the tropopause fold. This stratospheric intrusion descended to the
10 summit of Angel Peak on June 11, but **Figure C-8** shows that O₃ remained moderate in Clark
11 County. The CRDS instrument on Angel Peak recorded an MDA8 O₃ concentration of 66 ppbv,
12 but the highest regulatory MDA8 O₃ values recorded in the LVV were 60 ppbv at Apex and 63
13 ppbv at Indian Springs. The concentrations at all of the Clark County monitors fell to much
14 lower levels (40-50 ppbv) on June 12 as the intrusion moved south and east into Arizona and
15 was followed by the clean Pacific air.

16
17 **Figures C-9** through 11 show that the 1-min O₃ peaks of 74 ppbv at 1221 PDT and 82 ppbv at
18 2300 PDT followed the strong southwesterly winds of the approaching trough and were
19 accompanied by very dry air with low NO_y. This O₃ was also anticorrelated with CO, CH₄, and
20 N₂O; hallmarks of stratospheric intrusions. The two O₃ peaks were separated by moister mid-
21 tropospheric Pacific air that was also anticorrelated with the surface air markers. Clean middle
22 and upper tropospheric Pacific air followed behind the intrusion on June 11-12 as the synoptic
23 flow rotated to the northwest behind the receding trough.

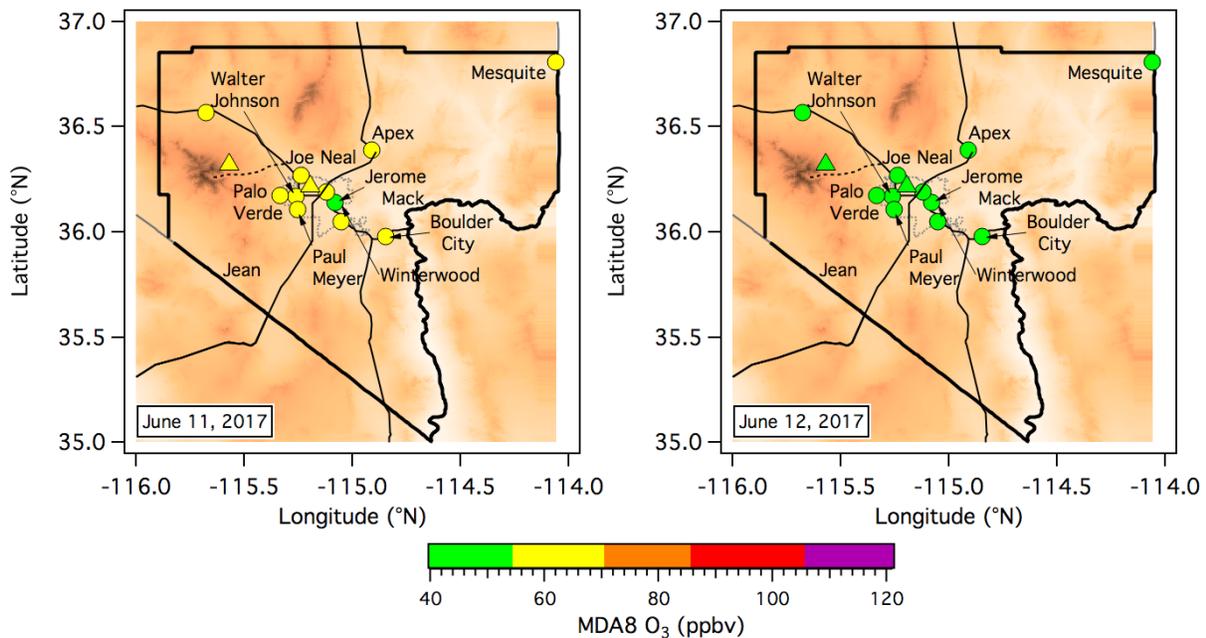
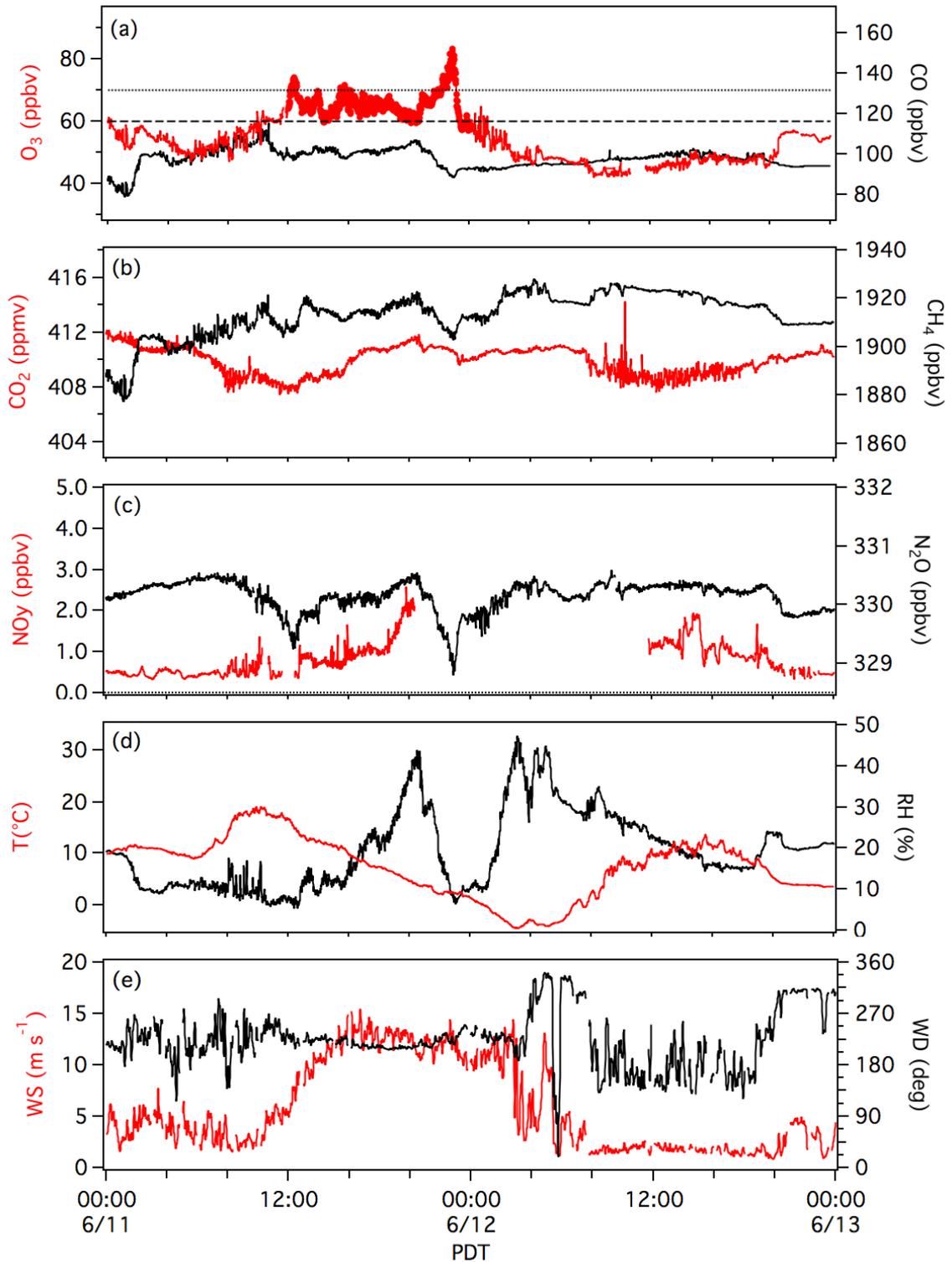


Figure C-8. Topographic maps of Clark County showing the June 11-12 MDA8 O₃ distributions.

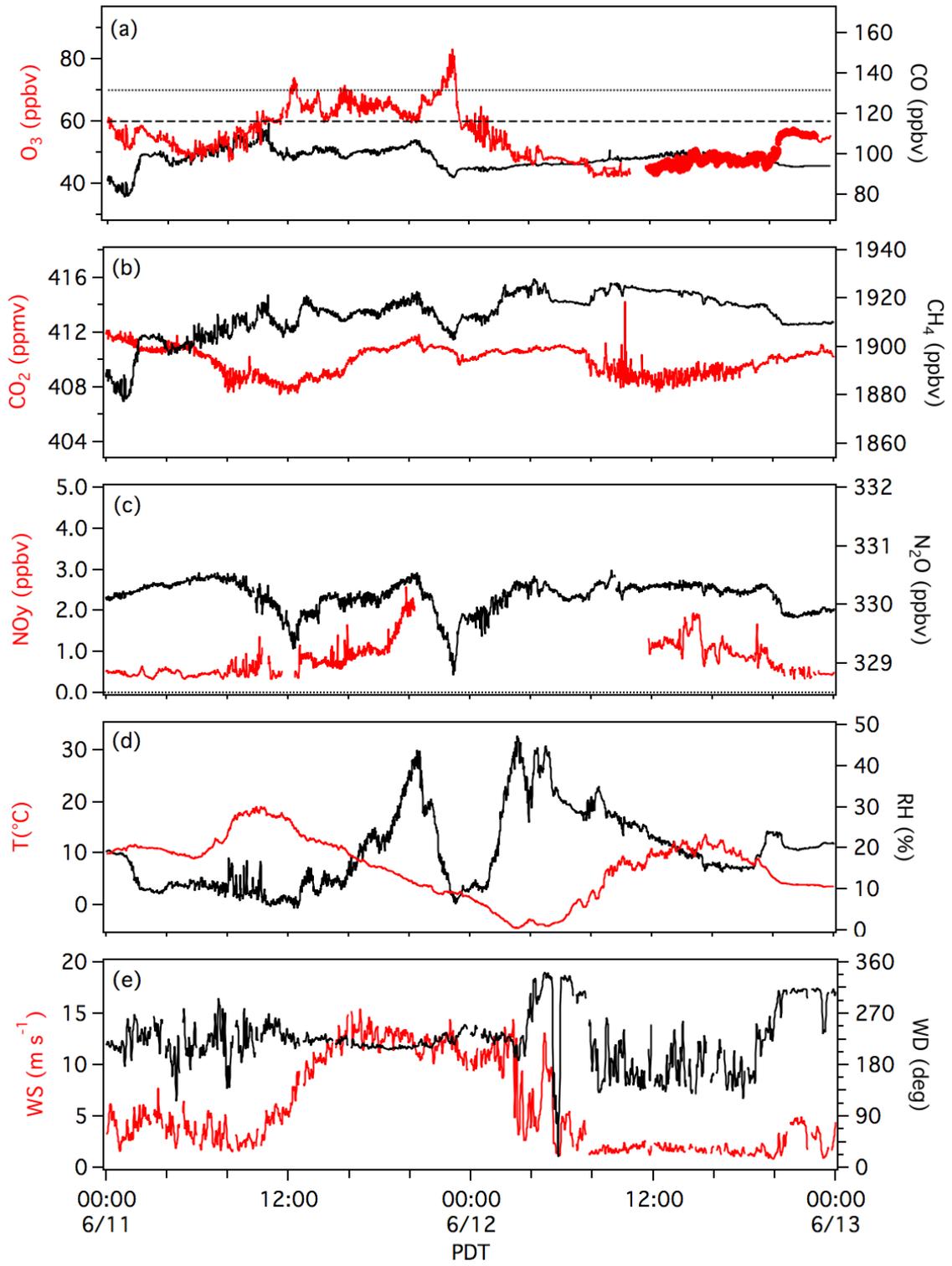
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Figure C-9. Time series of the Angel Peak mobile laboratory measurements on June 11.

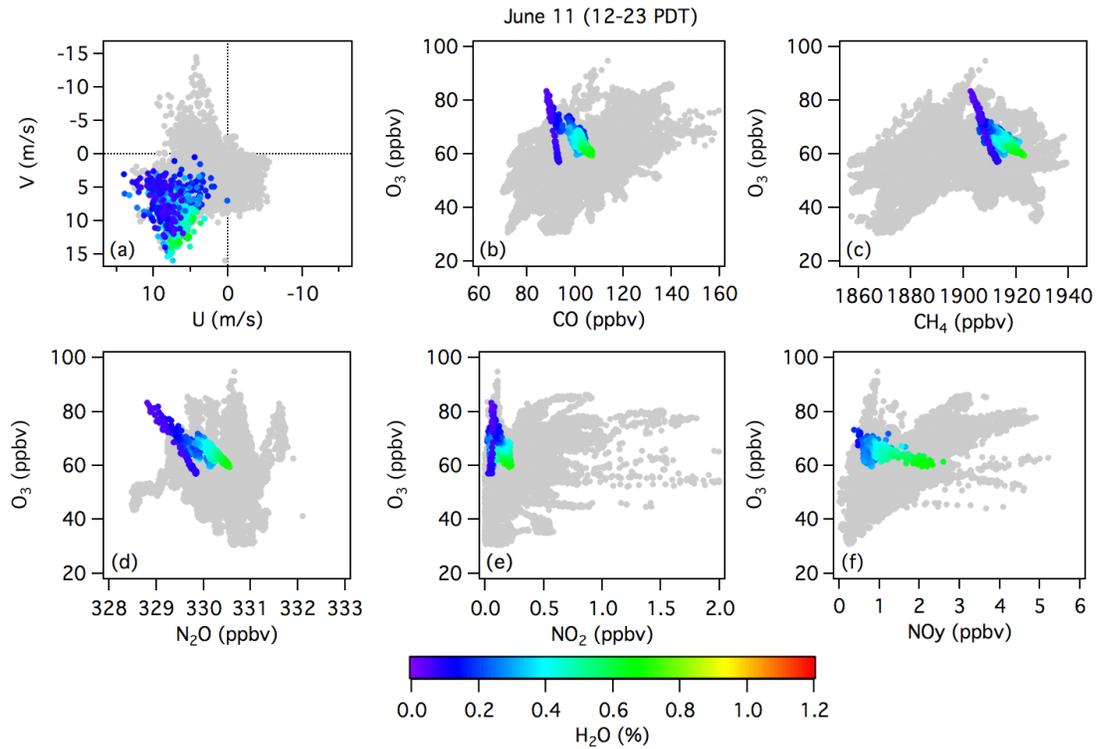
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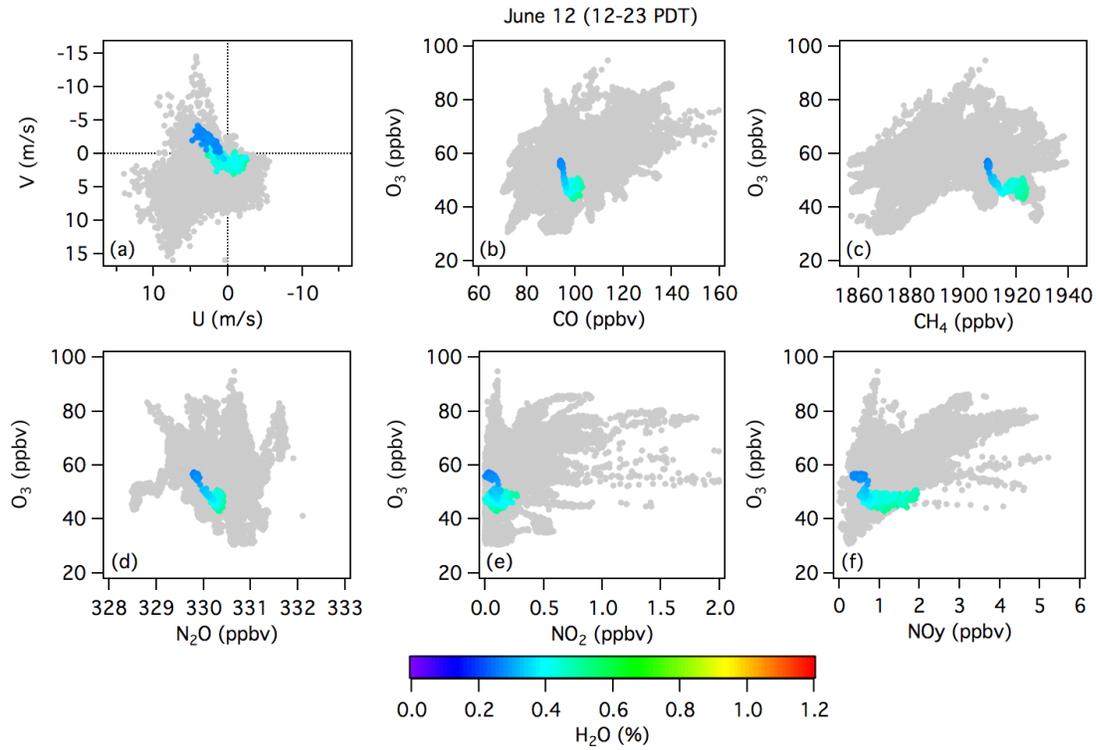
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Figure C-10. Time series of the Angel Peak mobile laboratory measurements on June 12.

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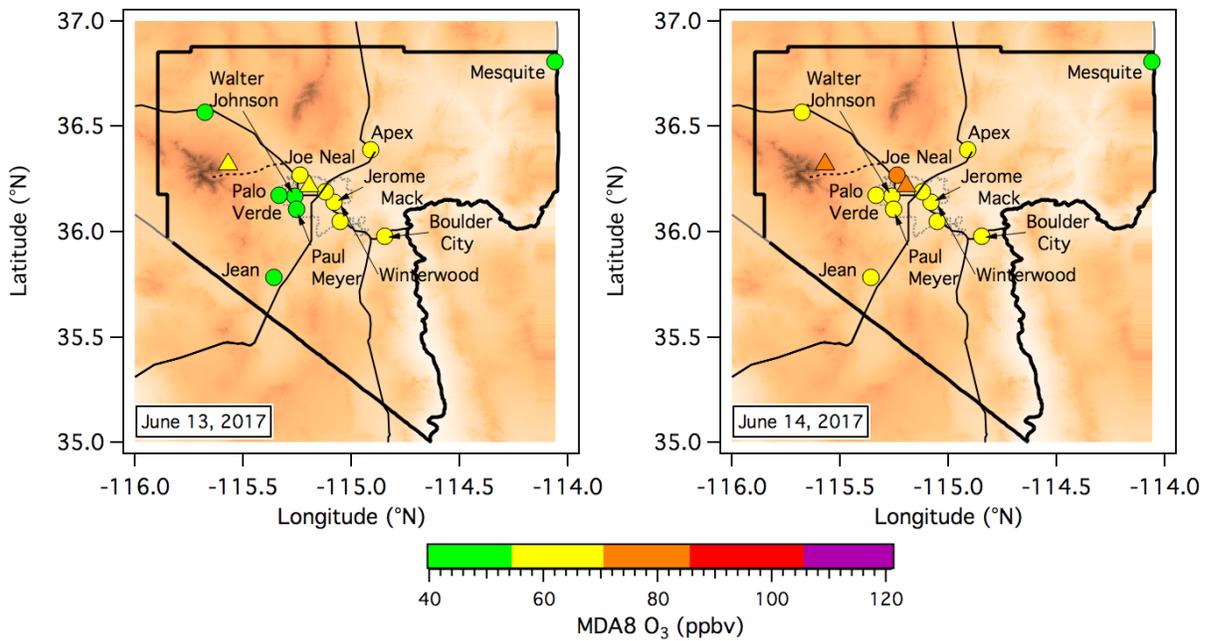
4

5 **Figure C-11.** Scatter plots showing the relationships between O_3 and the other parameters

6 measured by the mobile laboratory on June 11 and 12.

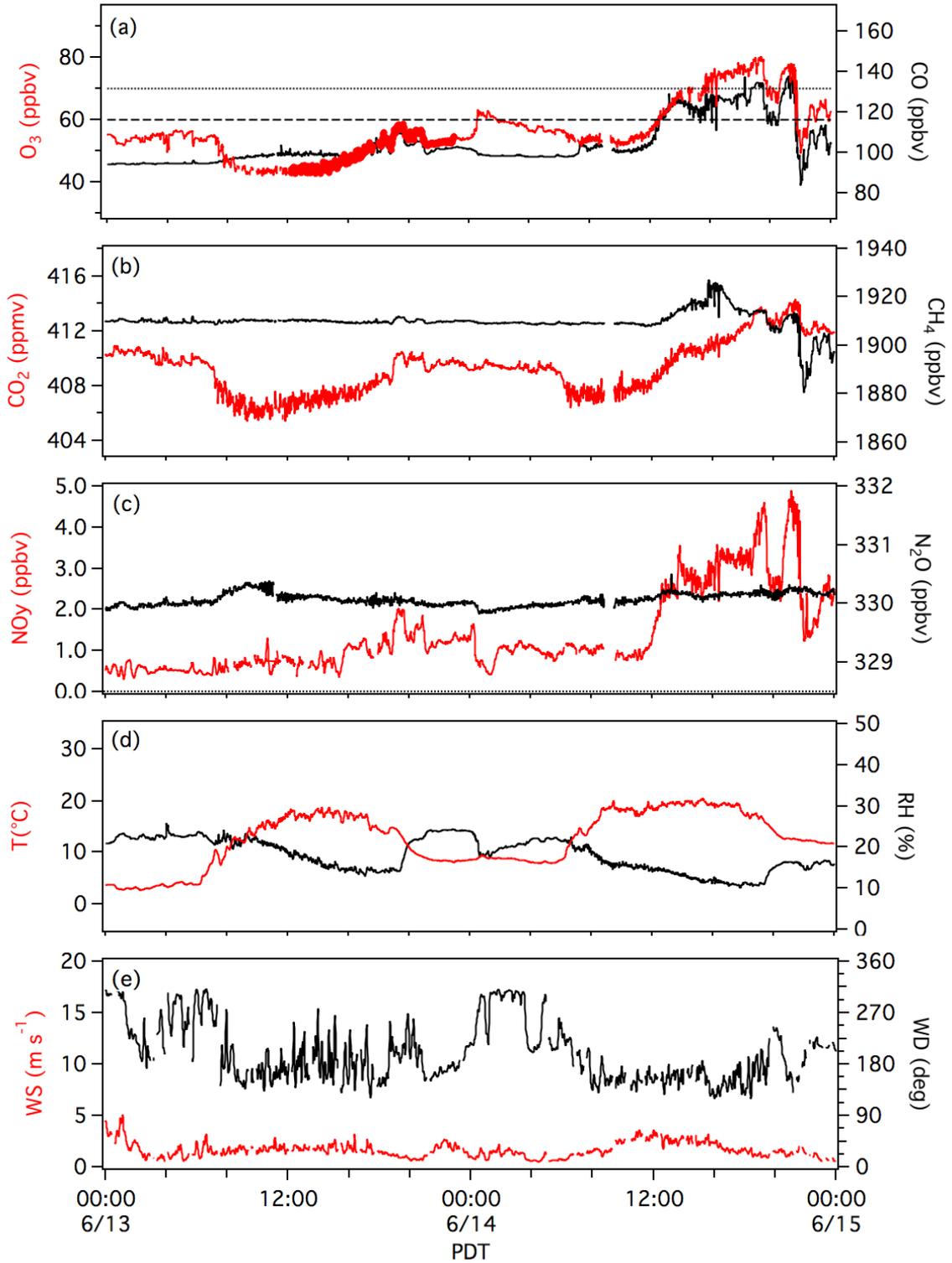
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2 **June 13-14**
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4 Surface O₃ slowly increased across the LVV on June 13 as warmer temperatures followed the
5 passage of the upper level low and cold front and promoted local photochemical production.
6 **Figure C-12** shows that MDA8 O₃ remained good-to-moderate on June 13, but increased
7 significantly on June 14 when the NAAQS was exceeded by the Joe Neal monitor with 73 ppbv,
8 and equaled by the Walter Johnson monitor with 70 ppbv. The NLVA monitor recorded an
9 MDA8 of 74 ppbv and the CRDS instrument on AP 71 ppbv. The correlations between O₃ and
10 CO, CH₄, and NO_y became stronger as moister polluted air was transported from the valley to
11 AP by the afternoon upslope. The AP measurements show no evidence of undiluted lower
12 stratospheric air.
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14
15 **Figure C-12.** Topographic maps of Clark County showing the June 13-14 MDA8 O₃ distributions.
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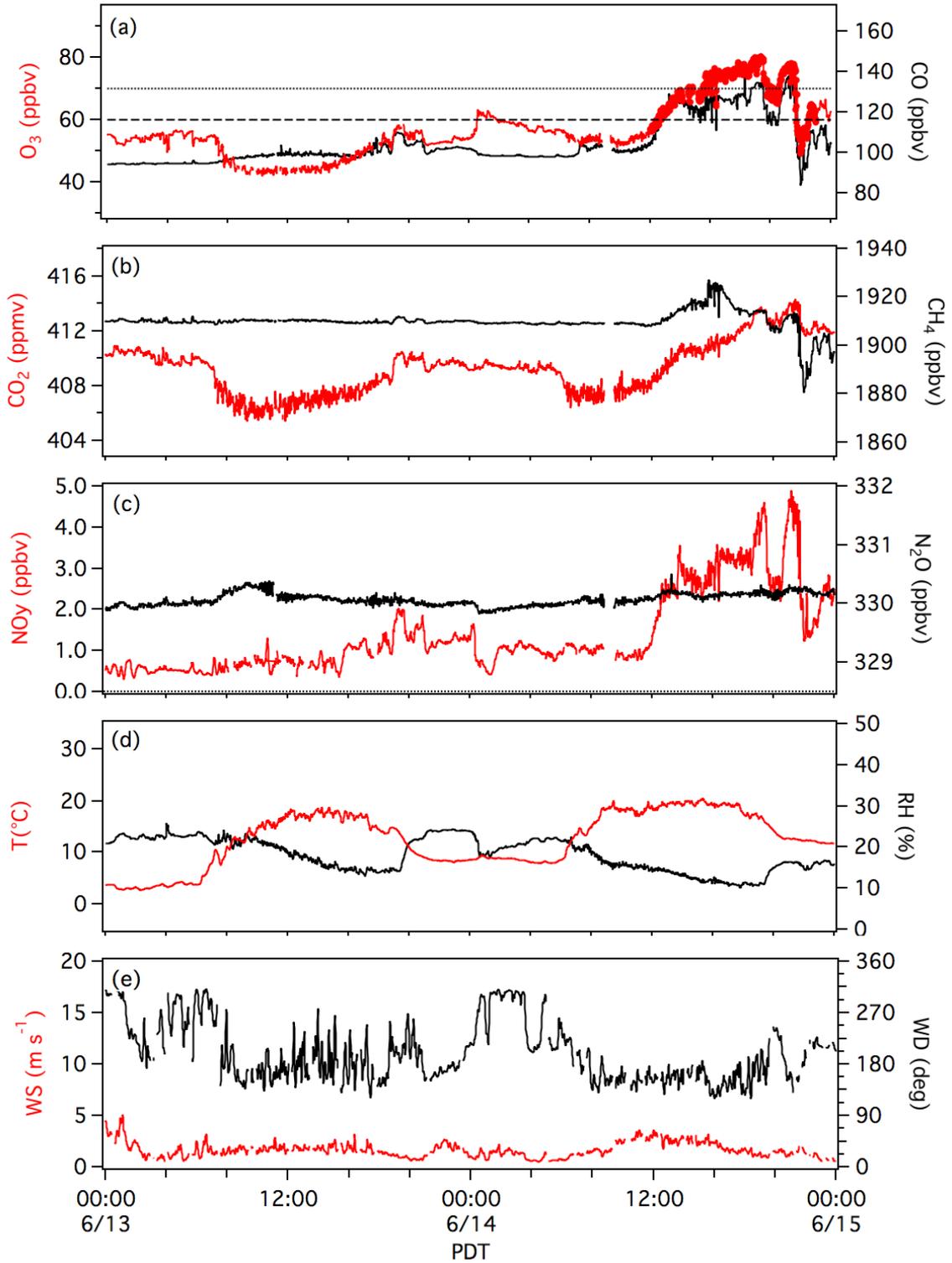
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Figure C-13. Time series of the Angel Peak mobile laboratory measurements on June 13.

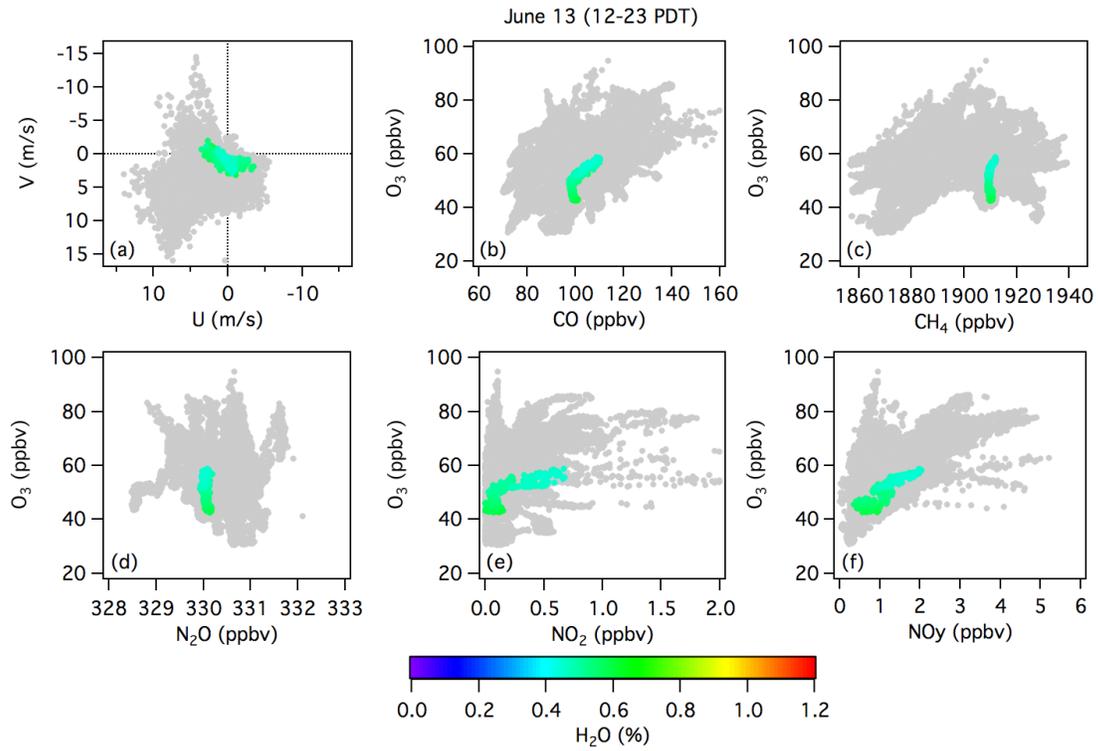
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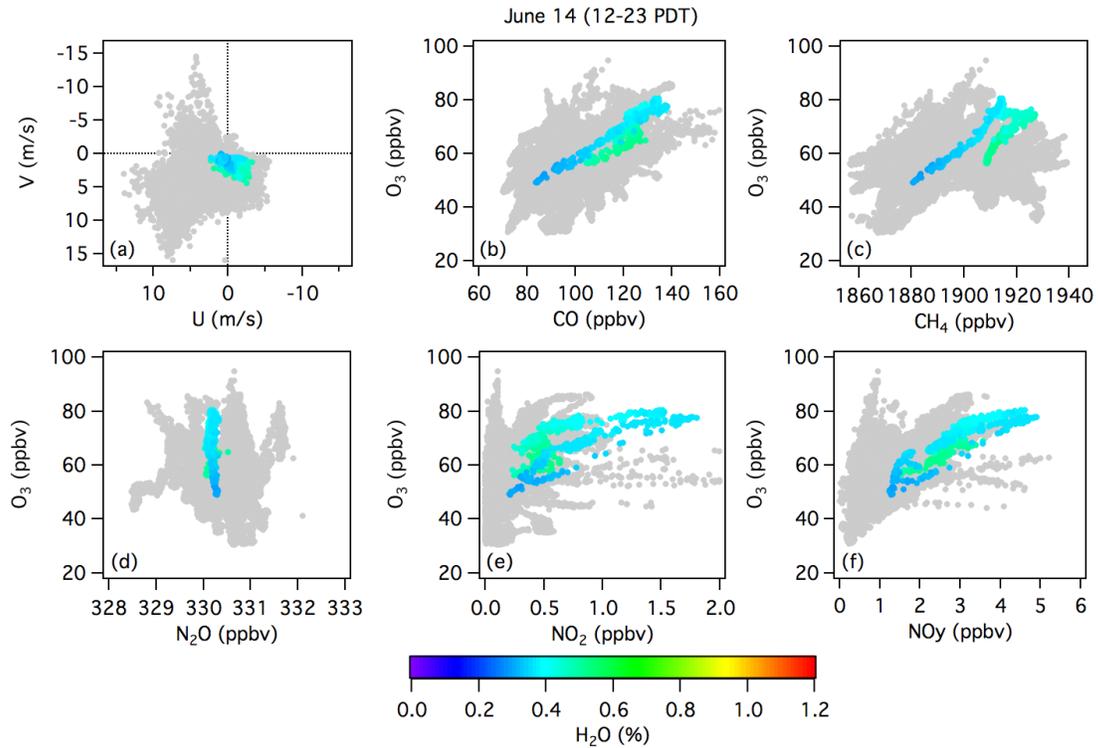
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Figure C-14. Time series of the Angel Peak mobile laboratory measurements on June 14.

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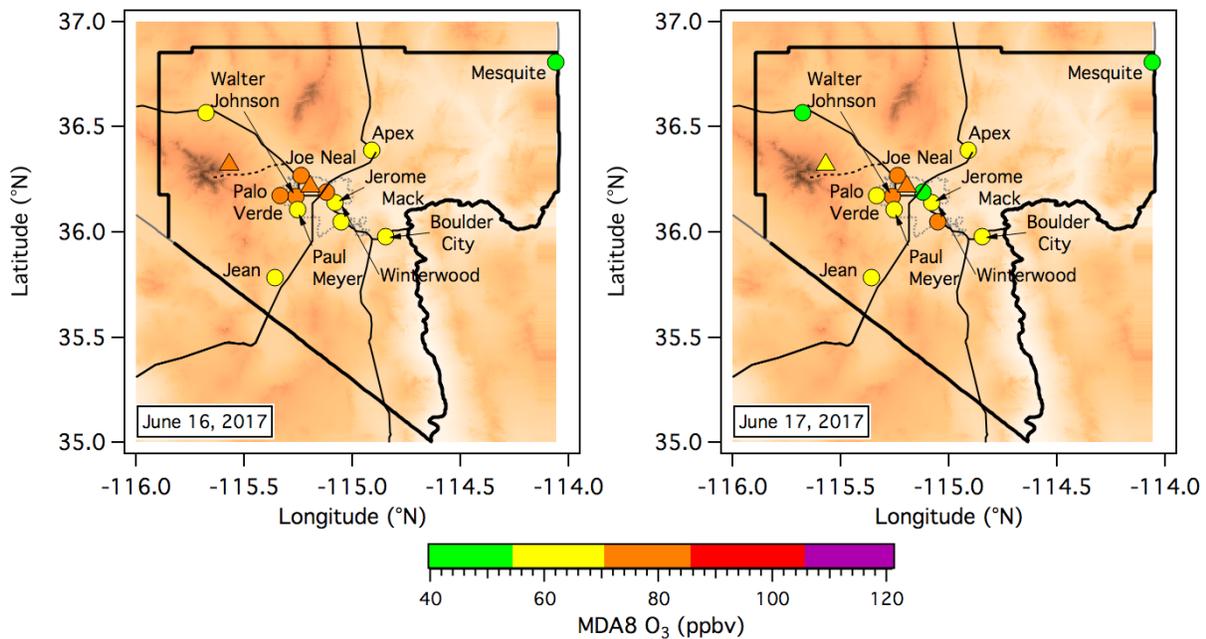
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4 **Figure C-15.** Scatter plots showing the relationships between O_3 and the other parameters
5 measured by the mobile laboratory on June 13 and 14.

6

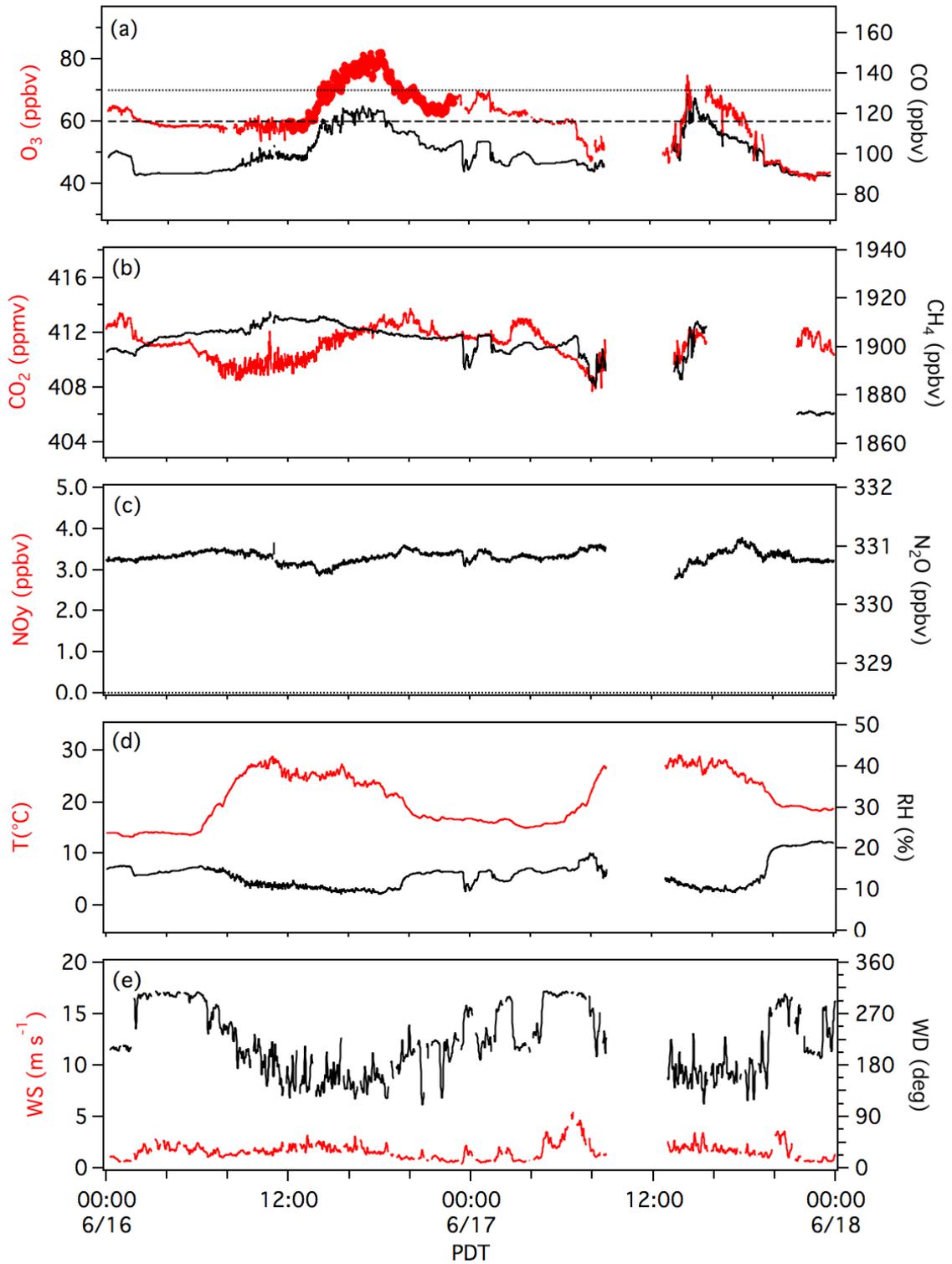
1
2 **June 16-17**
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4 High ozone was present all across the LVV on both June 16-17 with MDA8 concentrations of 63
5 ppbv or more at all of the CCDAQ monitors except Mesquite, and exceedances at 3 of the
6 monitors on the 16th and 4 on the 17th. The Joe Neal monitor measured the highest MDA8 O₃
7 with 74 ppbv on June 16 and 72 ppbv on June 17. The NLVA monitor measured 76 and 73 ppbv,
8 respectively, on these two days, and the CRDS instrument on Angel Peak measured 74 and 61
9 ppbv. The AP time series and scatter plots show that O₃ was positively correlated with CO on
10 both days, and the highest concentrations occurred during periods of upslope flow from the
11 valley. Ozone was uncorrelated with N₂O on both days, suggesting that there were no fire
12 influences. Ozone was also uncorrelated with CH₄ on the 16th, but positively correlated on the
13 17th. Unfortunately, there were no NO_y measurements at AP during this period, but the
14 available measurements suggest that the high O₃ in both the LVV and AP was associated with
15 regional pollution with no fire, long-range transport, or stratospheric influences.
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21 **Figure C-16.** Topographic maps of Clark County showing the June 13-14 MDA8 O₃ distributions.
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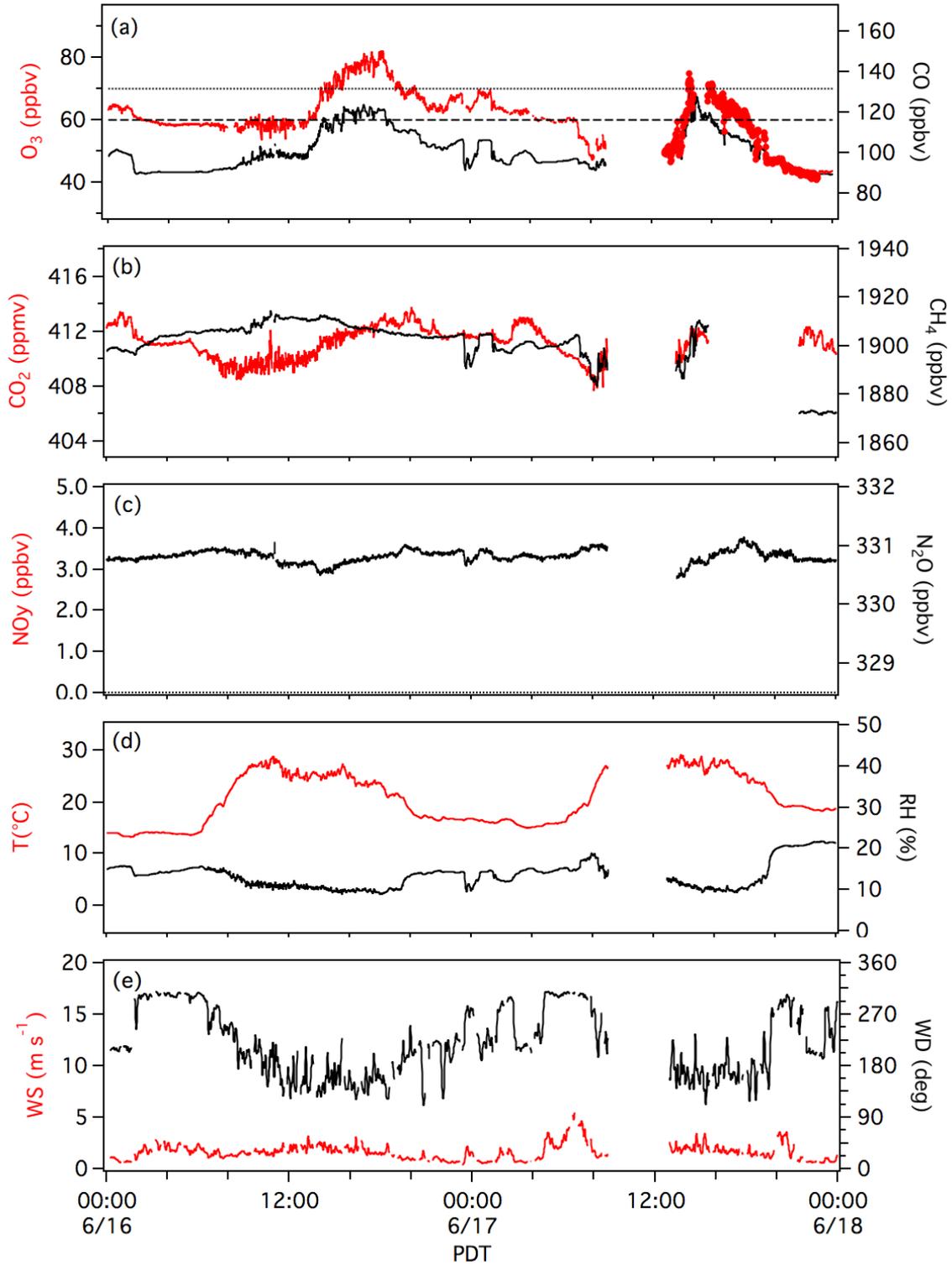
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Figure C-17. Time series of the Angel Peak mobile laboratory measurements on June 16.

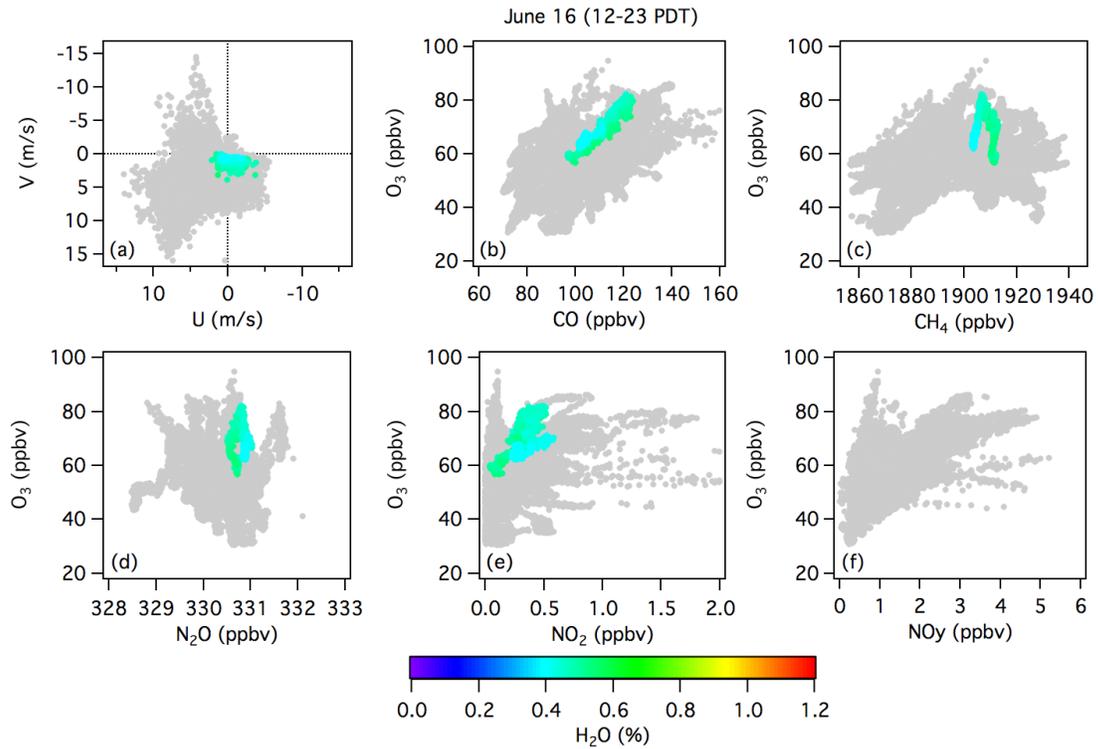
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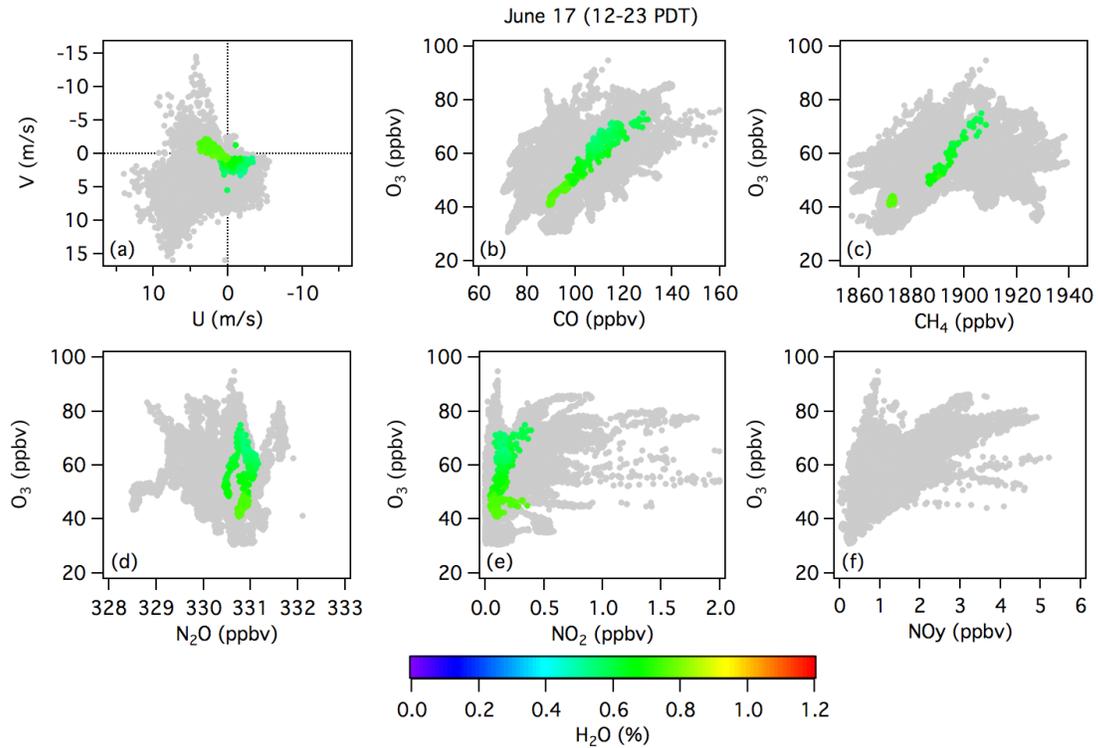
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Figure C-18. Time series of the Angel Peak mobile laboratory measurements on June 17.

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5 **Figure C-19.** Scatter plots showing the relationships between O_3 and the other parameters

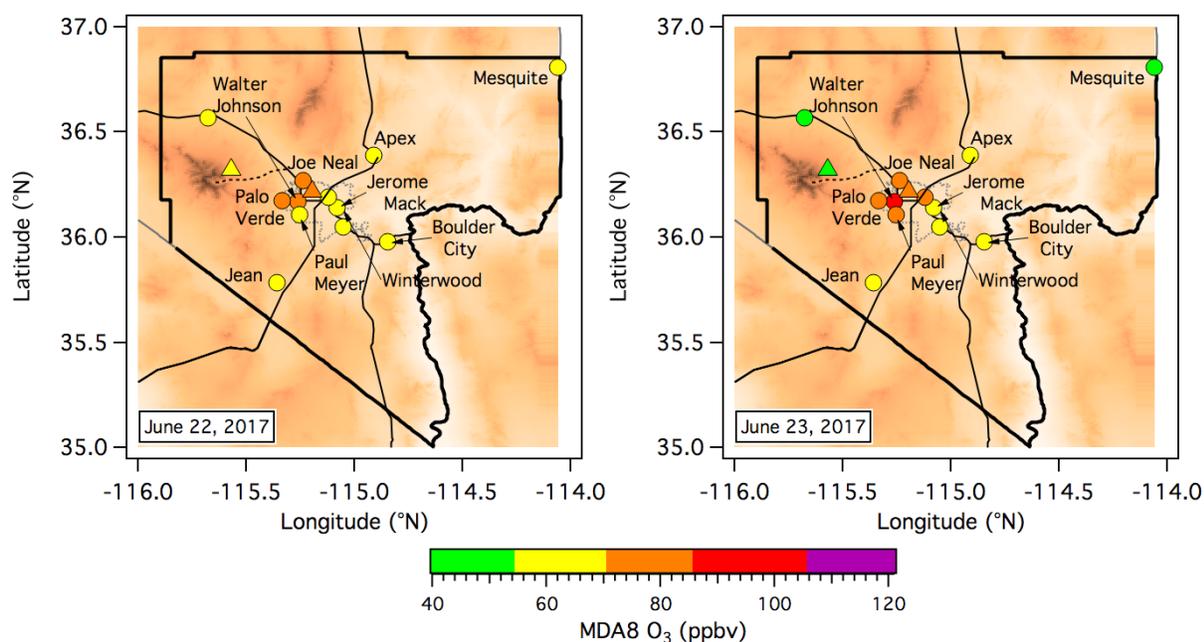
6 measured by the mobile laboratory on June 16 and 17.

1
2 **June 22-23**
3

4 The hottest temperatures of the FAST-LVOS campaign were measured on June 22 when the all-
5 time Las Vegas record of 117°F was tied at McCarran International Airport. The highest O₃ of
6 the campaign was also measured on June 22-23, with peak MDA8 concentrations of 77 ppbv at
7 Joe Neal on both June 22 and 23. The Walter Johnson monitor reported 87 ppbv on June 23,
8 but this number was based on only 6 hours of measurements and the actual value should be
9 somewhat lower. The NLVA monitor measured 76 and 83 ppbv on June 22 and 23, respectively,
10 but the CRDS instrument on Angel Peak measured only 63 and 58 ppbv.

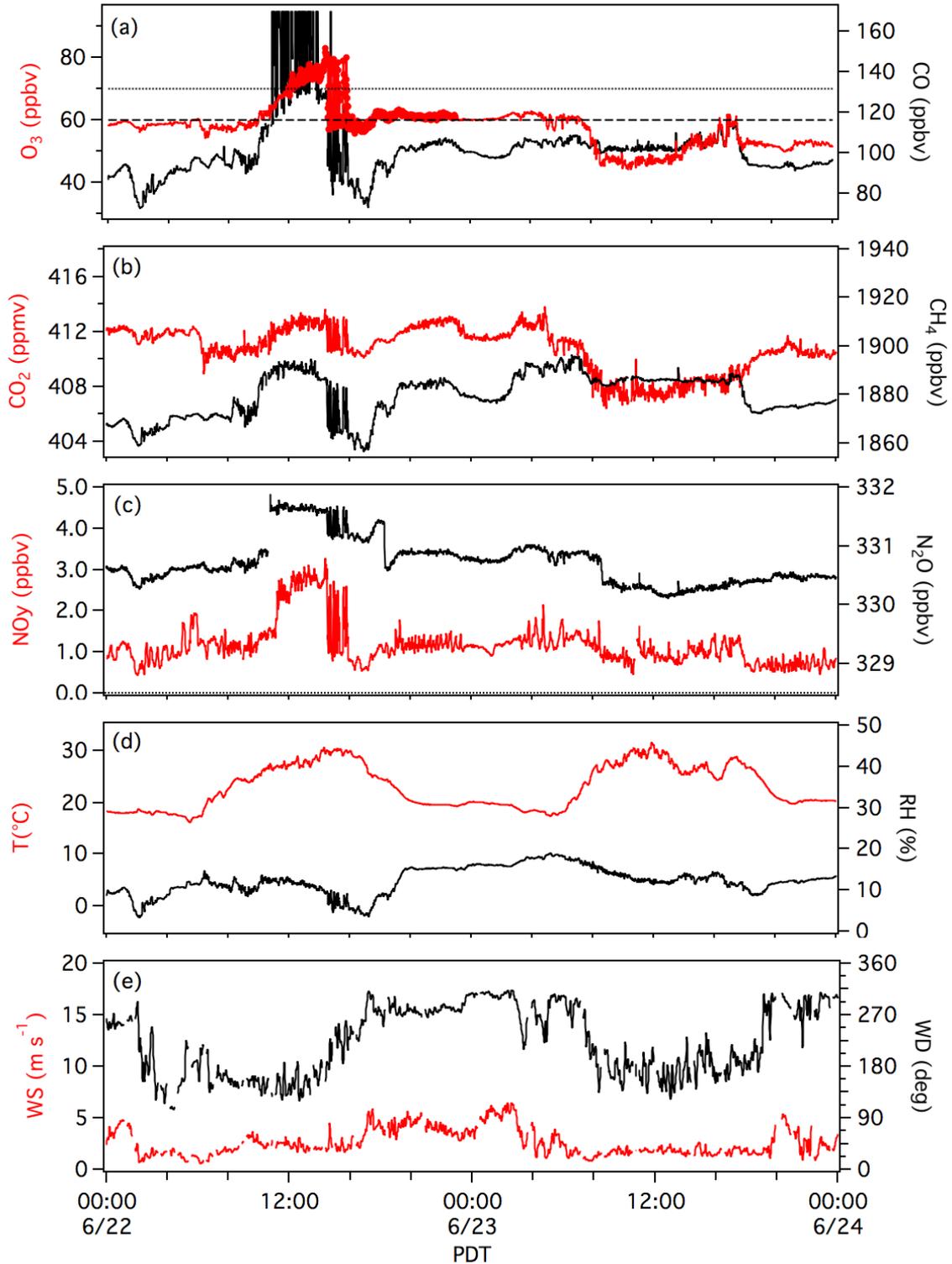
11
12 The instruments in the mobile laboratory sampled a distinctive plume with very high CO (≈450
13 ppbv) and elevated O₃ (≈80 ppbv) and around midday on June 22 (**Figure C-21**). This plume,
14 which has been attributed to the Holcomb Fire near Big Bear, CA, was also enhanced in CO₂,
15 CH₄, NO_y, and N₂O, and disappeared when the winds shifted from southerly to northwesterly in
16 the afternoon, bringing much drier clean air to the summit (blue points in **Figure C-23**, top).

17
18 Ozone remained relatively low at AP on June 23, with the highest concentrations of ≈60 ppbv
19 appearing around 1700 PDT in the moister air transported to the summit by the afternoon
20 upslope flow. Ozone decreased again when the winds shifted from southeasterly to
21 northwesterly in the early evening.



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25 **Figure C-20.** Topographic maps of Clark County showing the June 13-14 MDA8 O₃ distributions.
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Figure C-21. Time series of the Angel Peak mobile laboratory measurements on June 22.

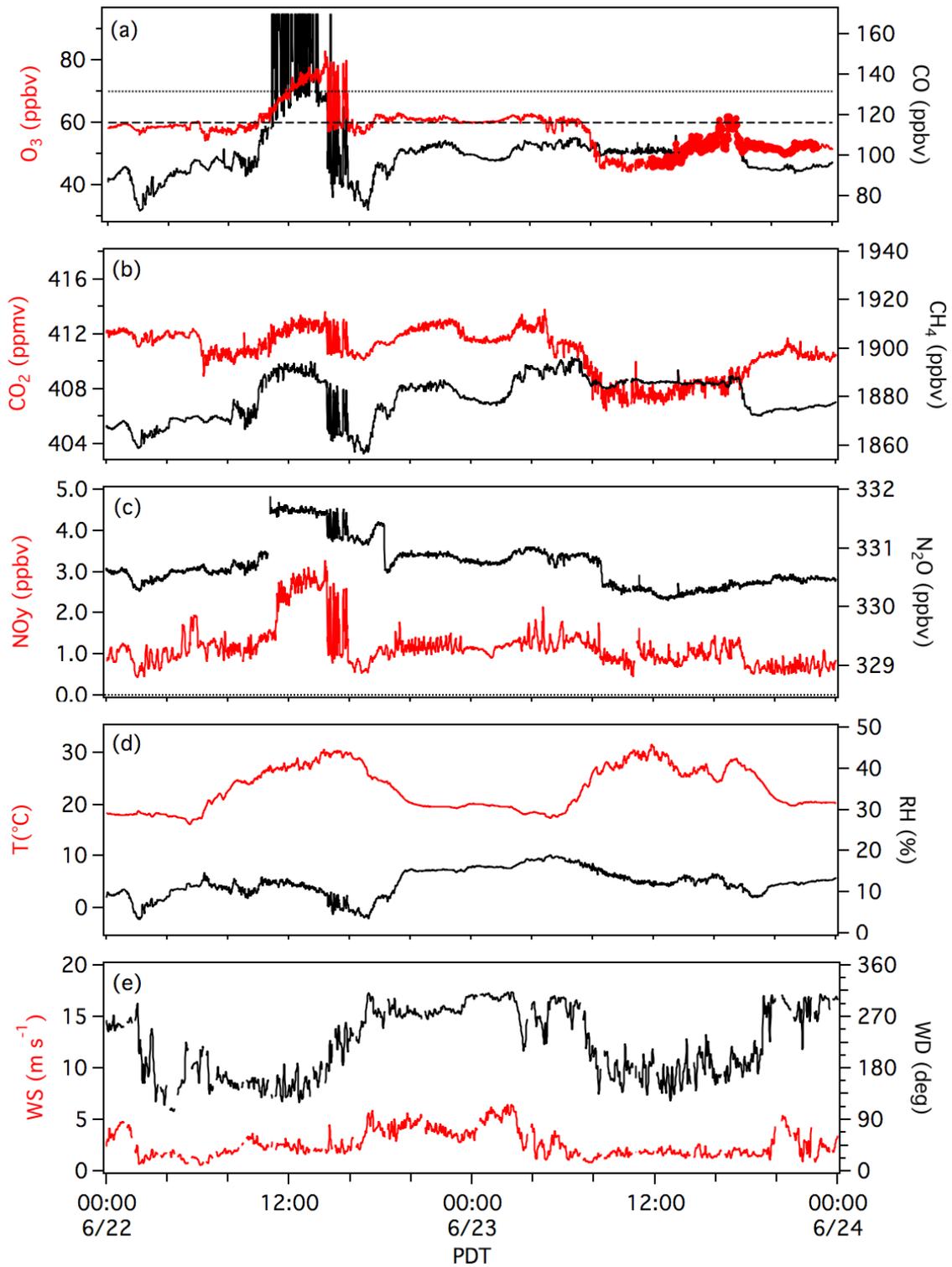
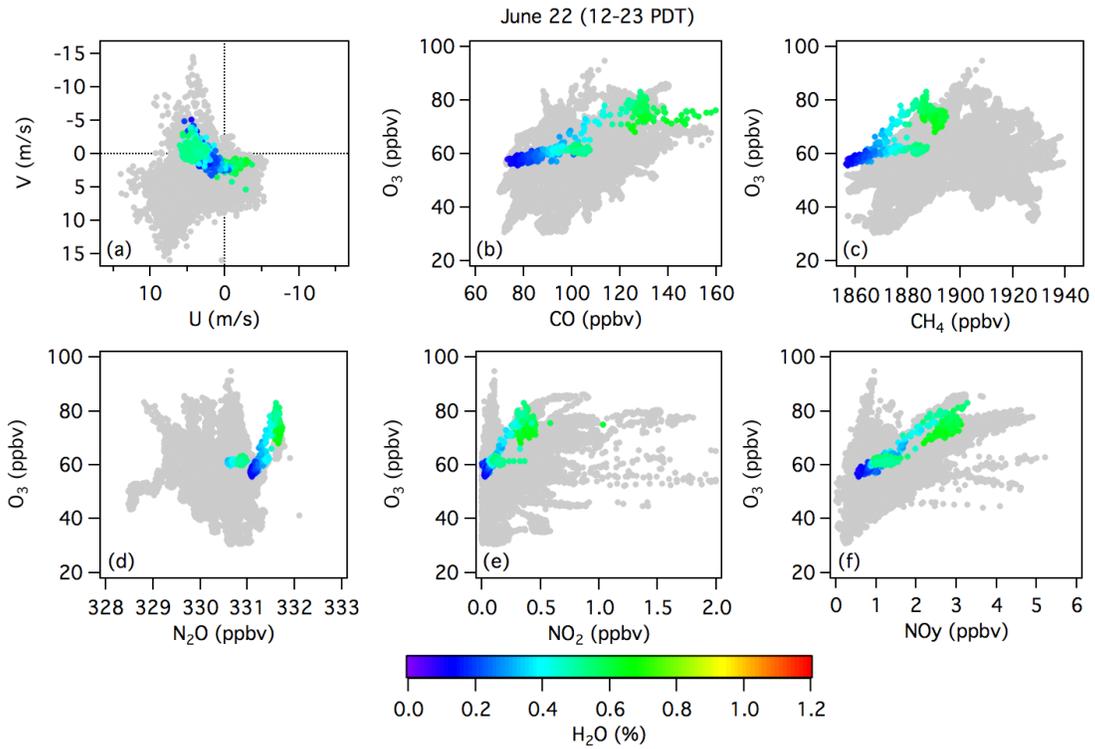
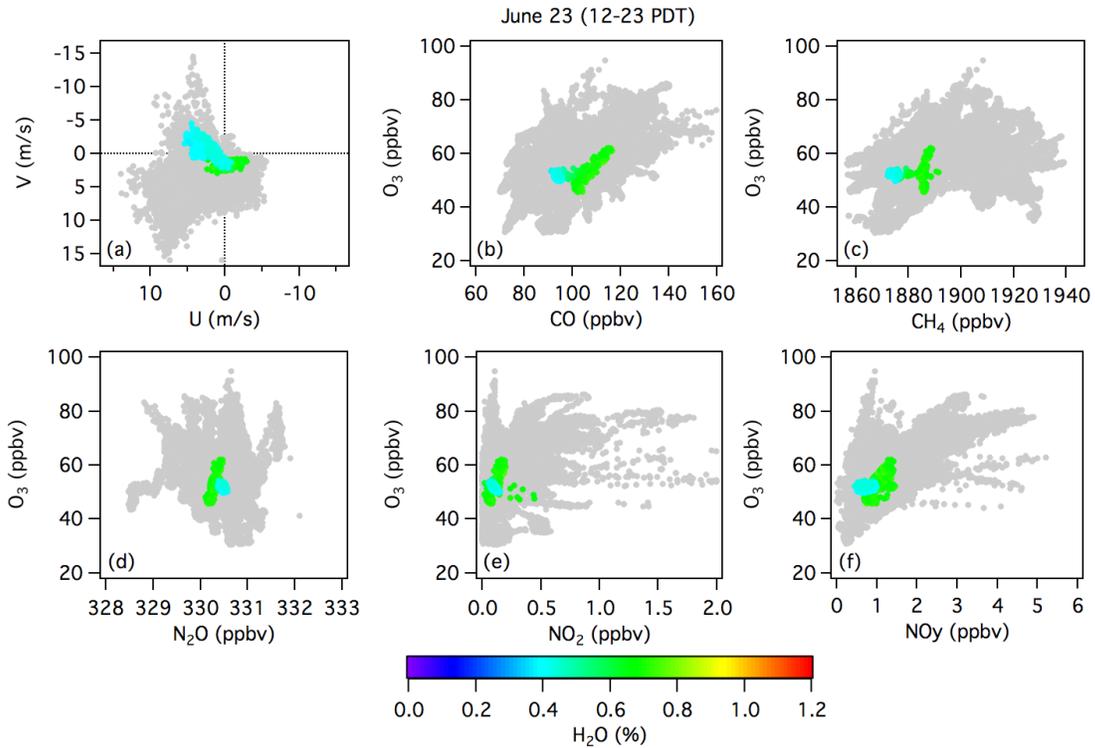
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Figure C-22. Time series of the Angel Peak mobile laboratory measurements on June 23.

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4 **Figure C-23.** Scatter plots showing the relationships between O_3 and the other parameters
5 measured by the mobile laboratory on June 22 and 23.

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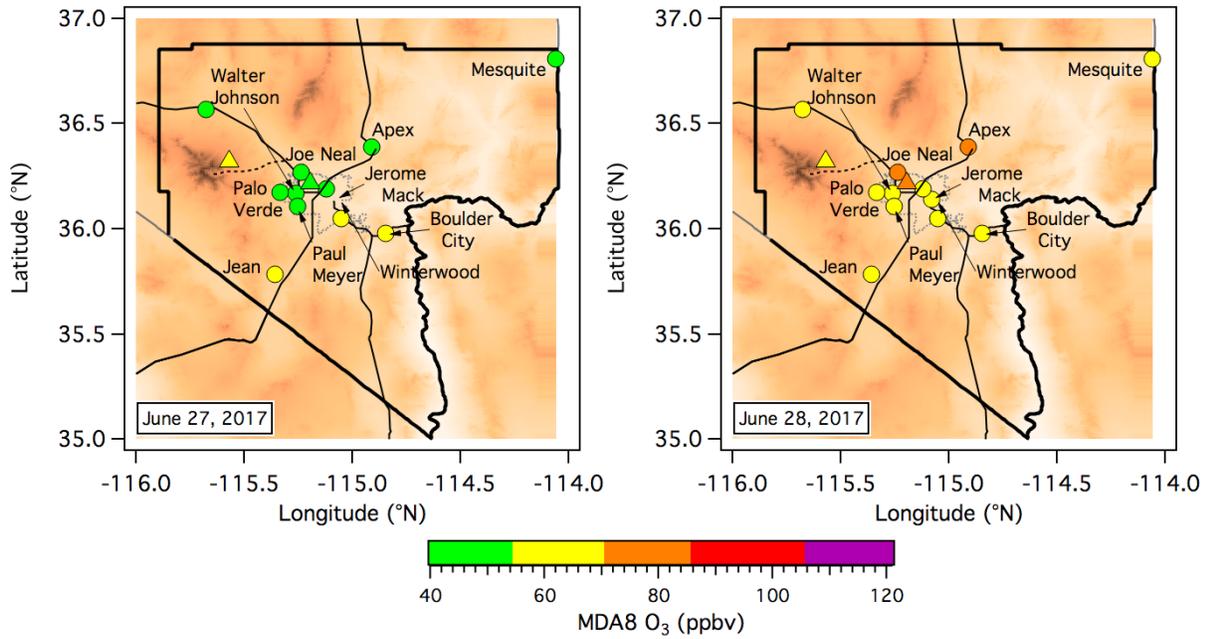
1
2 **June 28-30**
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4 The last three days of FAST-LVOS were also very hot with high O₃ measured throughout Clark
5 County (**Figure C-24**). Ozone was good to moderate on June 27, but MDA8 mixing ratios of 59
6 ppbv or more were recorded by all of the CCDAQ monitors, including the outlying Mesquite and
7 Jean CAMS, on June 28-30 days. The NAAQS was exceeded at Joe Neal on all 3 days, with MDA8
8 concentrations of 74,70, and 75 ppbv, respectively. The NOAA measurements at the NLVA were
9 very similar with 74,71, and 76 ppbv. The NOAA CRDS instrument on AP also measured high O₃
10 with 68, 71, and 74 ppbv, respectively.

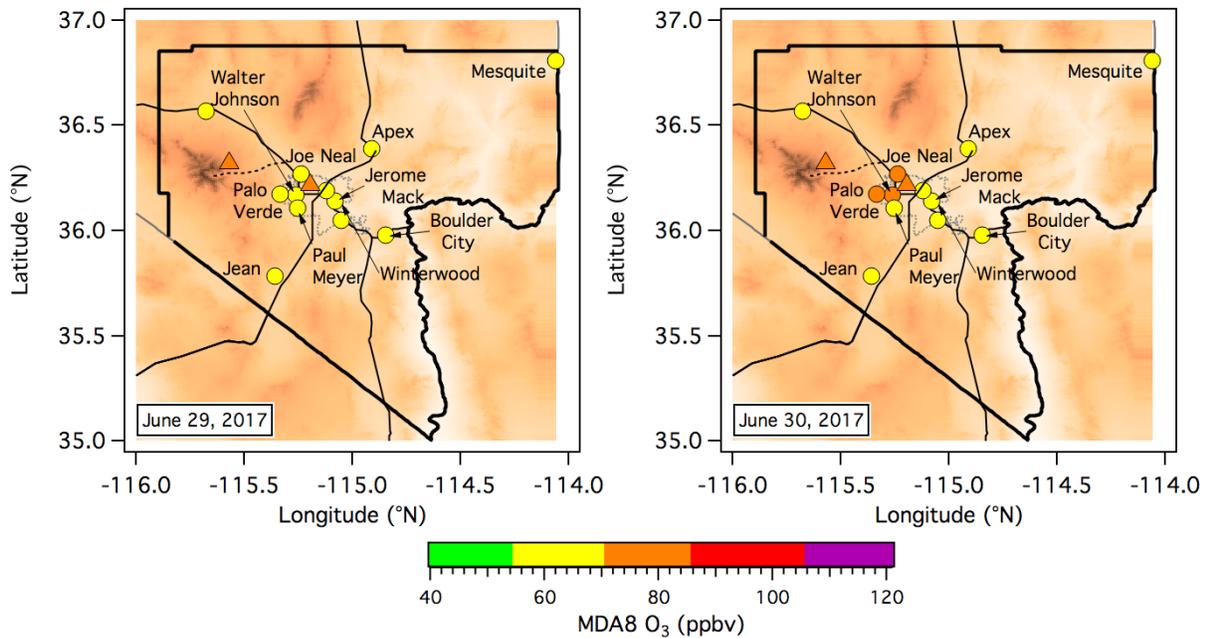
11
12 The TOPAZ measurements from June 28 (**Figure 9-42**) showed the entrainment of a suspected
13 fire plume by the convective mixed layer above the LVV near midday, but this plume remained
14 well above the summit of Angel peak and was not detected in the mobile laboratory in-situ
15 measurements. The highest O₃ appeared several hours later in a short-lived (50 min) spike of
16 dry air advected over the summit when the winds shifted from SSE to NW around 1830 PDT
17 (**Figure C-25**). The scatter plots show that the O₃ in this spike was positively correlated with
18 both CO and CH₄, but with a very steep slope. The low H₂O content suggests that the spike was
19 caused by a filament of Asian pollution that had descended from the upper troposphere.

20
21 These measurements stand in contrast to the high O₃ concentrations detected on June 29 and
22 30 that were positively correlated with both CO and NO_y, and uncorrelated with CH₄. This O₃
23 was embedded in moister air brought to AP by the afternoon upslope flow and is attributed to
24 regional pollution.
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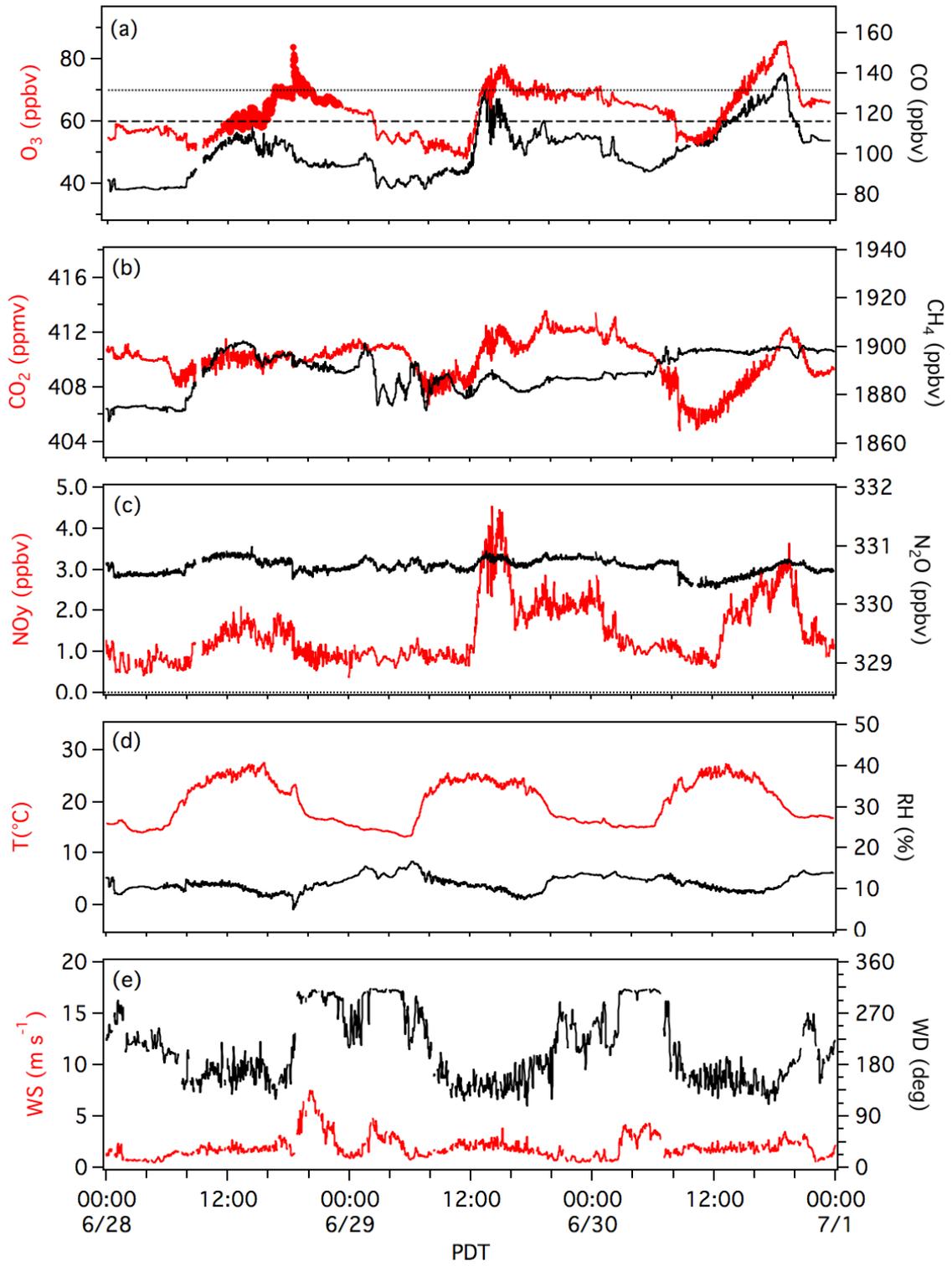
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Figure C-24. Topographic maps of Clark County showing the June 27-30 MDA8 O₃ distributions.

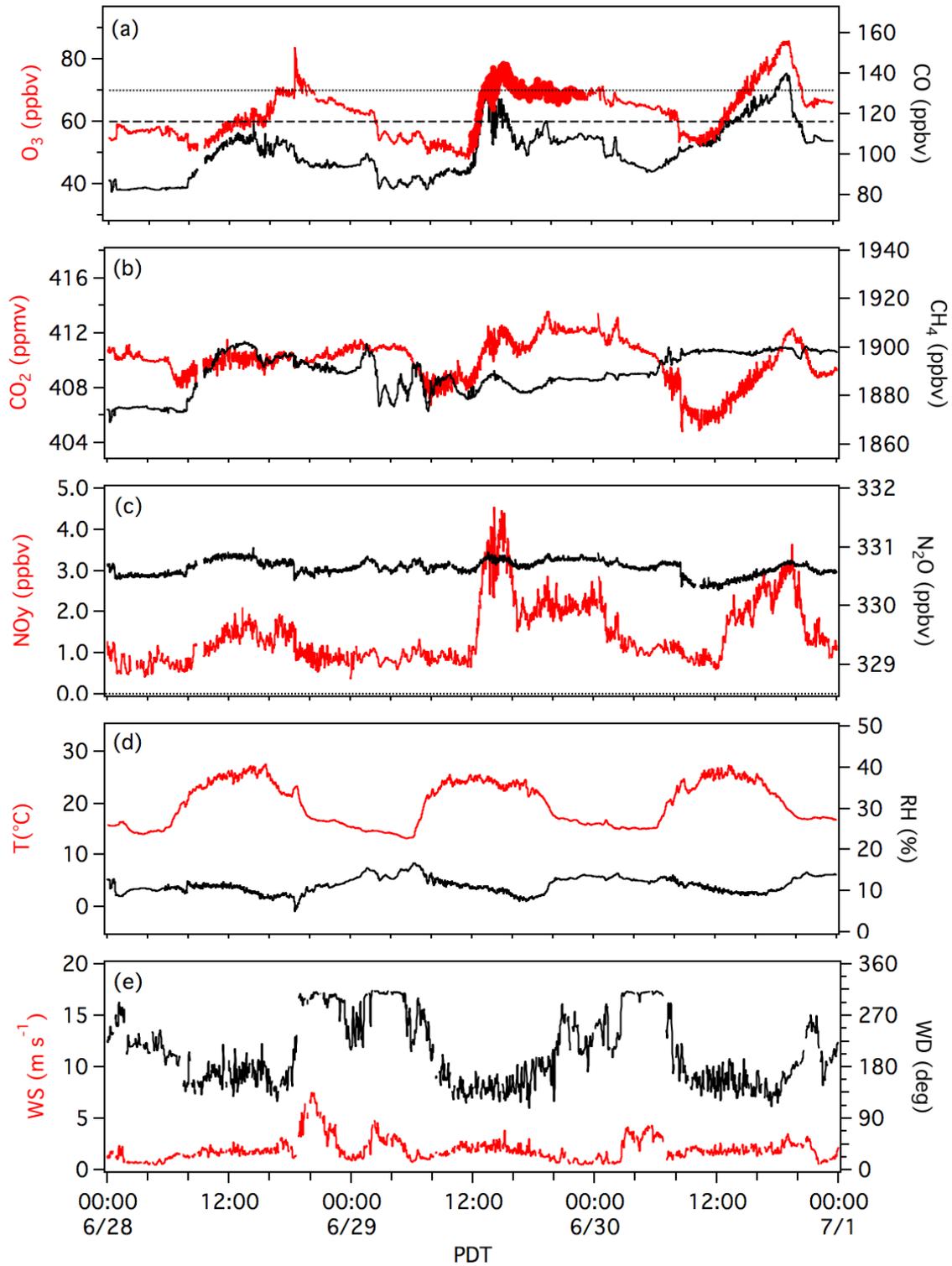
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Figure C-25. Time series of the Angel Peak mobile laboratory measurements on June 28.

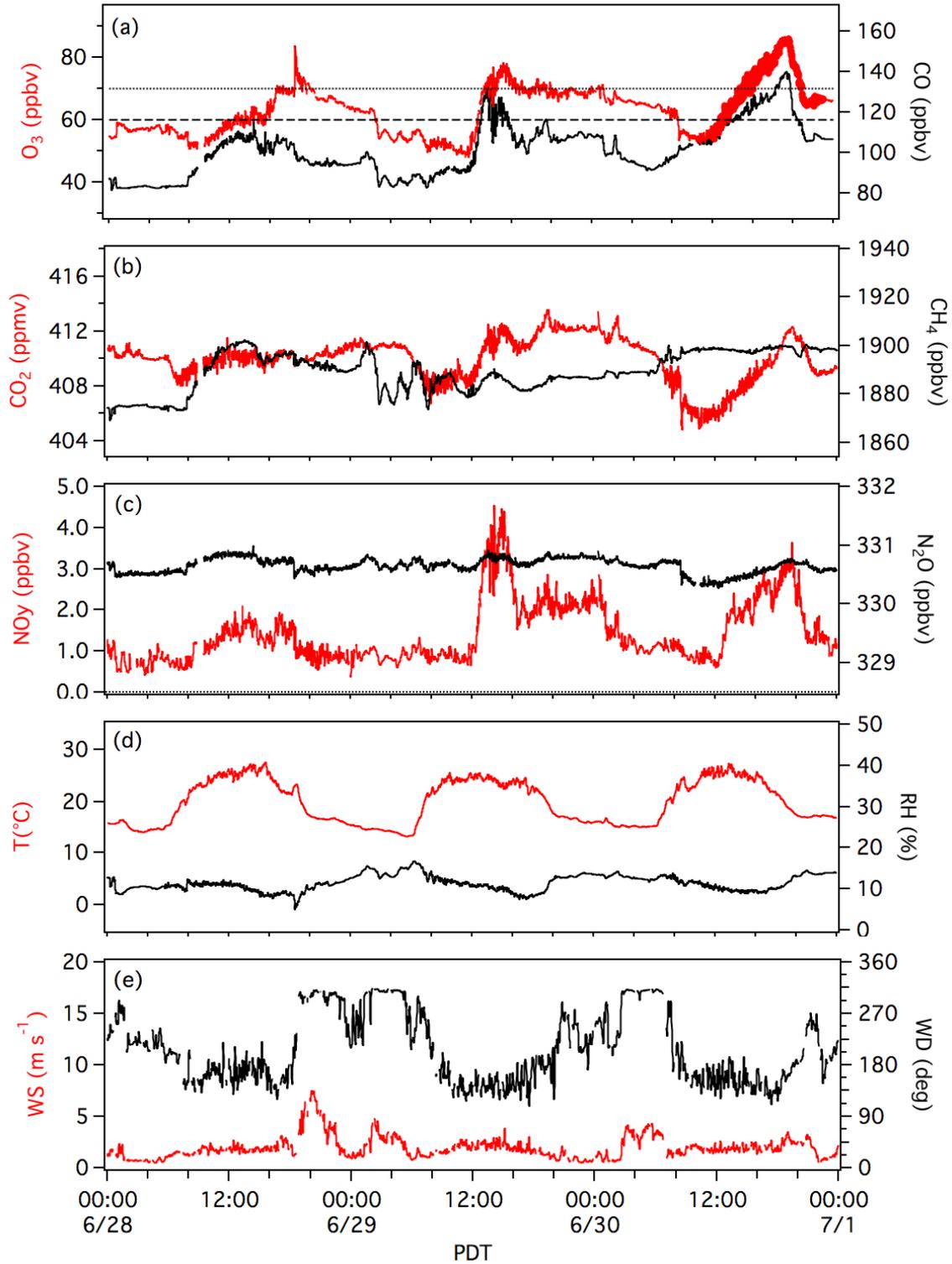
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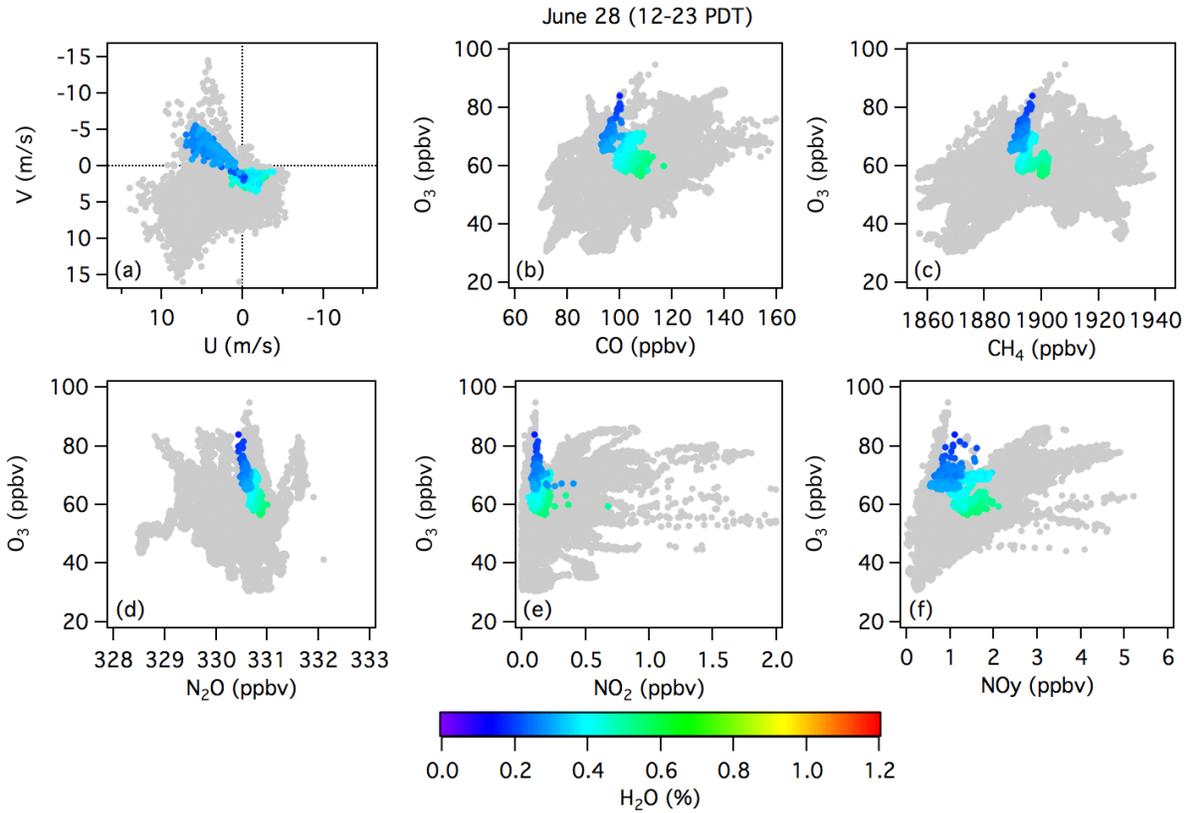
Figure C-26. Time series of the Angel Peak mobile laboratory measurements on June 29.

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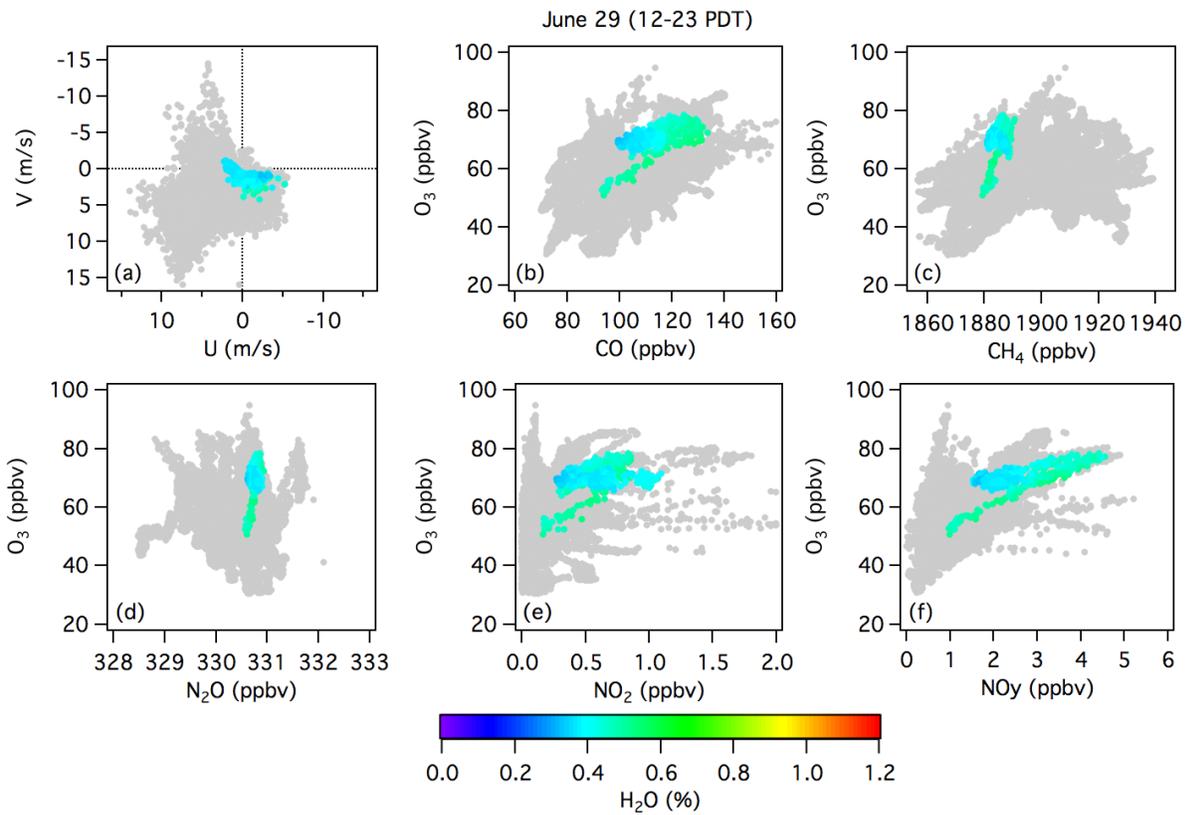


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Figure C-27. Time series of the Angel Peak mobile laboratory measurements on June 30.

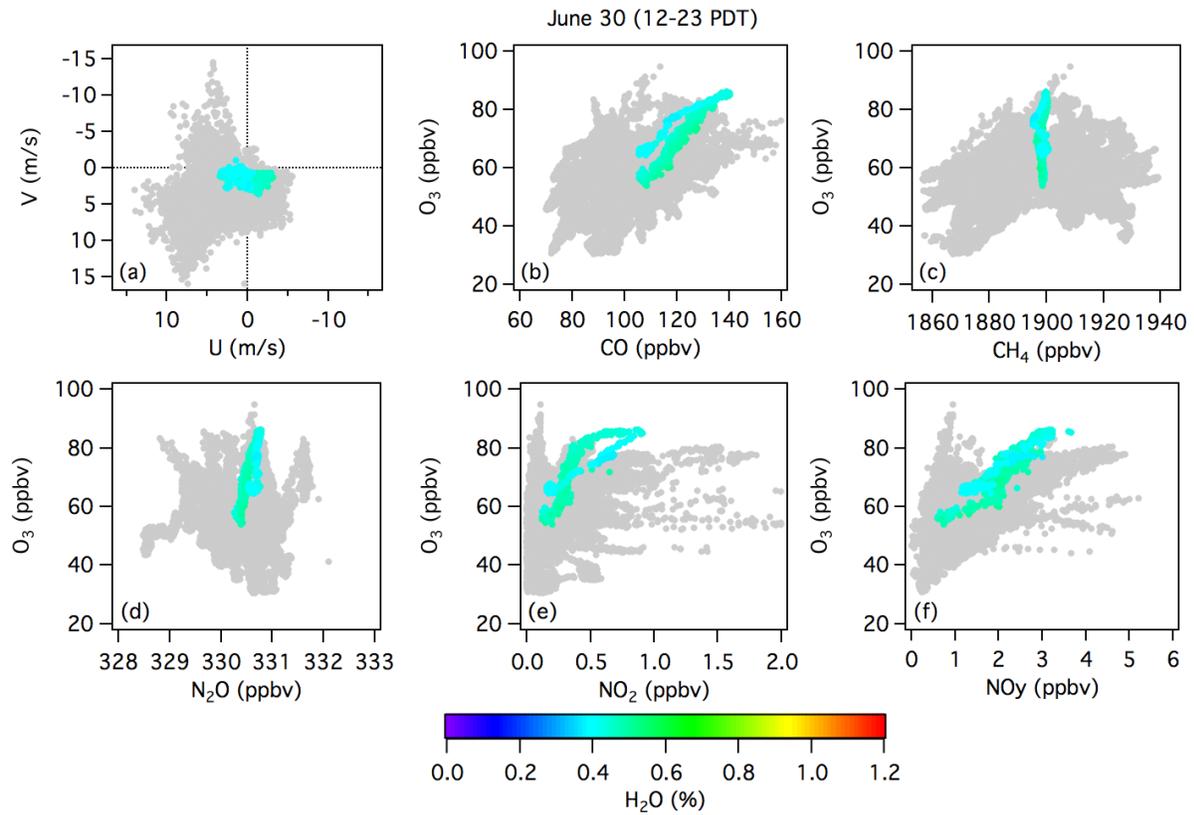


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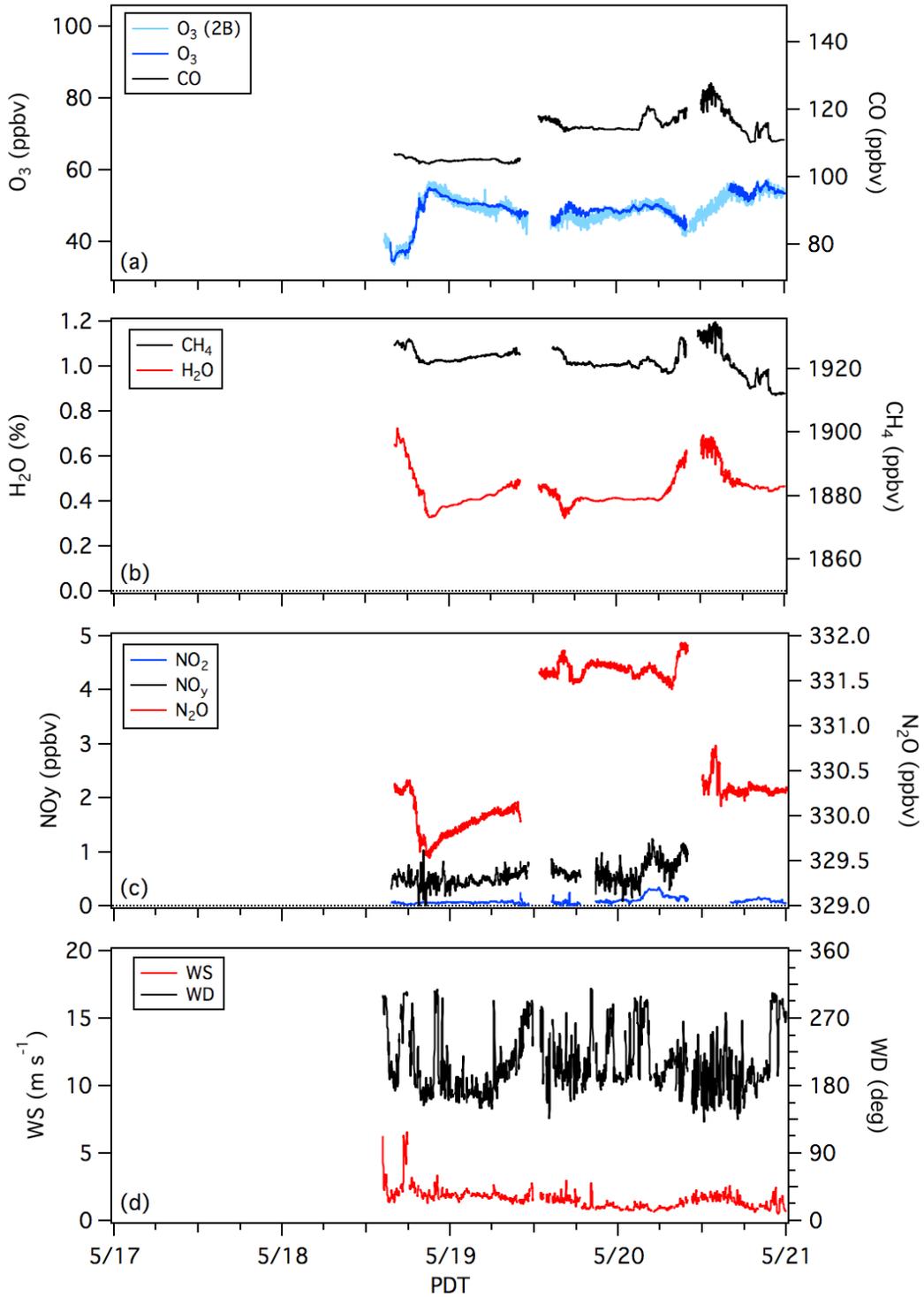


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 2 **Figure C-28.** Scatter plots showing the relationships between O₃ and the other parameters
 3 measured by the mobile laboratory on June 28, 29, and 30.
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1 **2. Angel Peak weekly summary plots**

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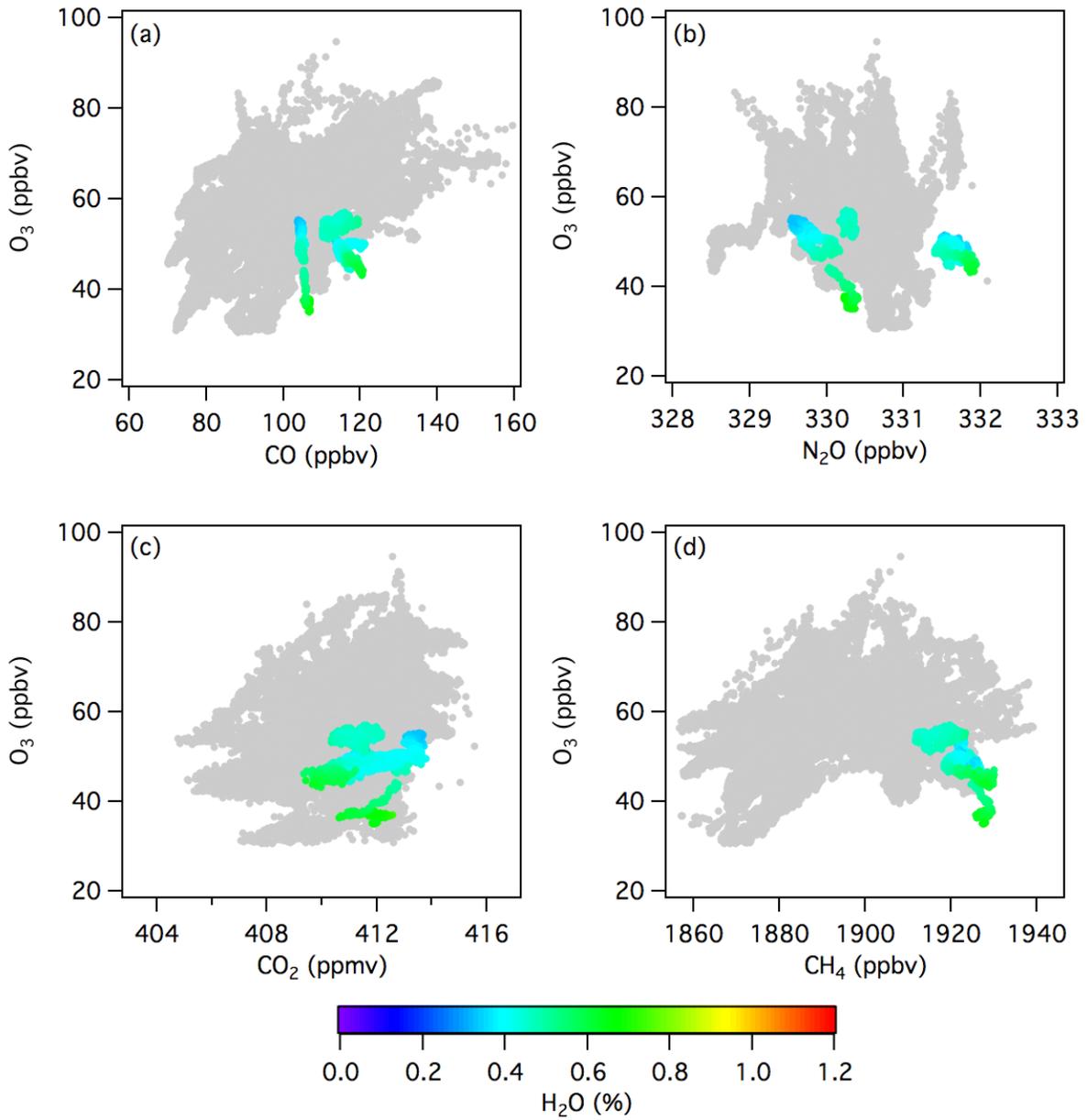
Week 1



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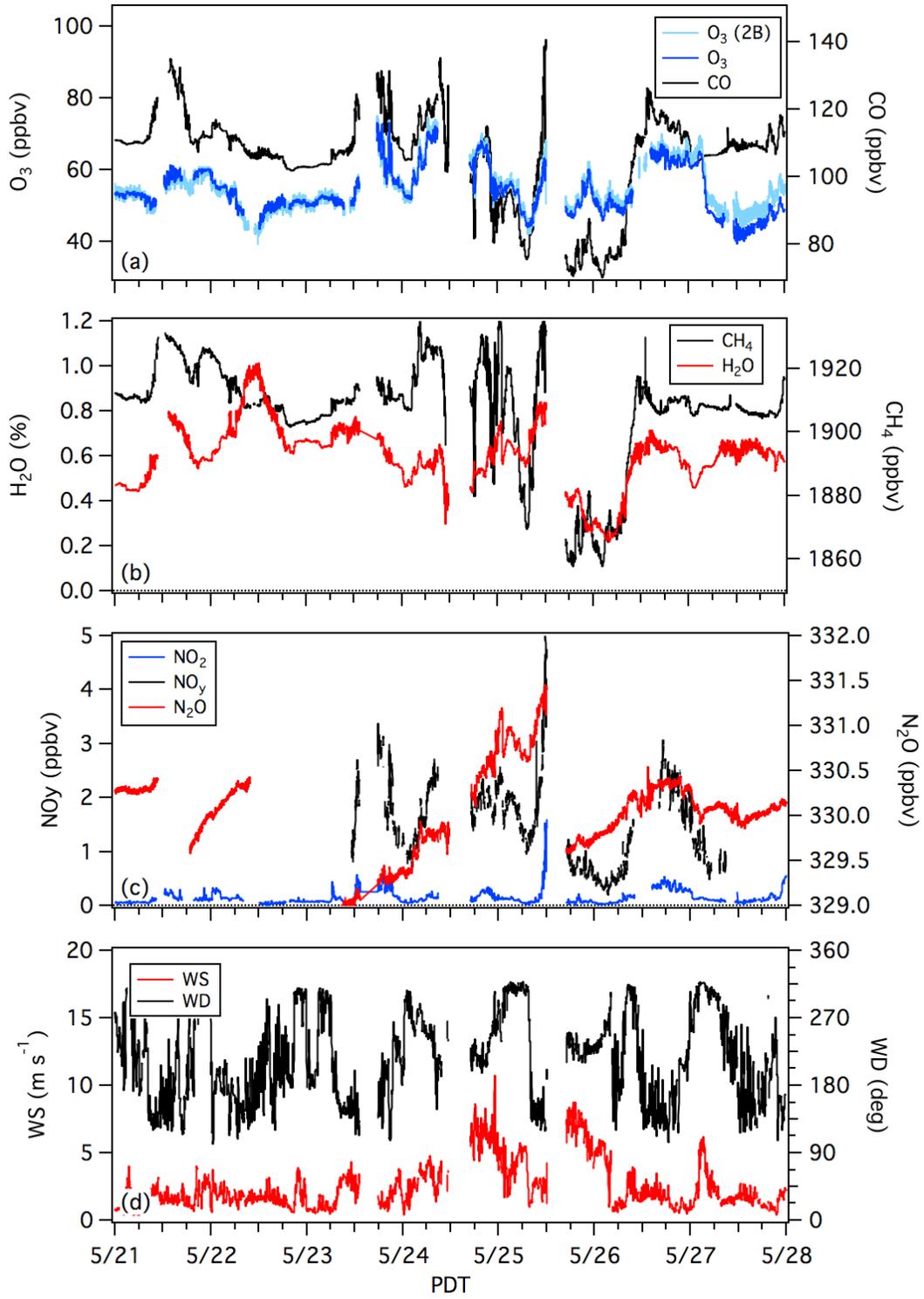
Week 1



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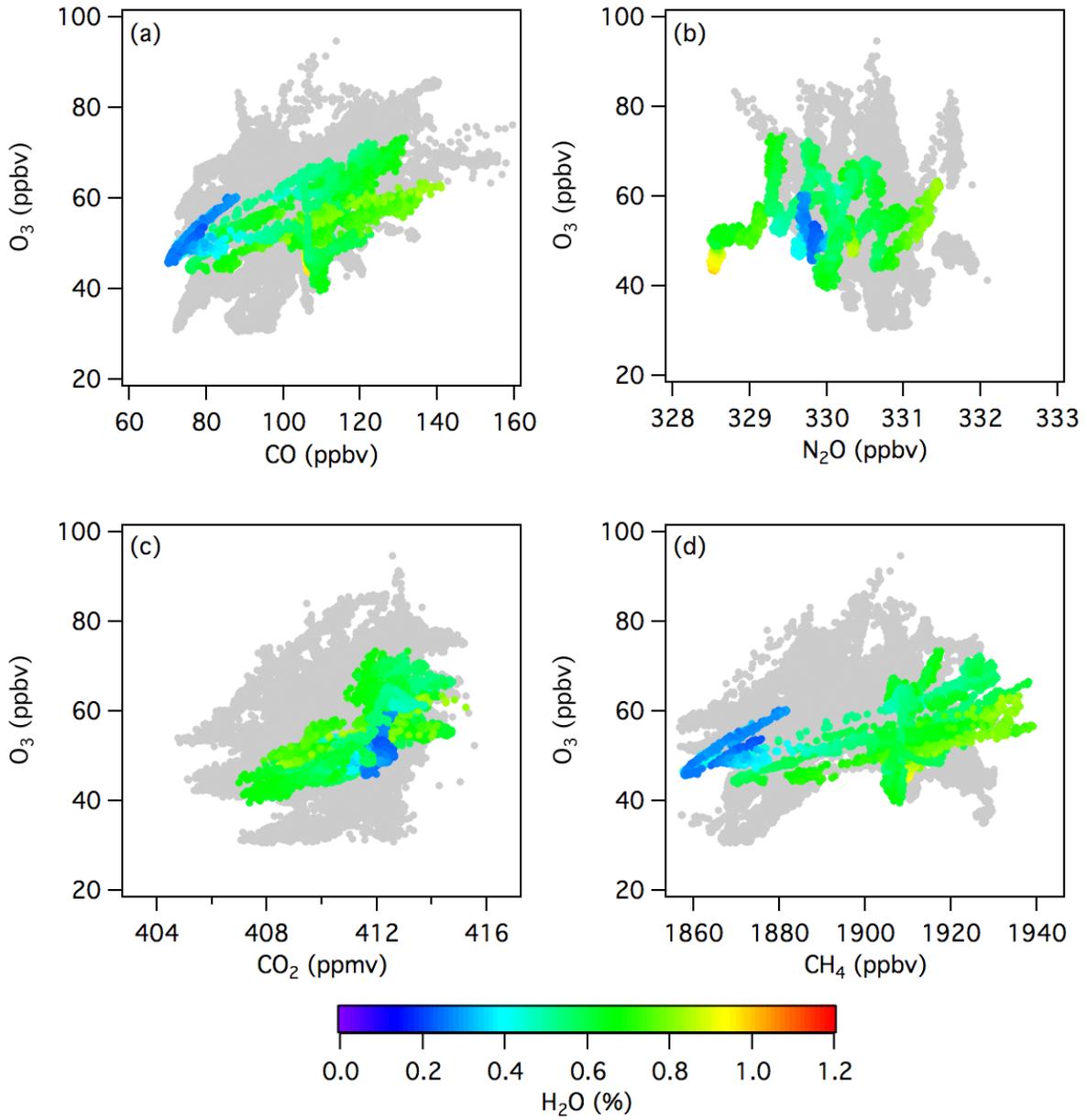
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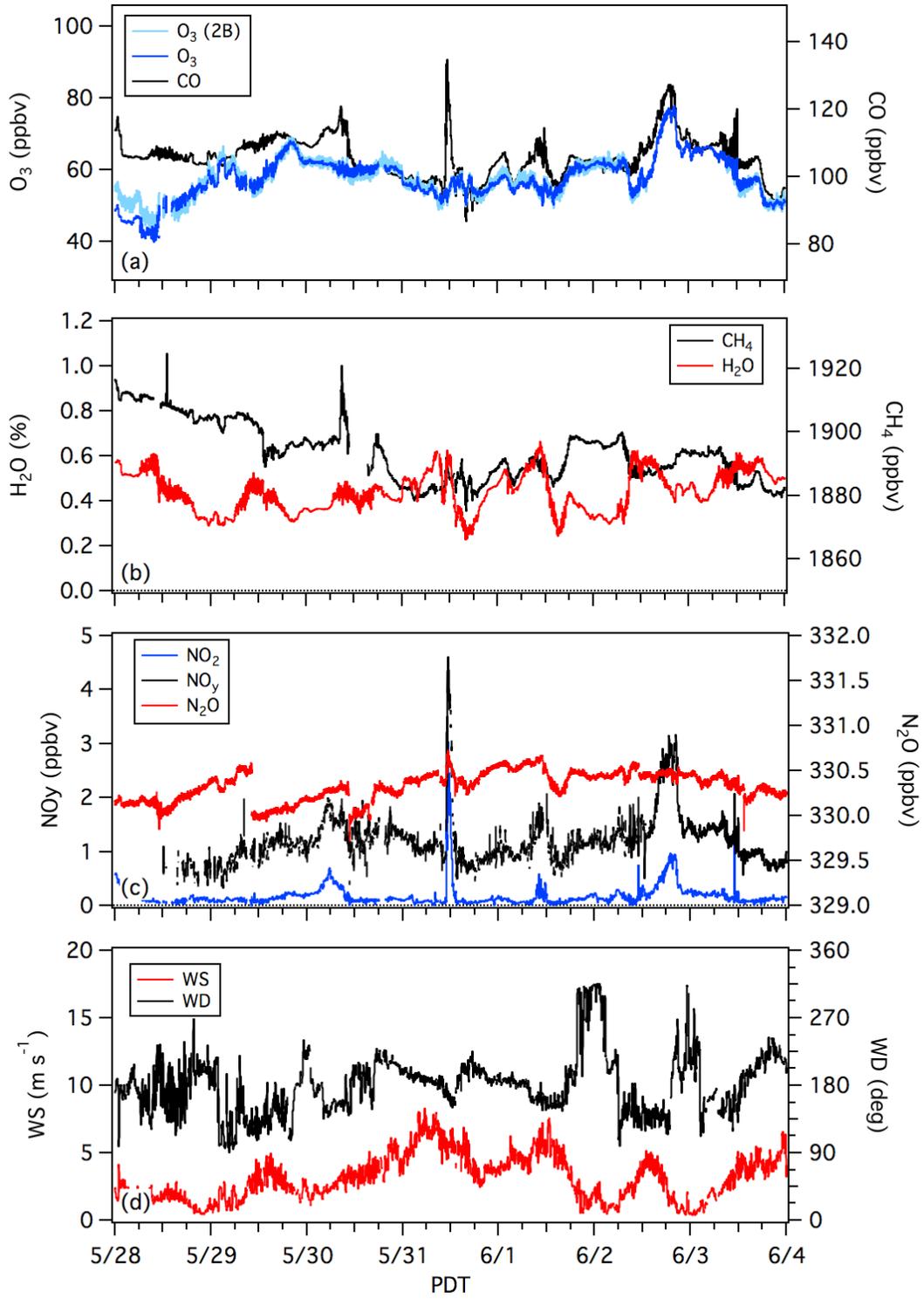
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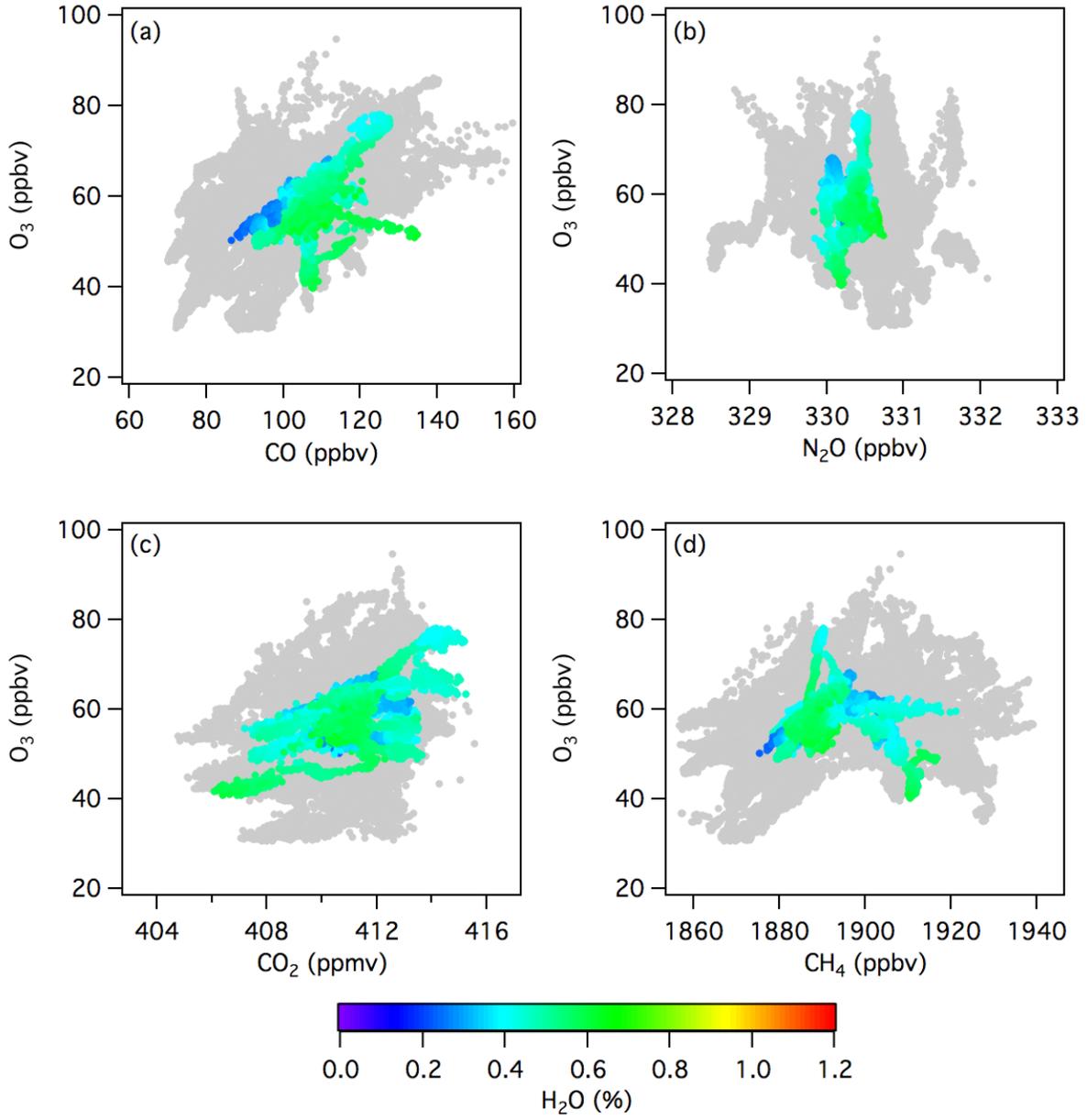
Week 3



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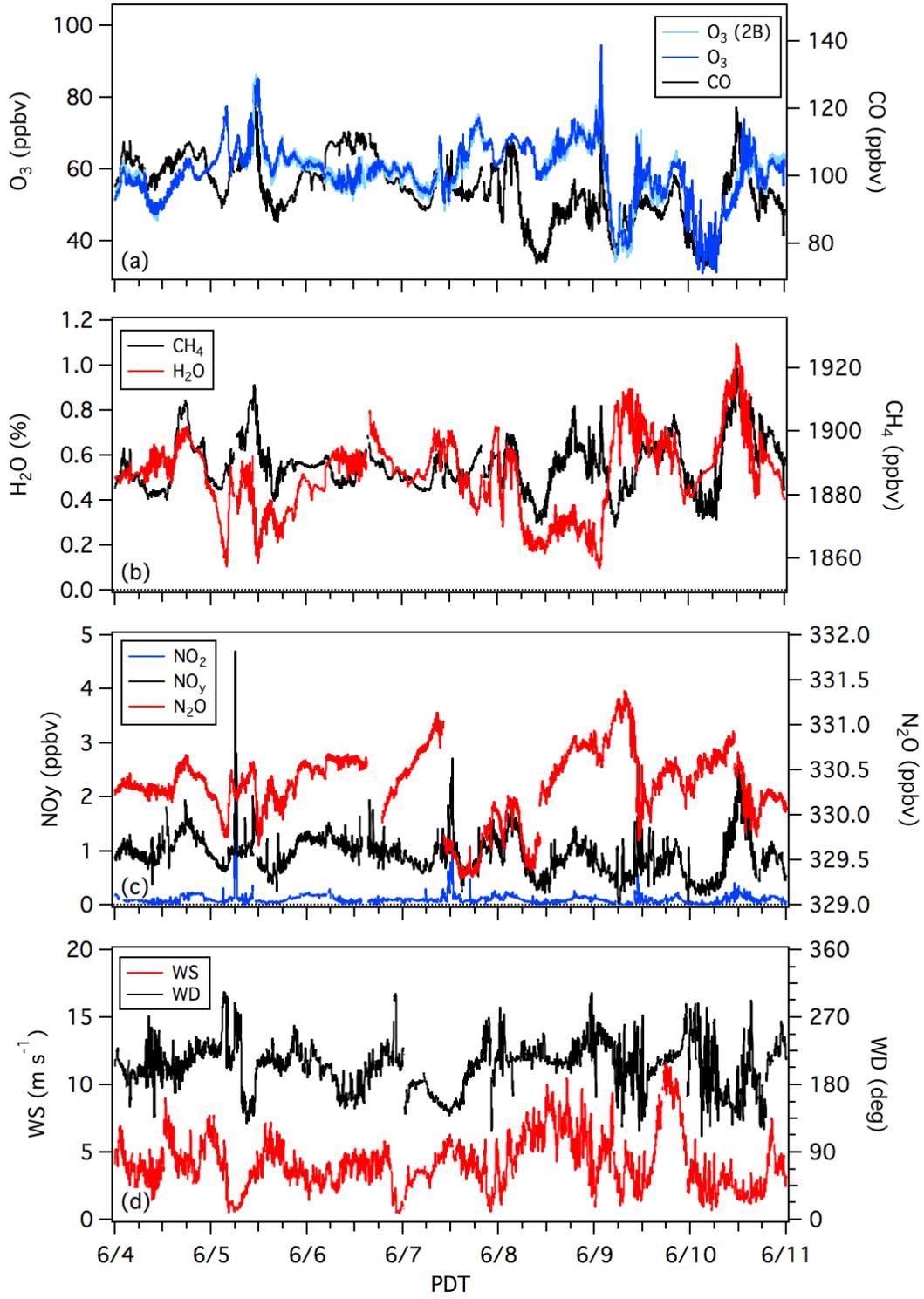
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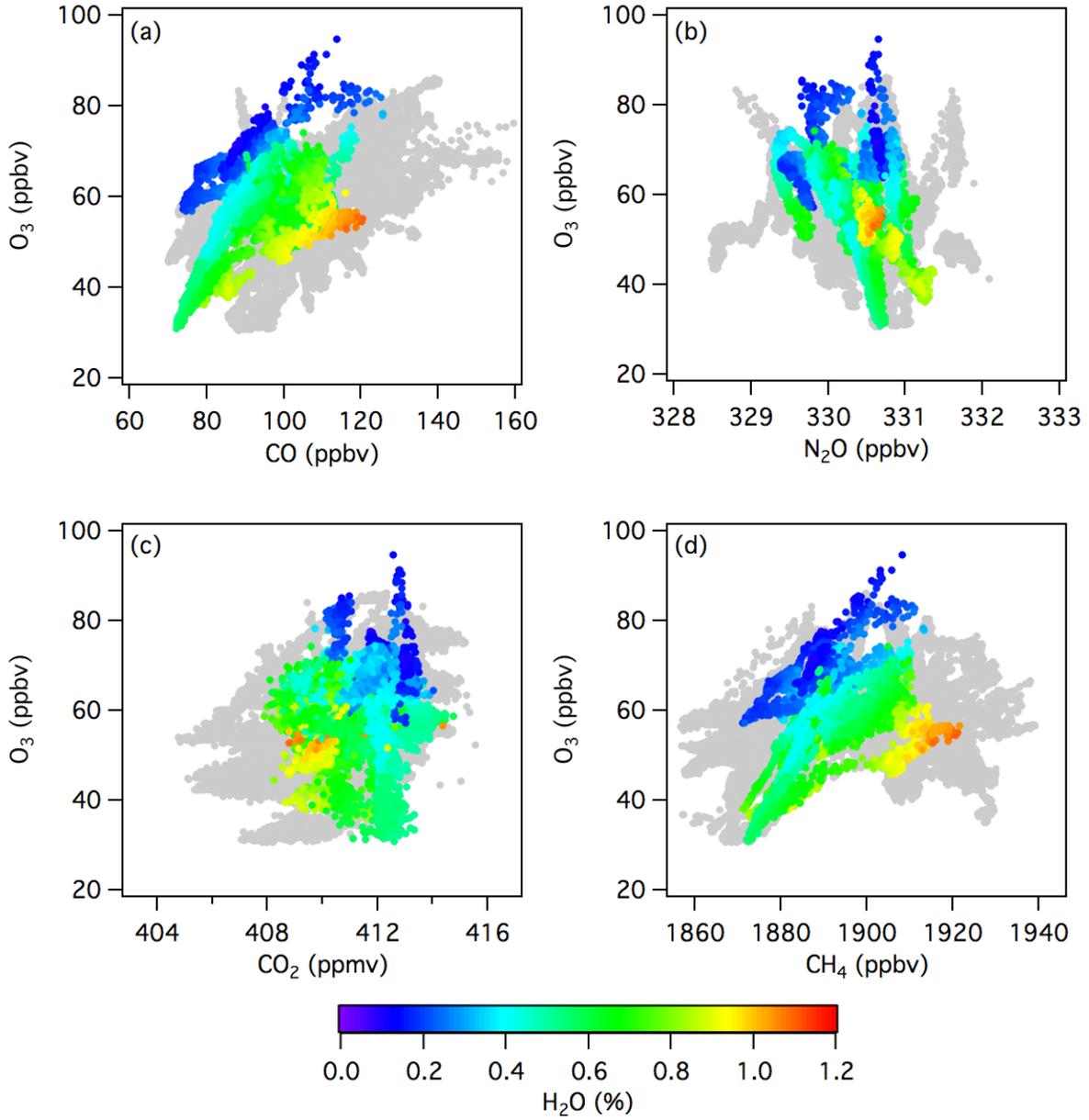
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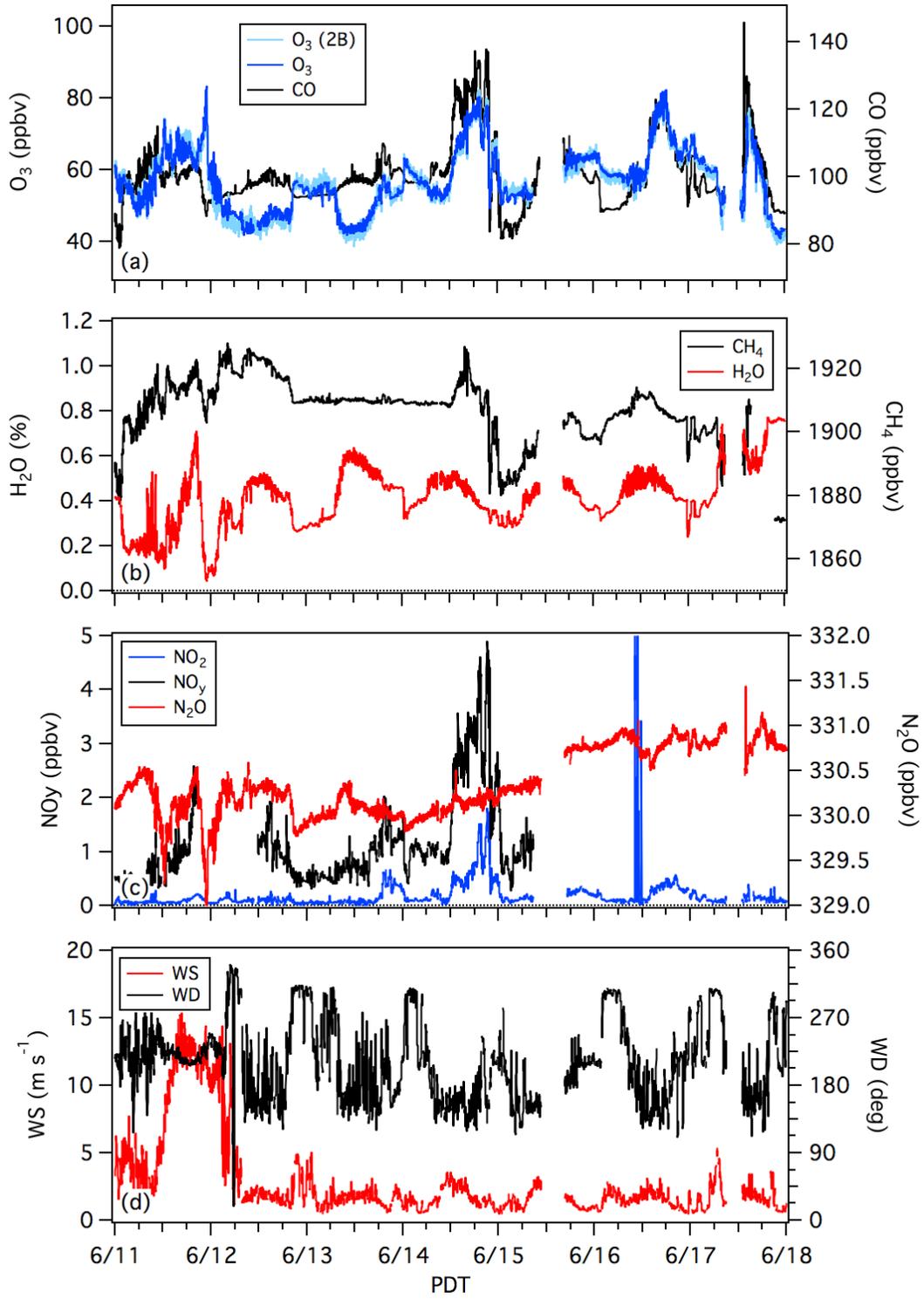
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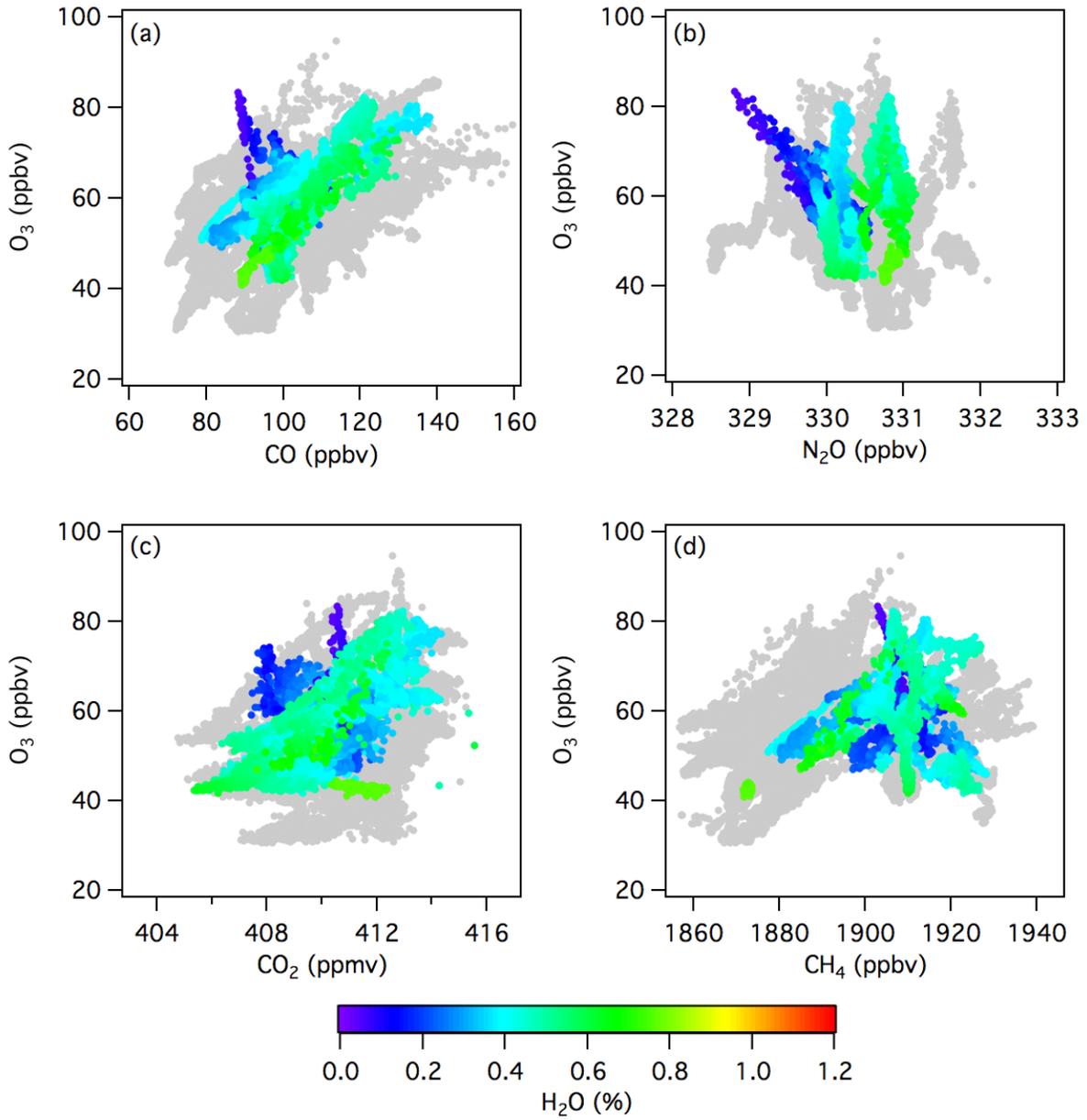
Week 5



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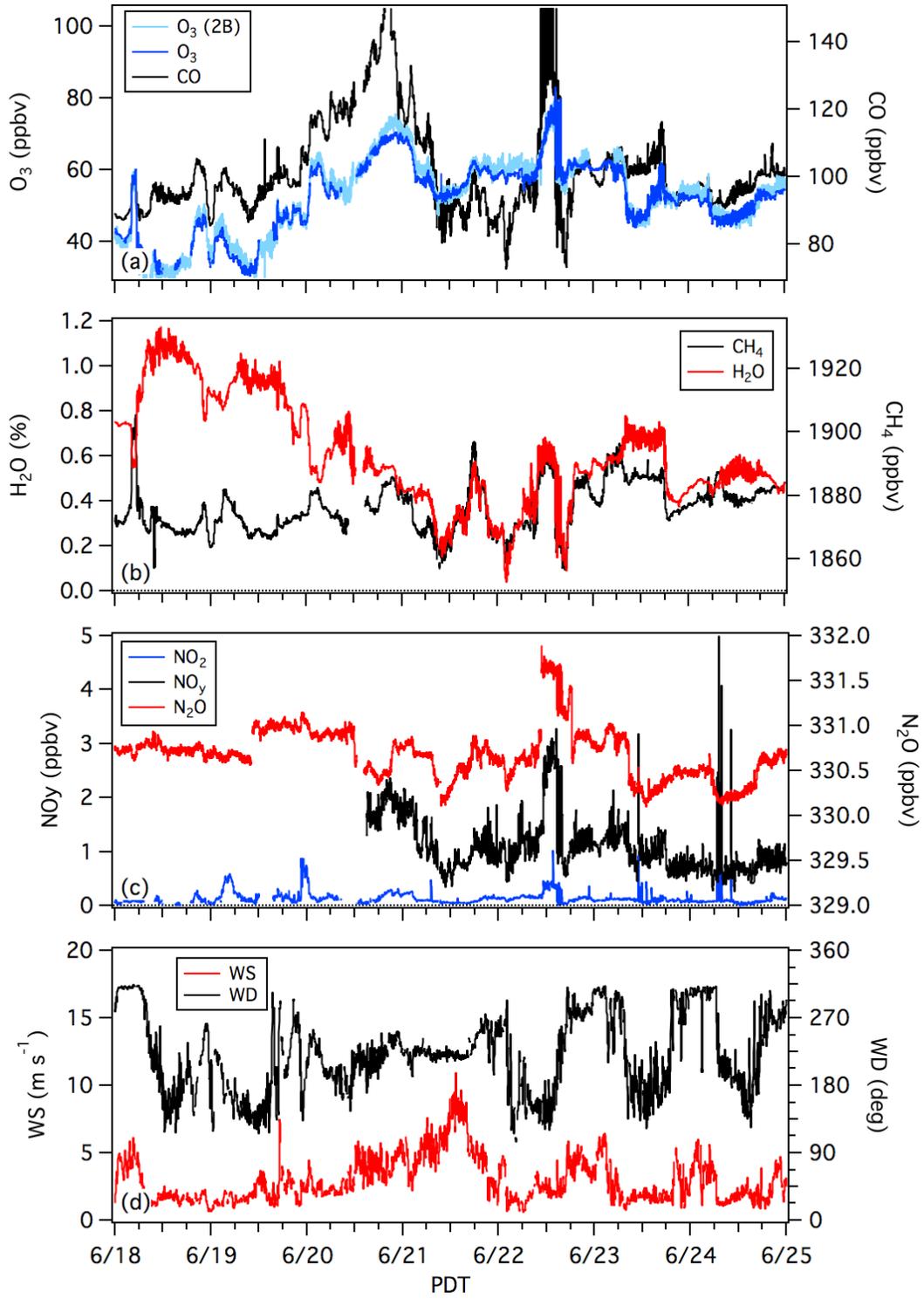
Week 5



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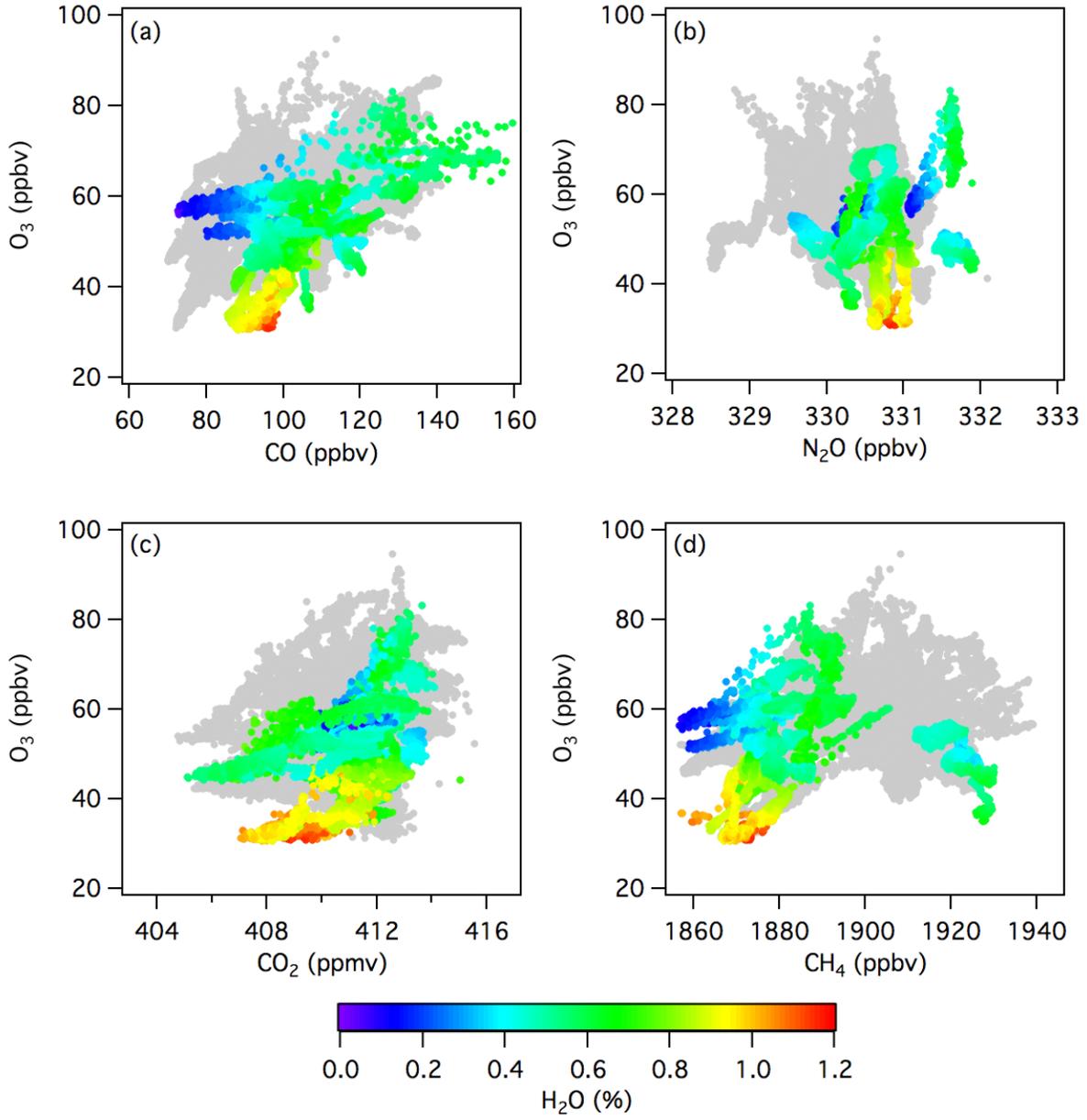
Week 6



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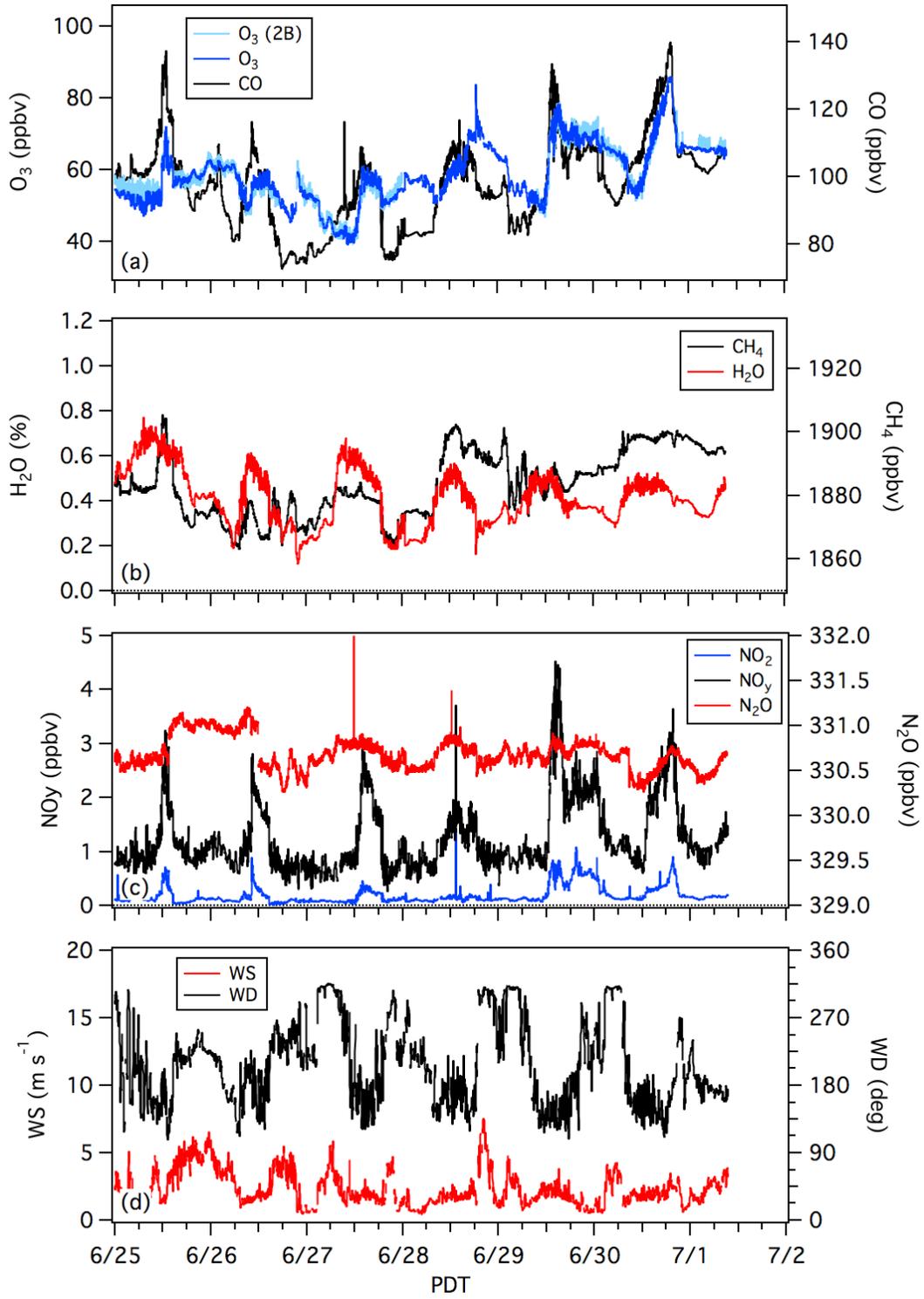
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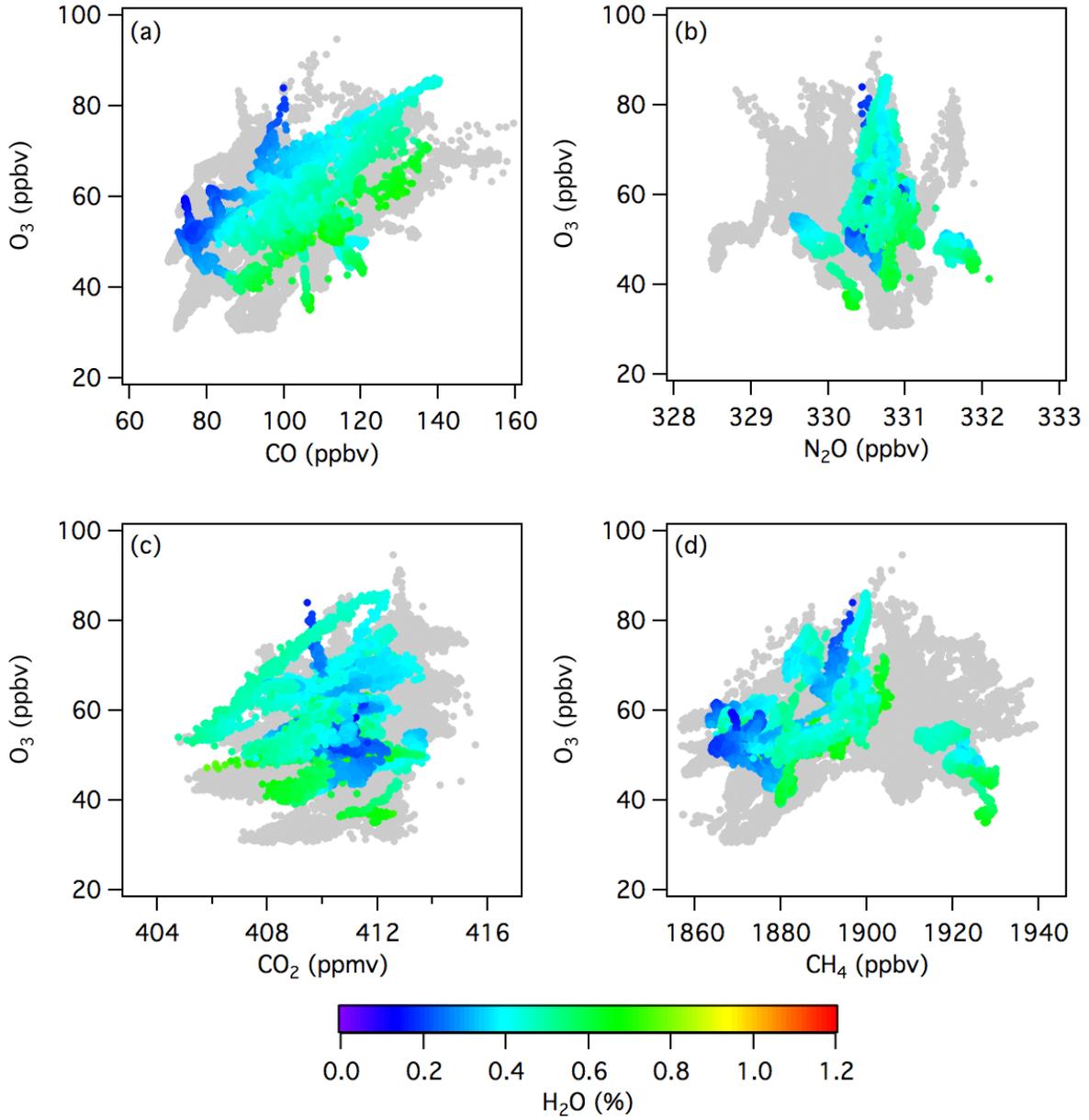
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Week 7



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1 **Appendix D: NOAA/ESRL/GMD Ozonesonde Profiles**

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3 This Appendix contains profile plots showing the O₃, RH, and potential temperature profiles
4 from the ozonesonde ascents during the FAST-LVOS IOPs. The horizontal dashed lines show the
5 top of boundary layer inferred from the measurements.

6

7 The complete data are available at:

8 <https://www.esrl.noaa.gov/csd/groups/csd3/measurements/fastlvos/ozonesondes/>

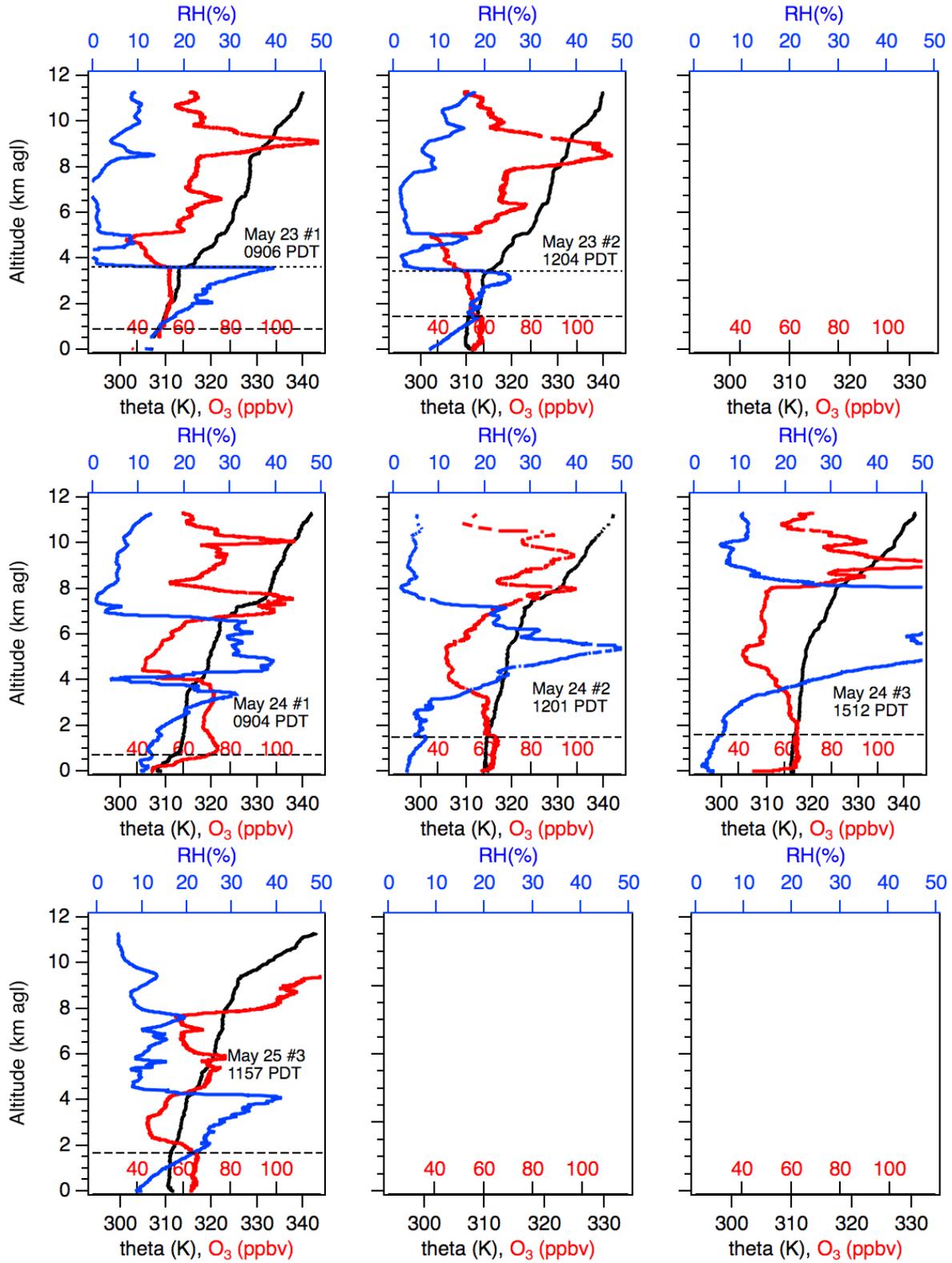
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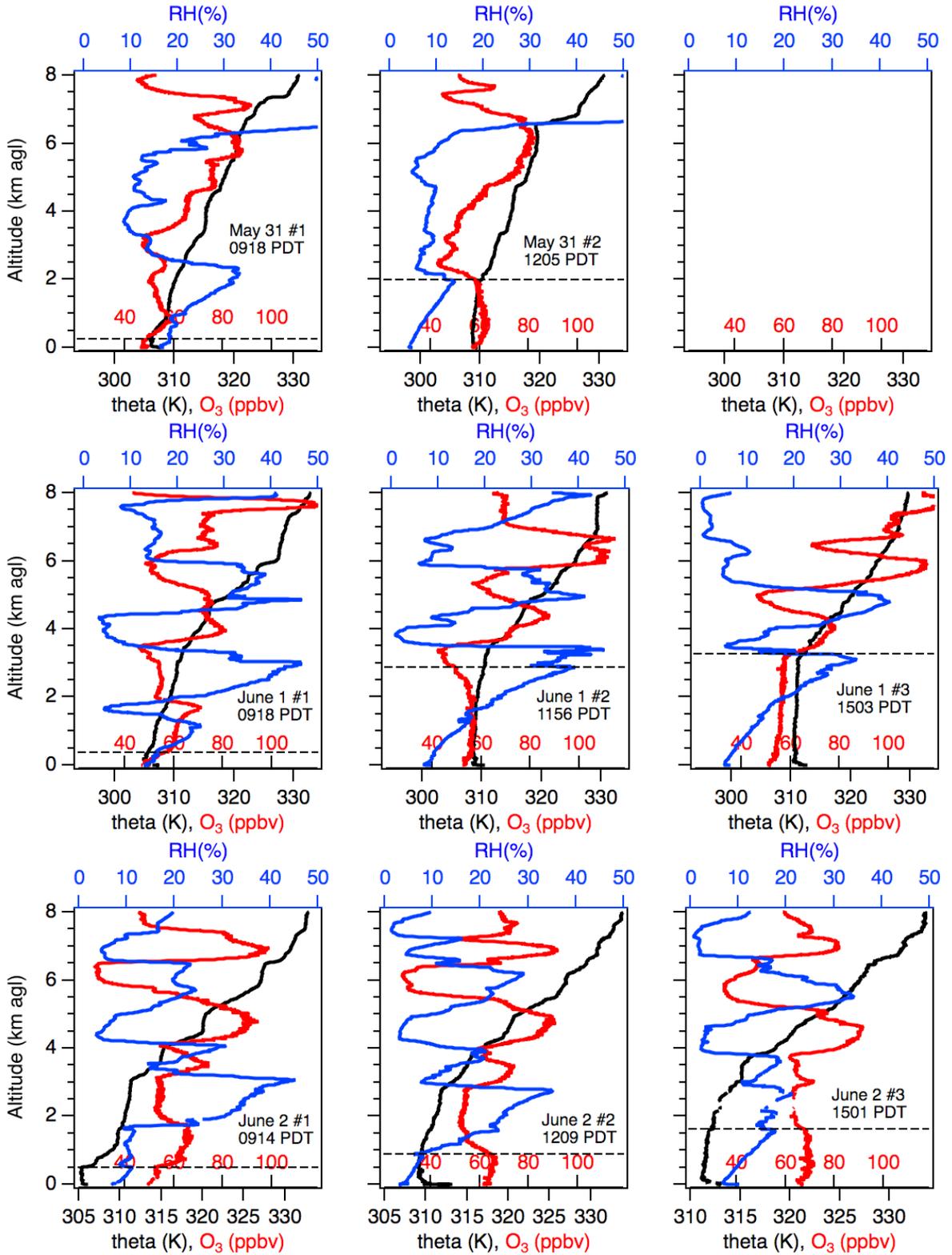
IOP 1



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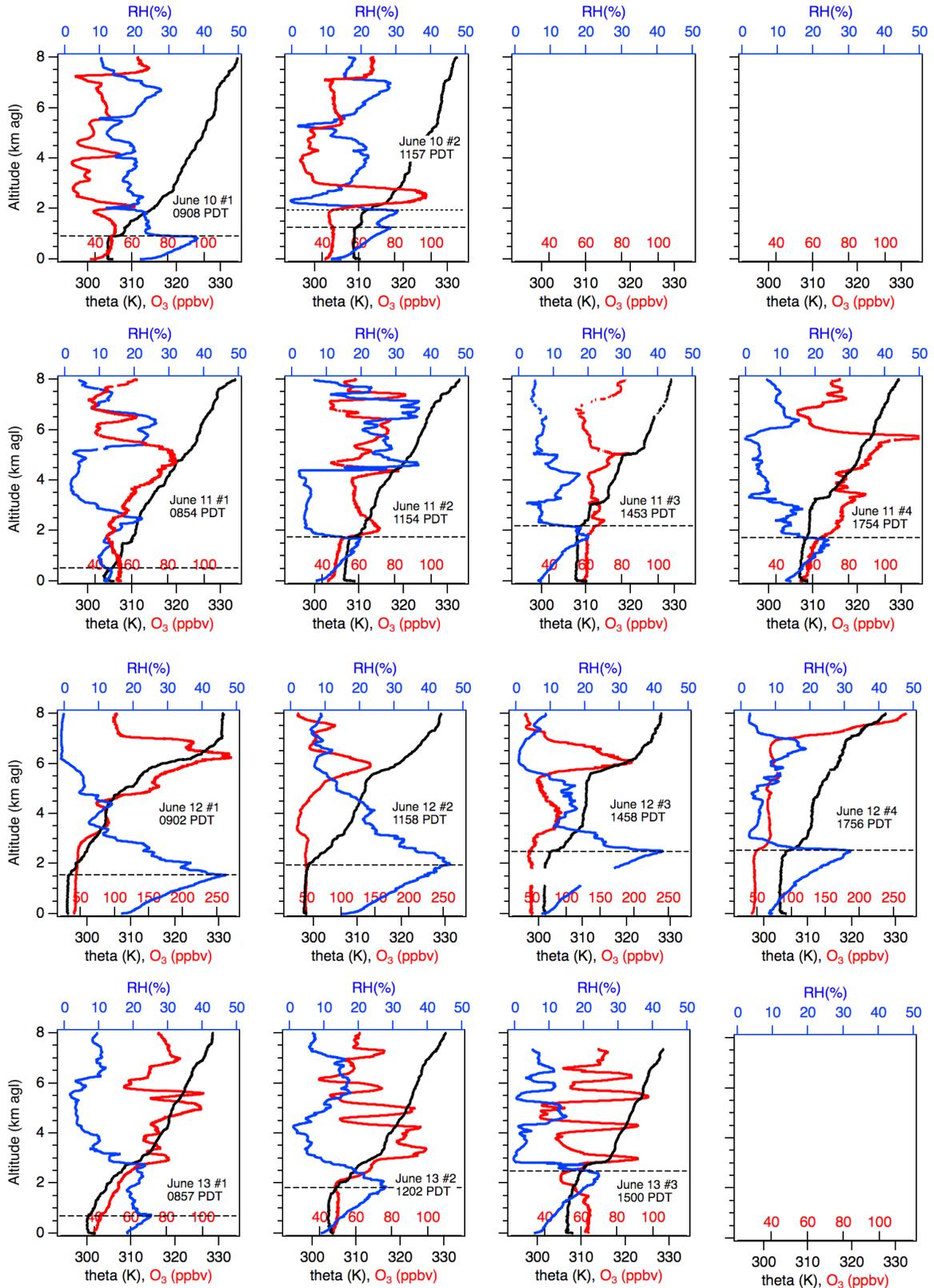
IOP2



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IOP3

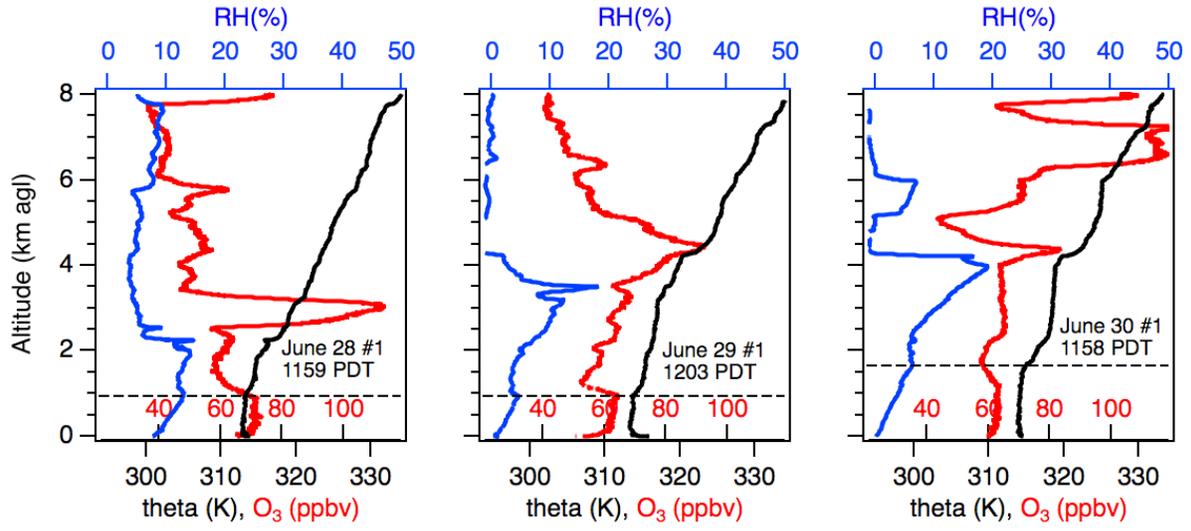


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IOP4



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1 **Appendix E: Scientific Aviation Airborne Measurements**

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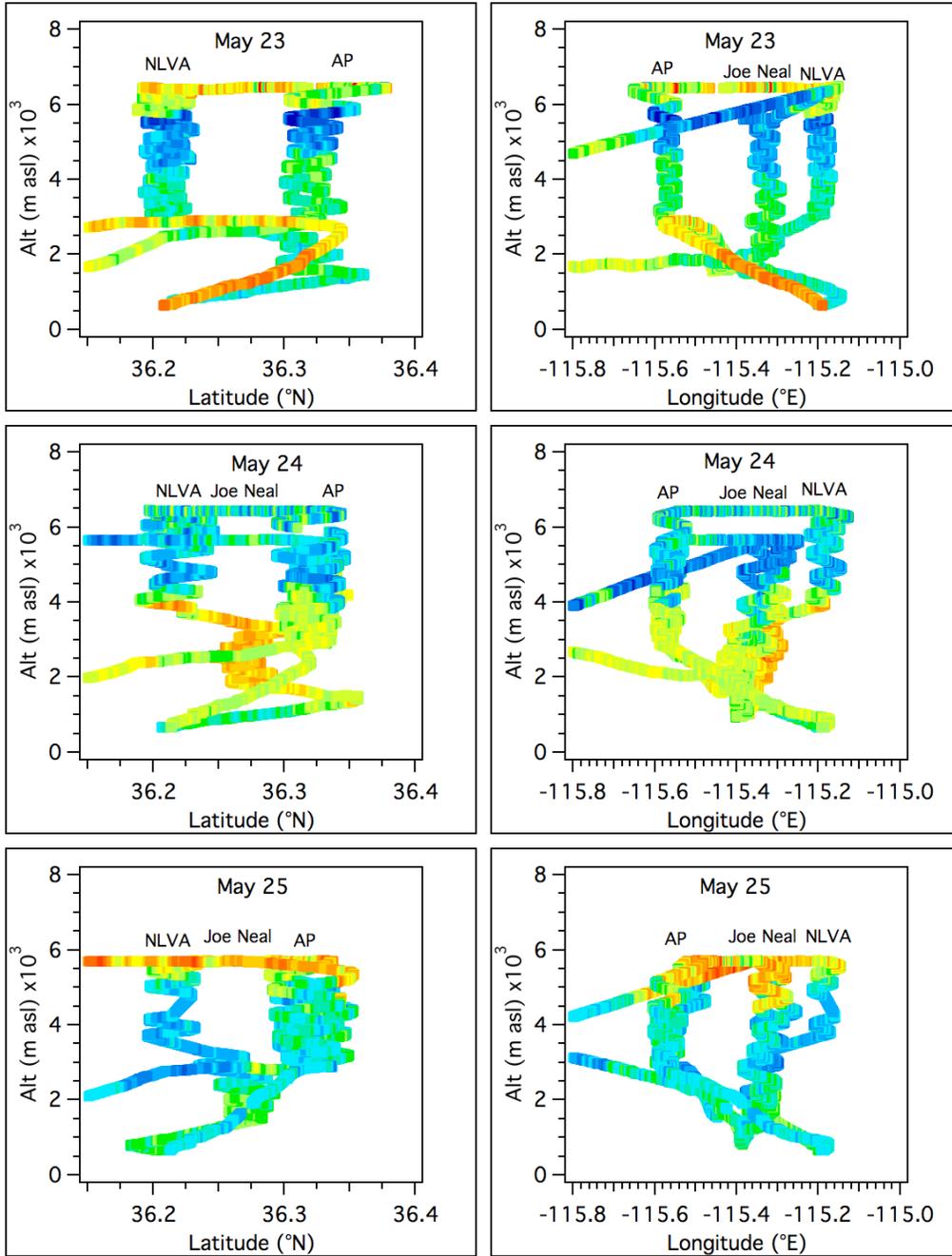
3 This Appendix contains flight track maps and profile plots showing the Scientific Aviation O₃
4 measurements during the FAST-LVOS IOPs. The daily maps shown first also display the MDA8
5 ozone measured by the regulatory monitors; both these and the flight tracks are colorized using
6 the TOLNet color scale. The altitude-latitude (northern Clark County only) and altitude-
7 longitude plots shown next are also colorized by ozone using this scale. The final set of plots for
8 each IOP shows the relative humidity, ozone, and methane profiles measured on the late
9 morning ascents from and afternoon descents into the NLVA.

10

11 The complete data files are available on request from andrew.o.langford@noaa.gov.

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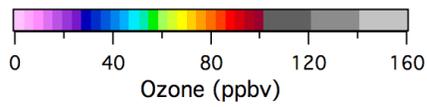
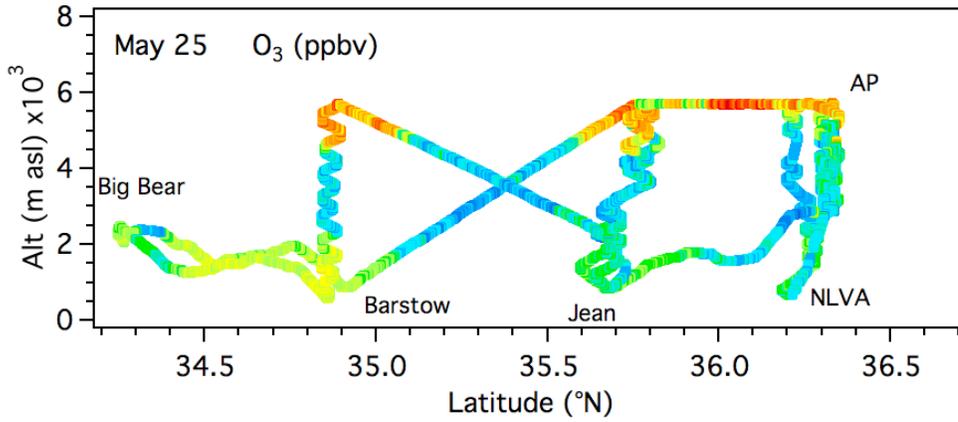
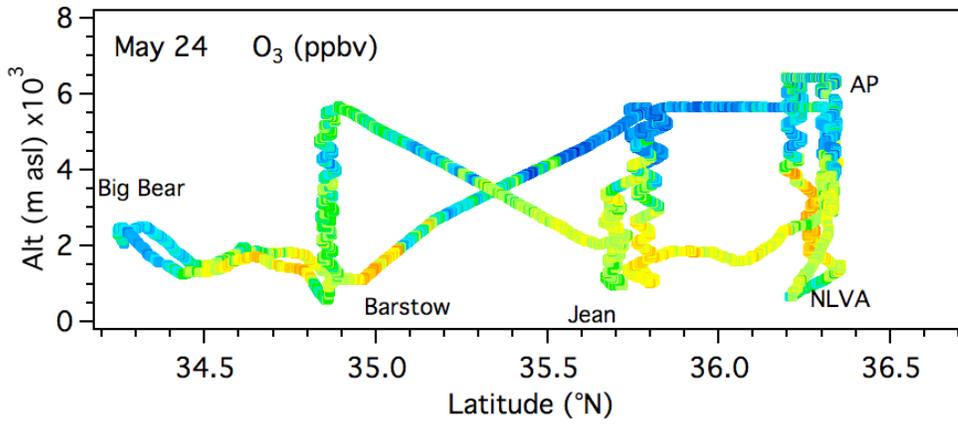
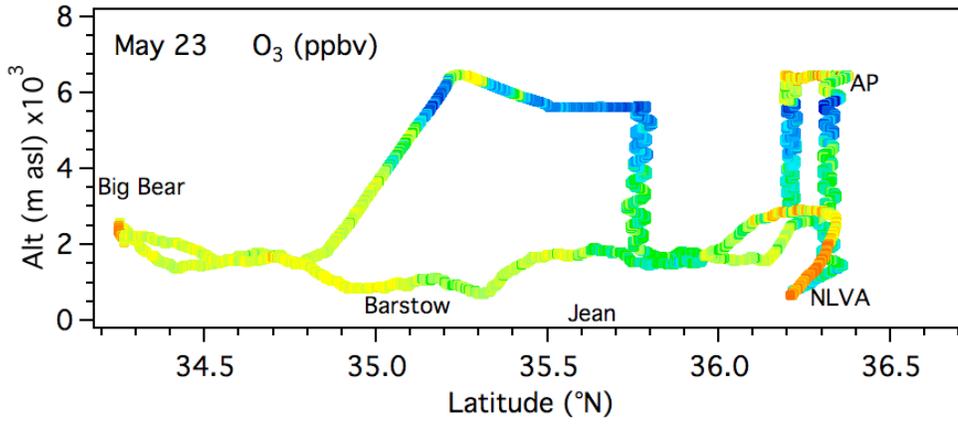
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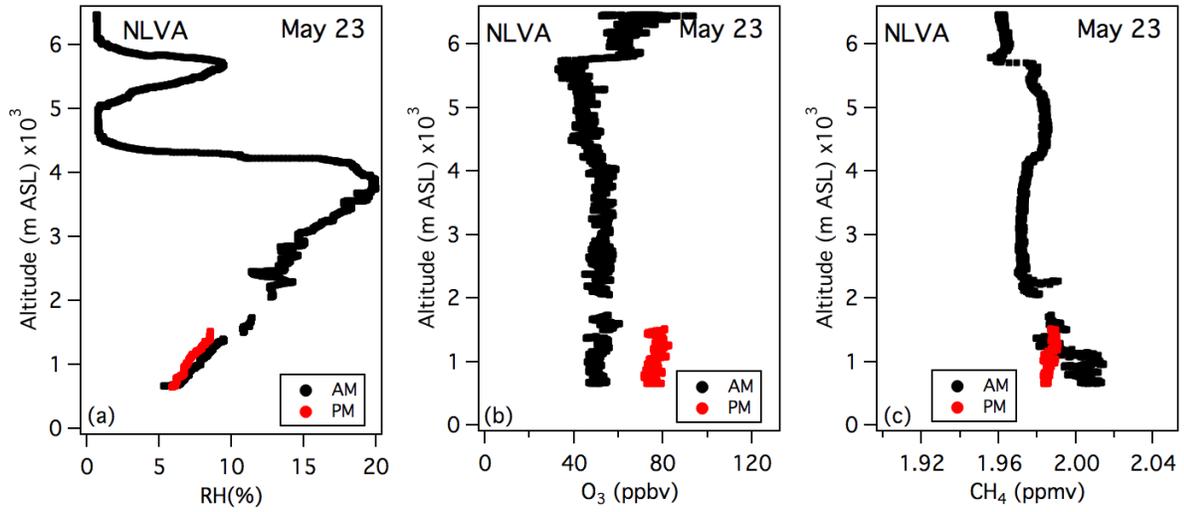
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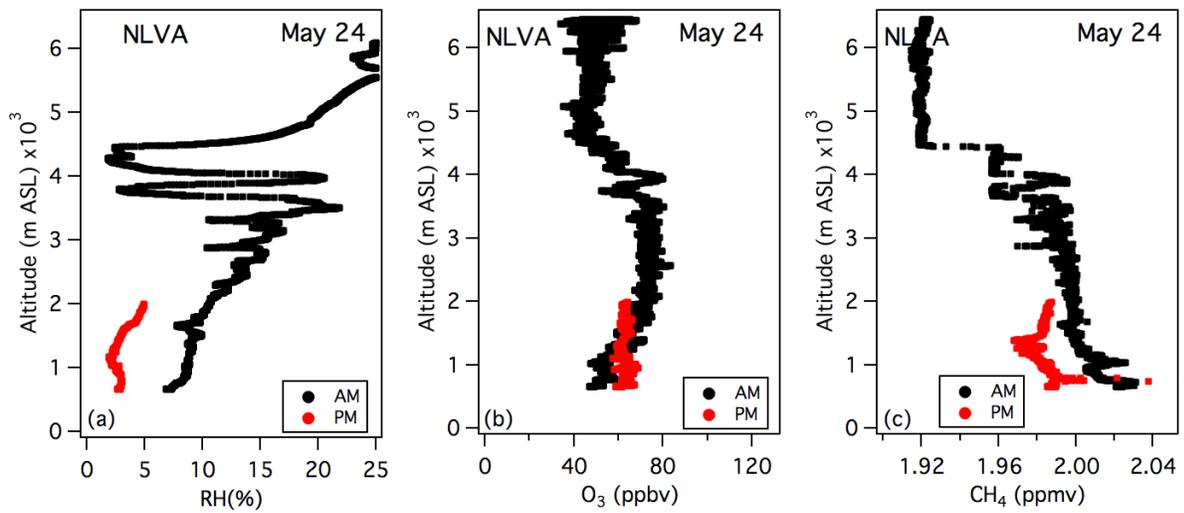
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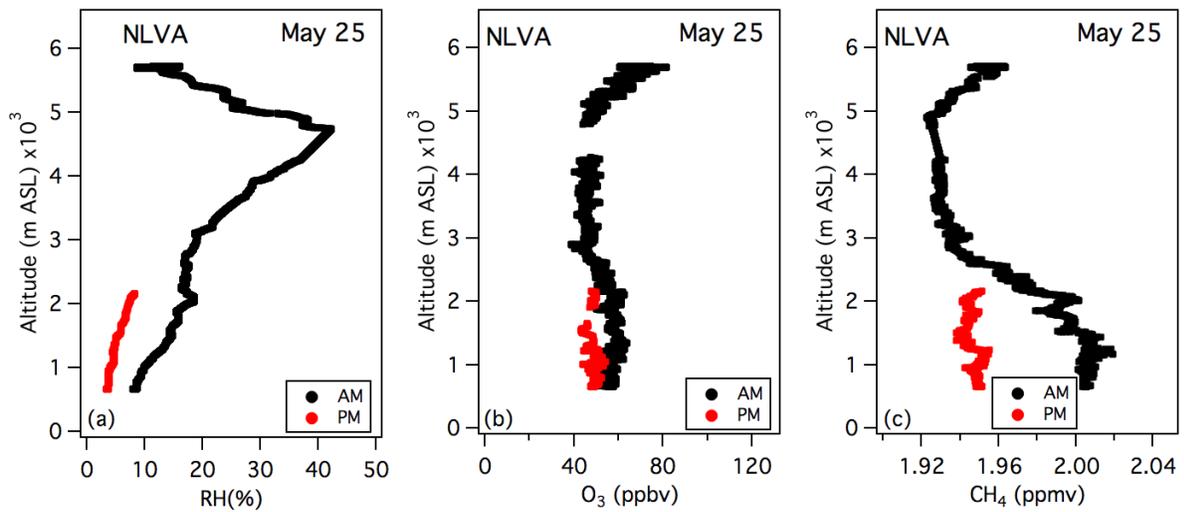
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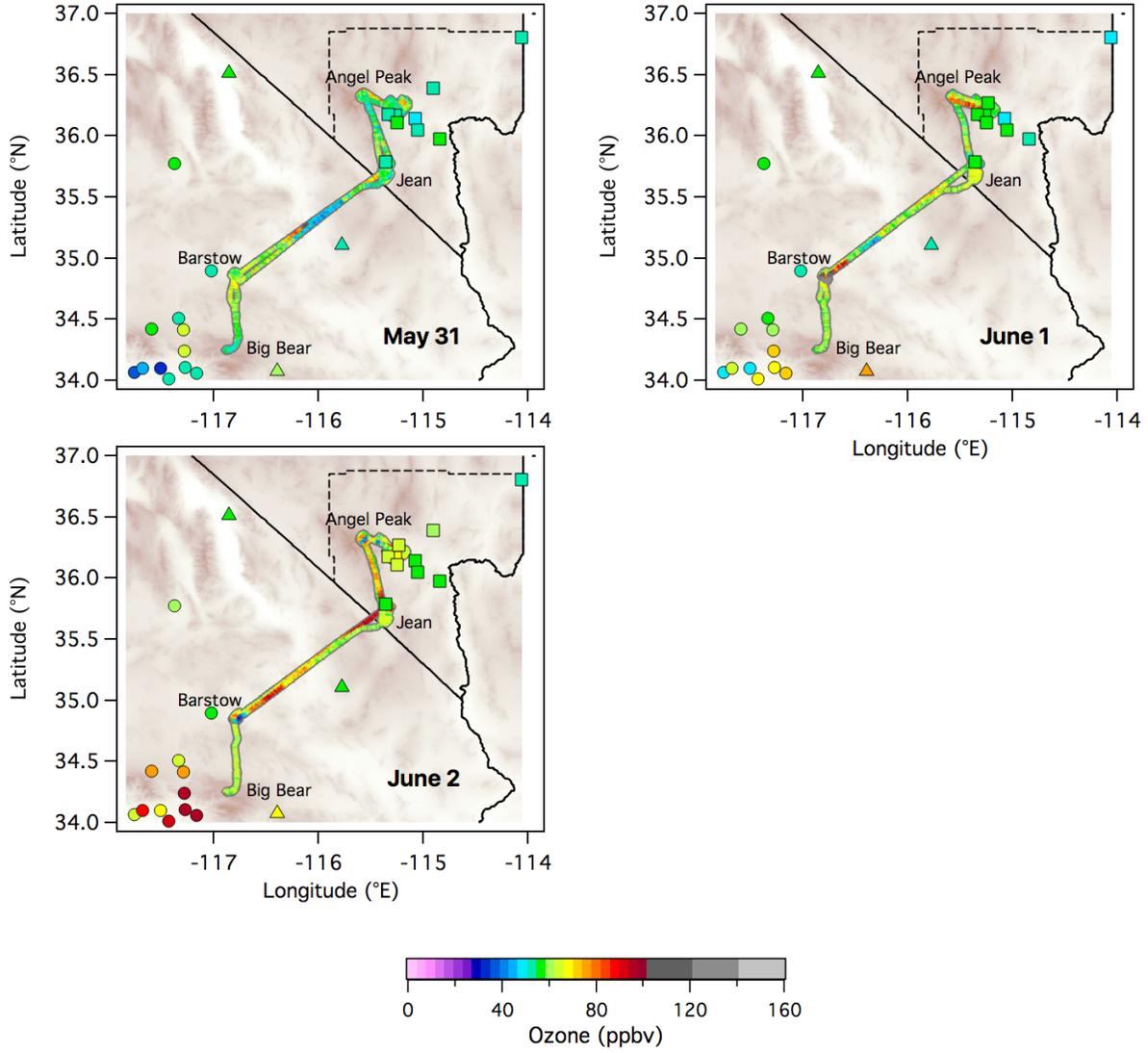
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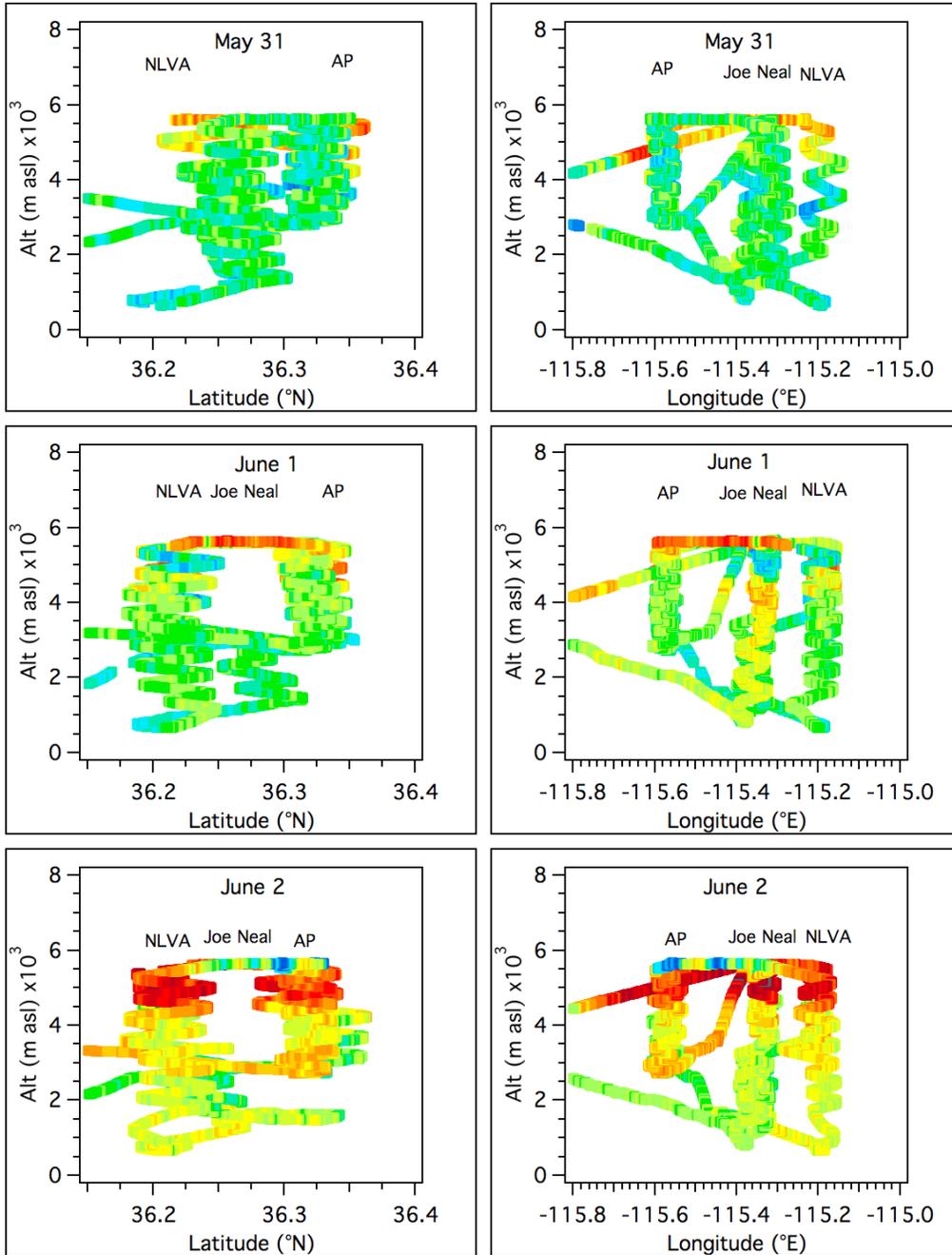
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IOP2



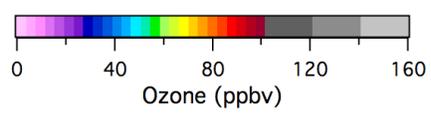
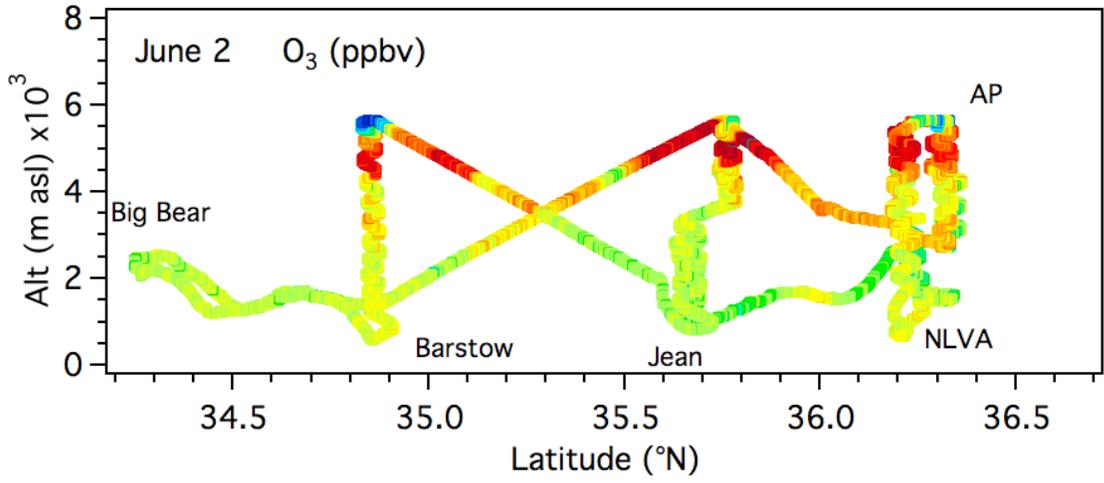
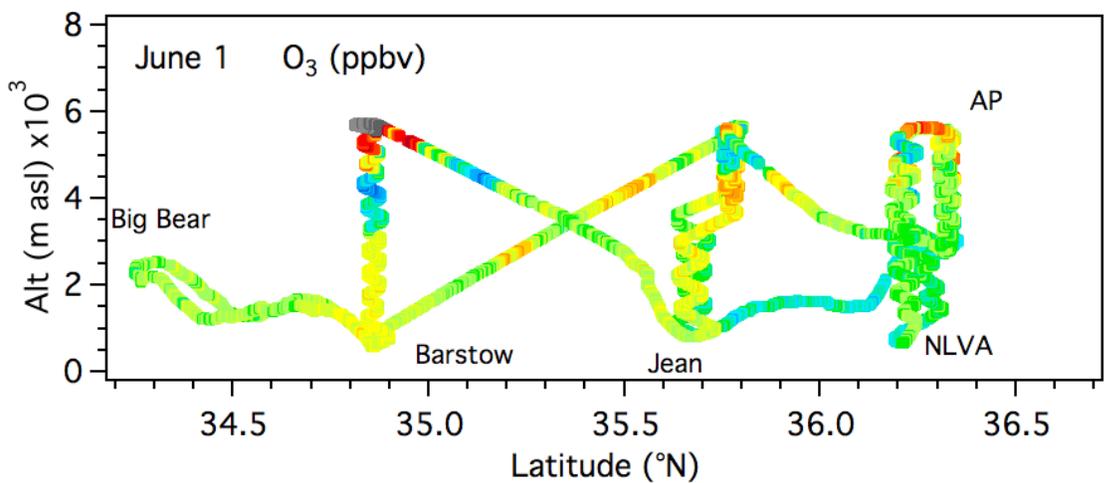
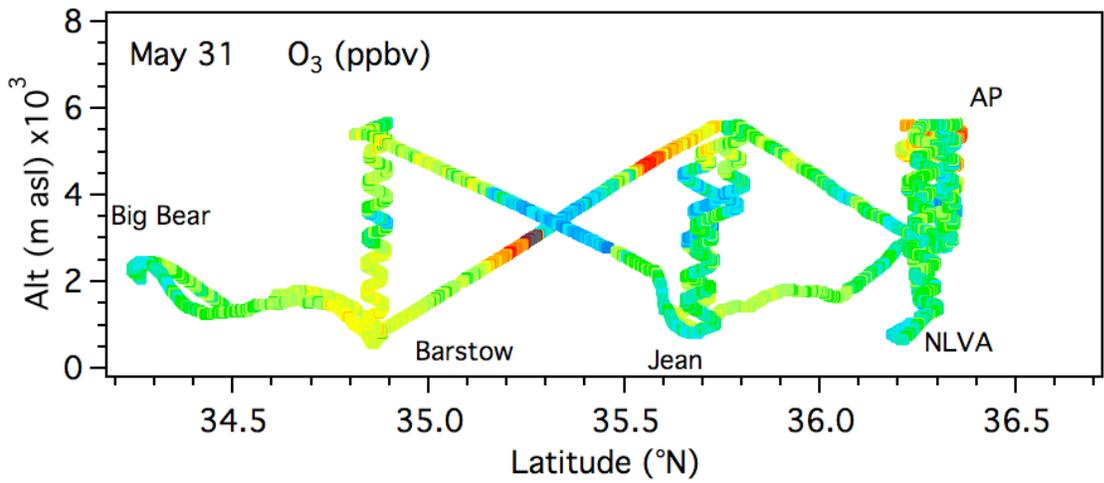
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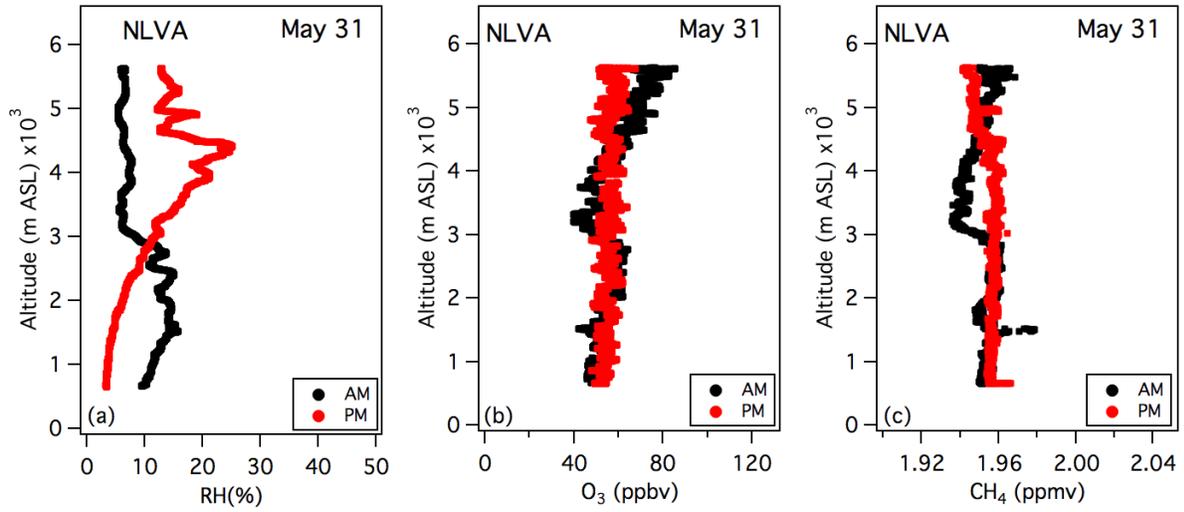


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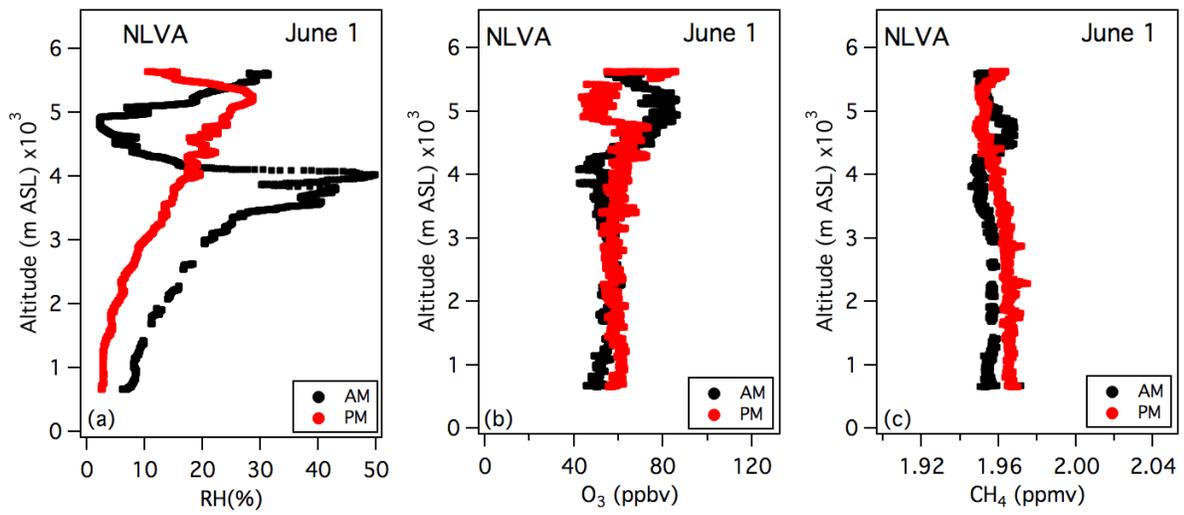
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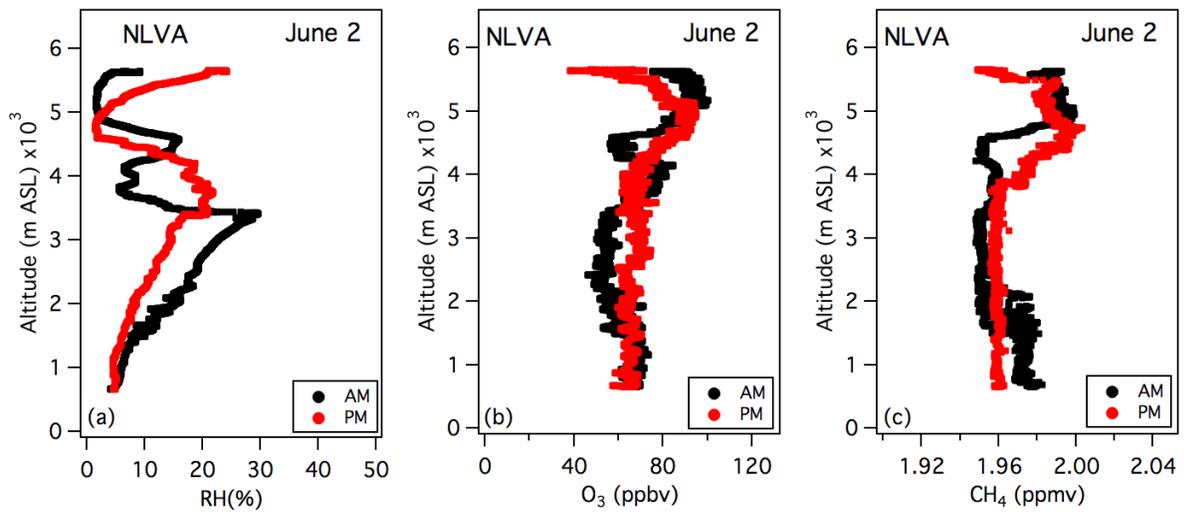
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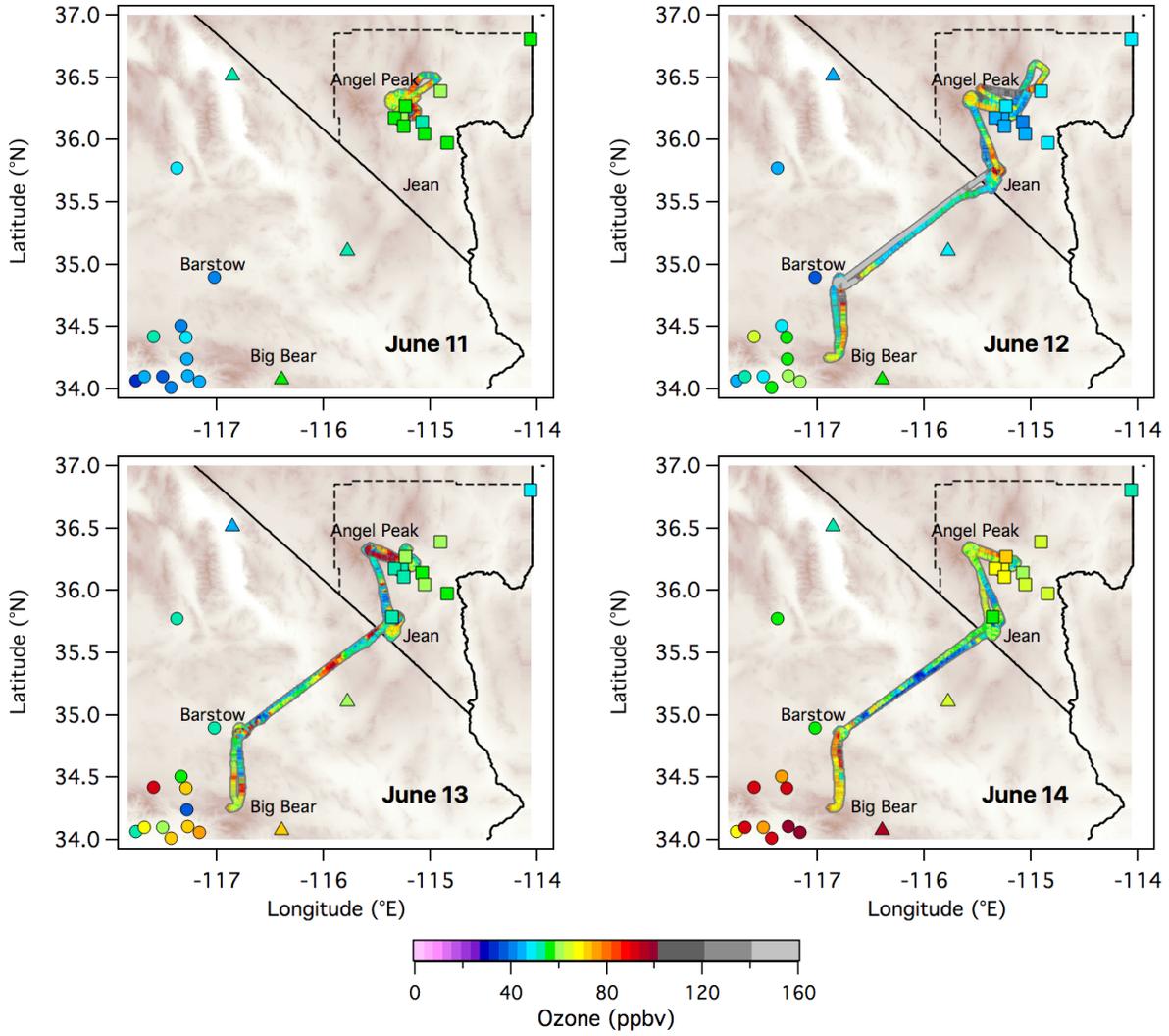
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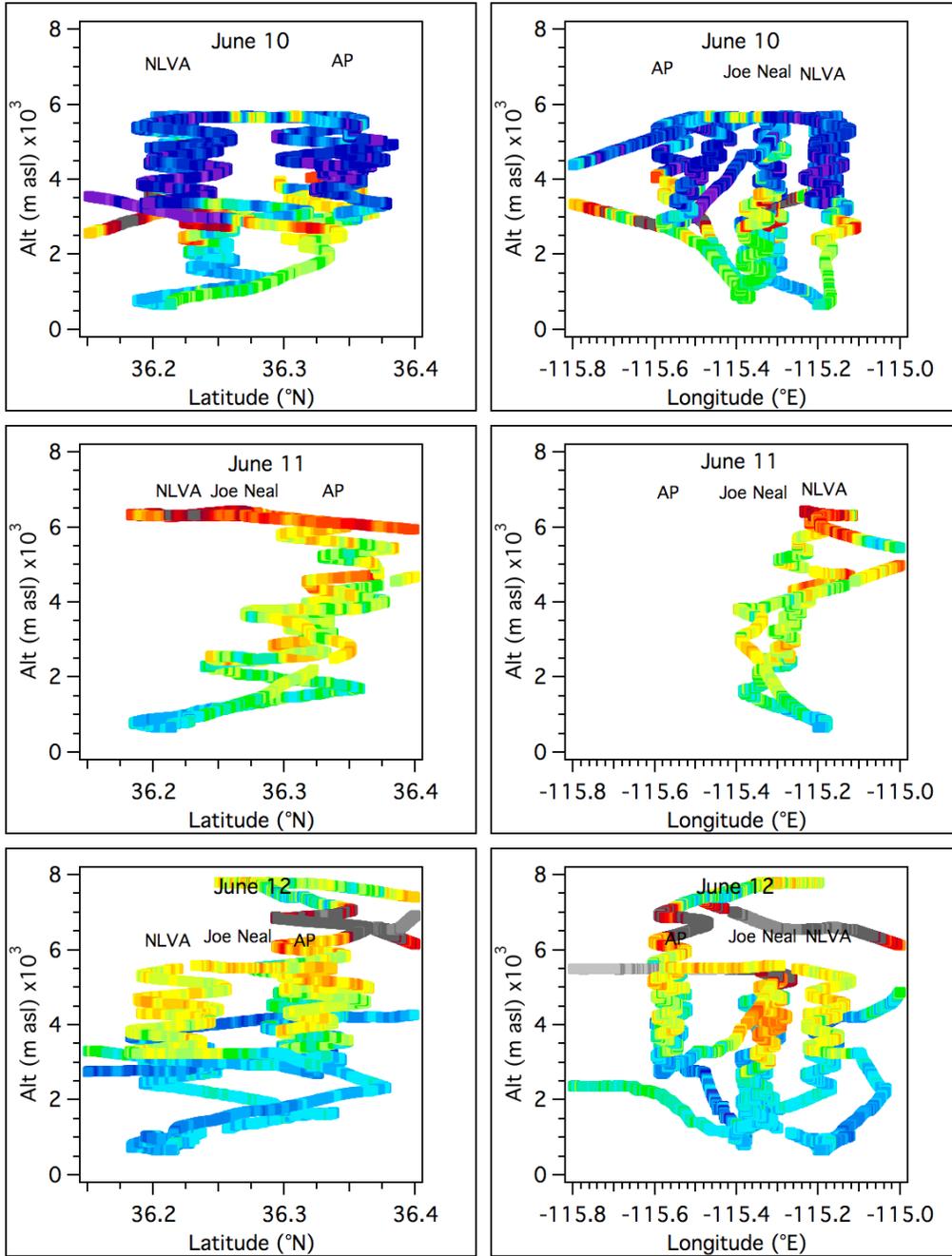
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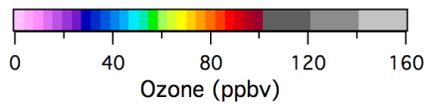
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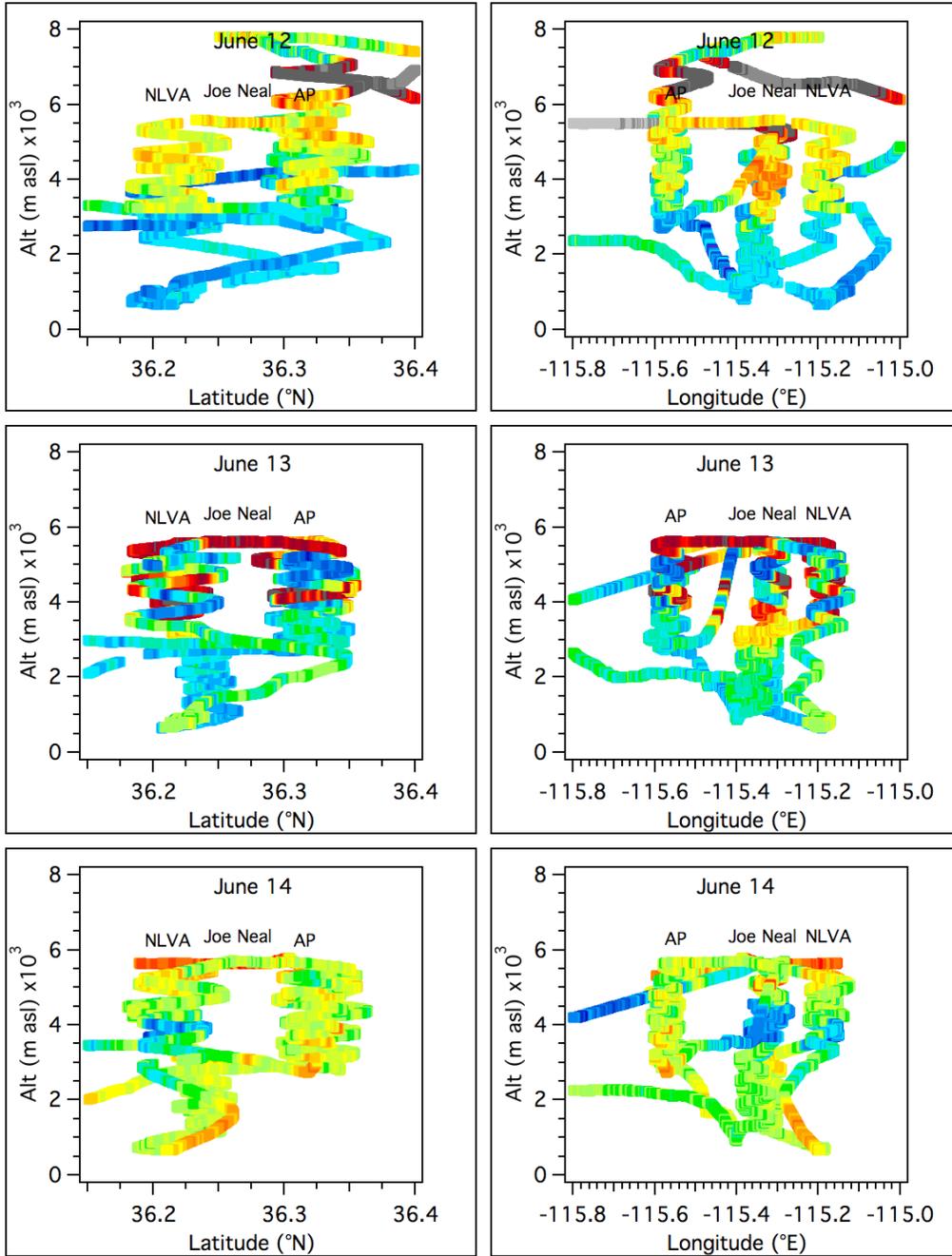


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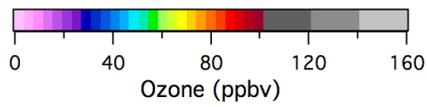
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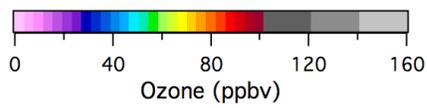
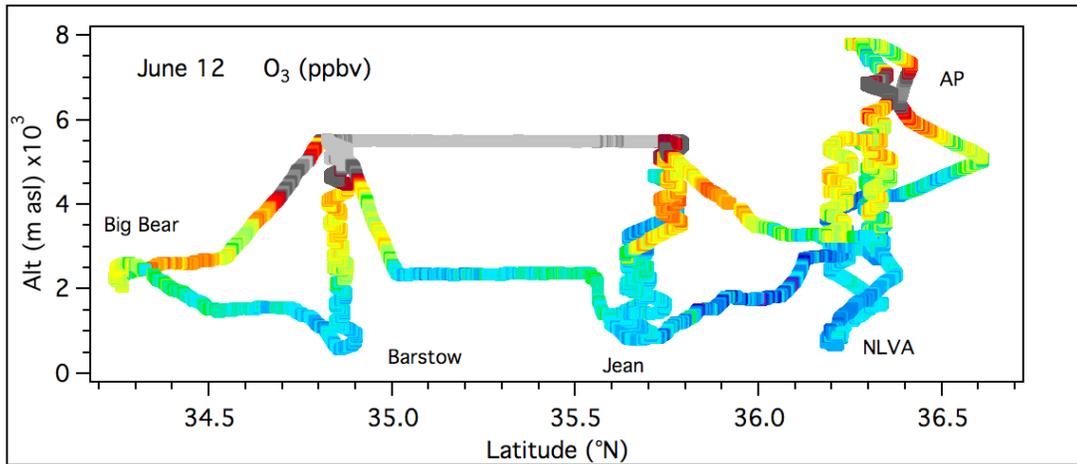
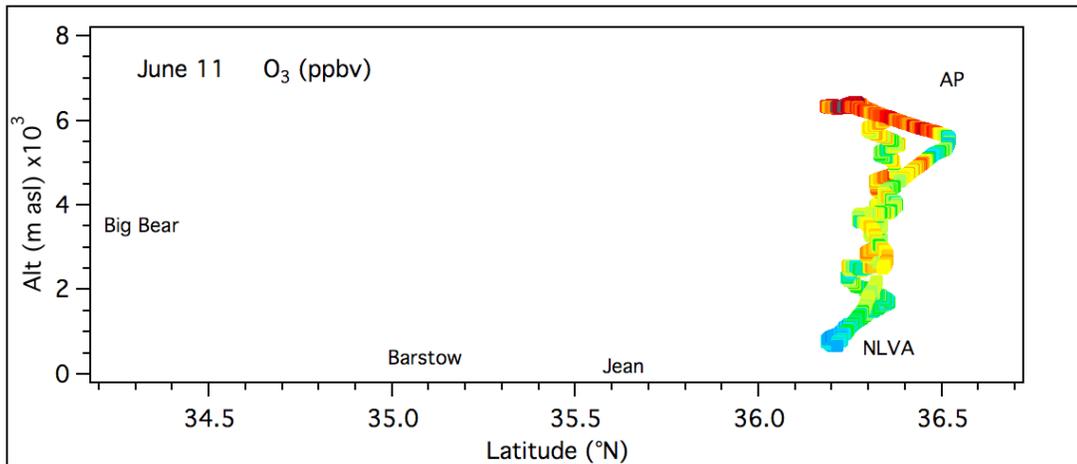
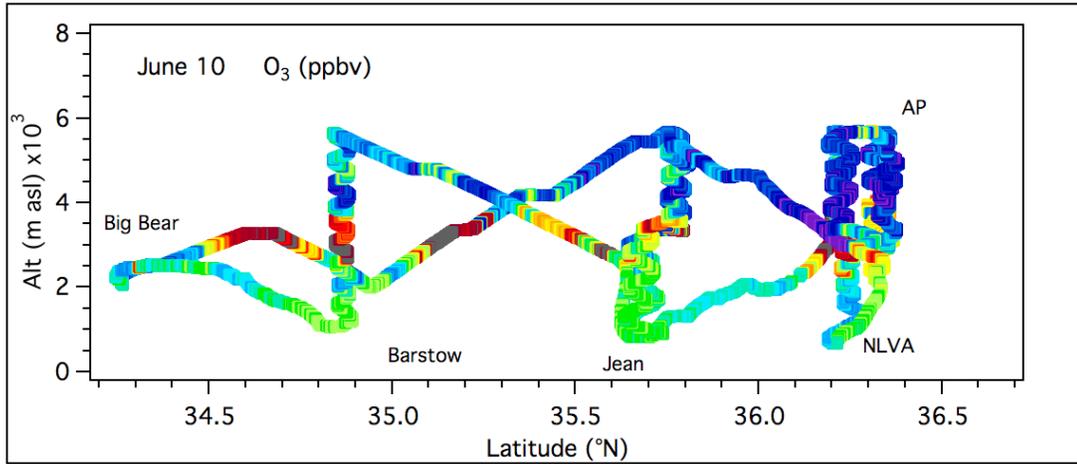


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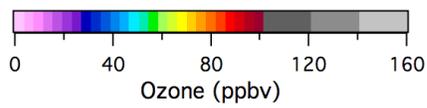
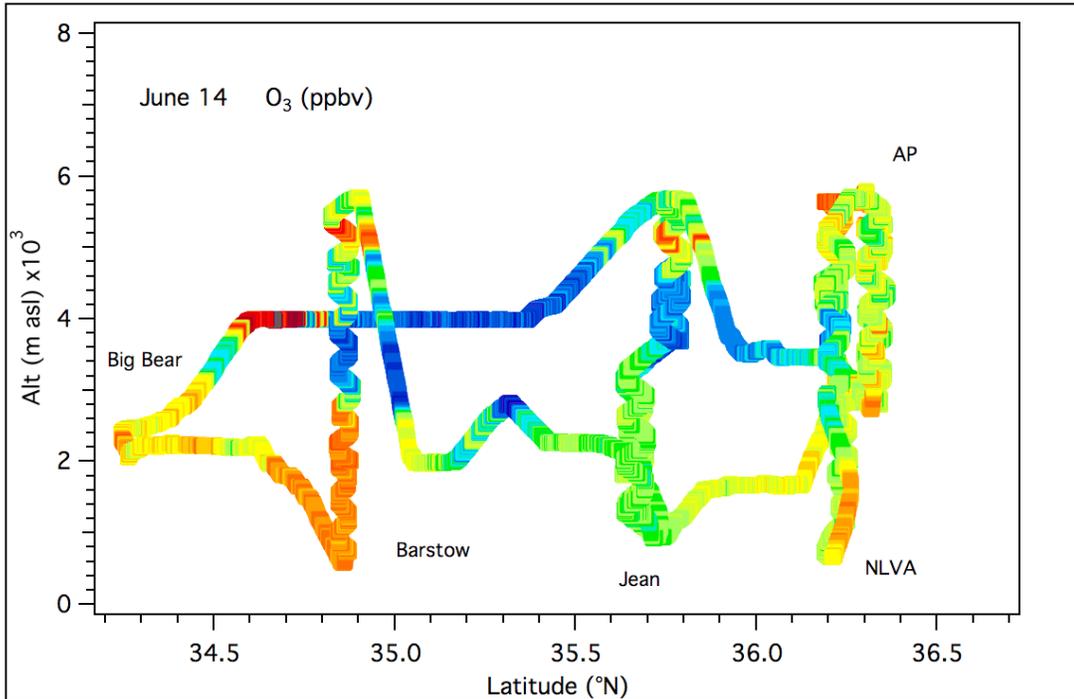
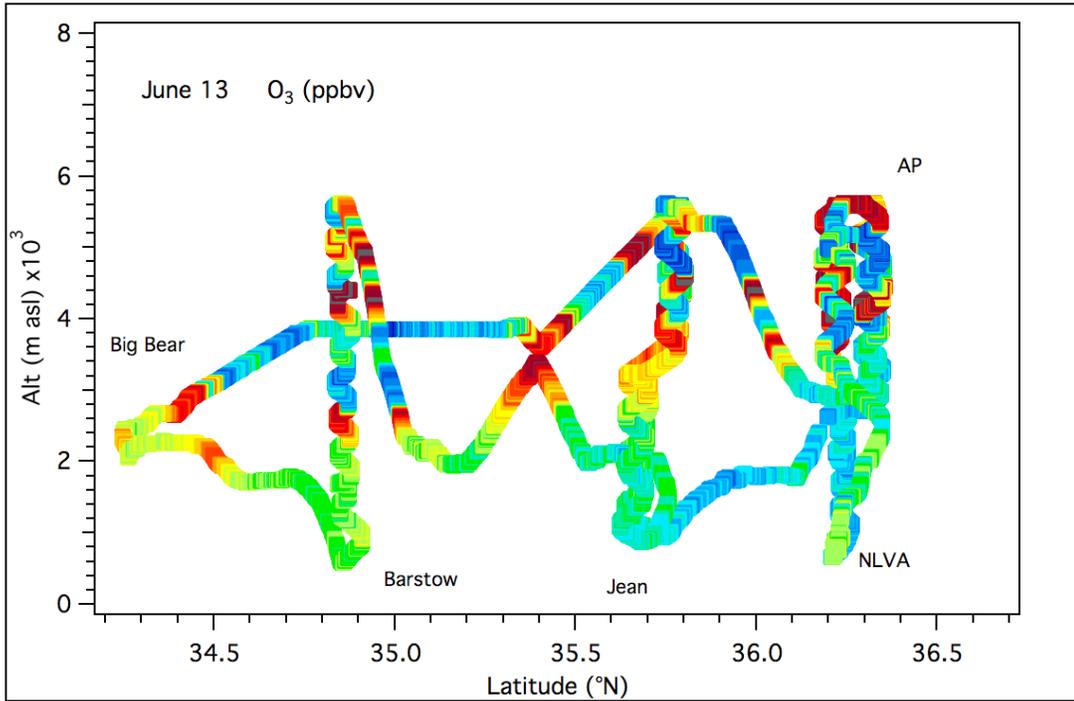
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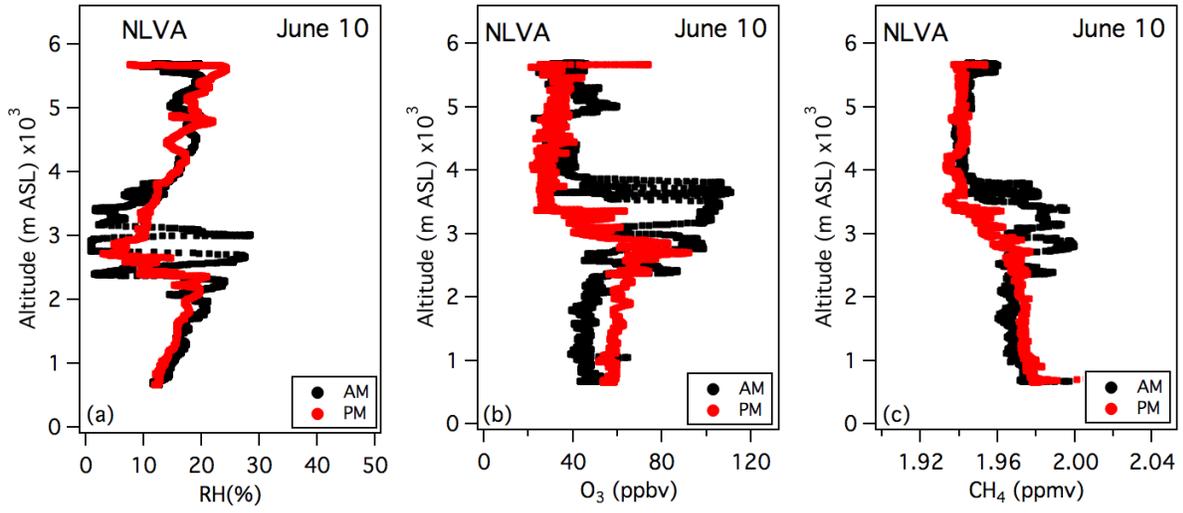
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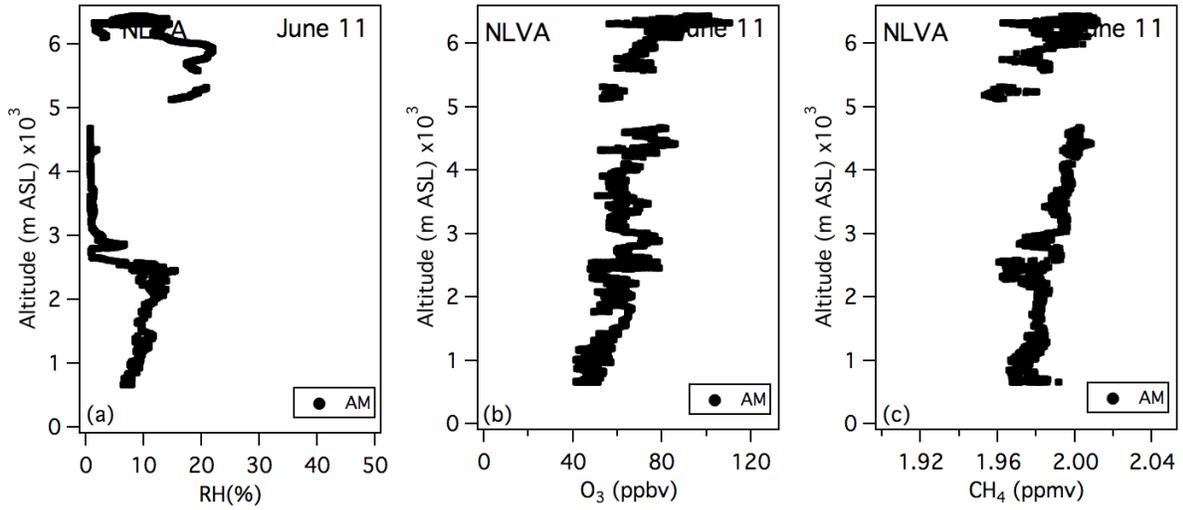


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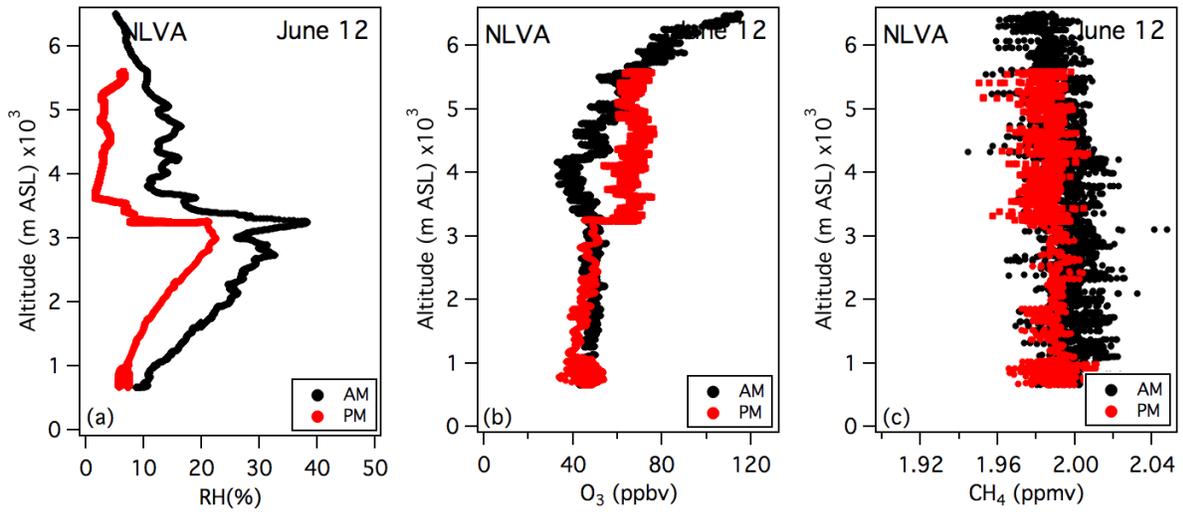
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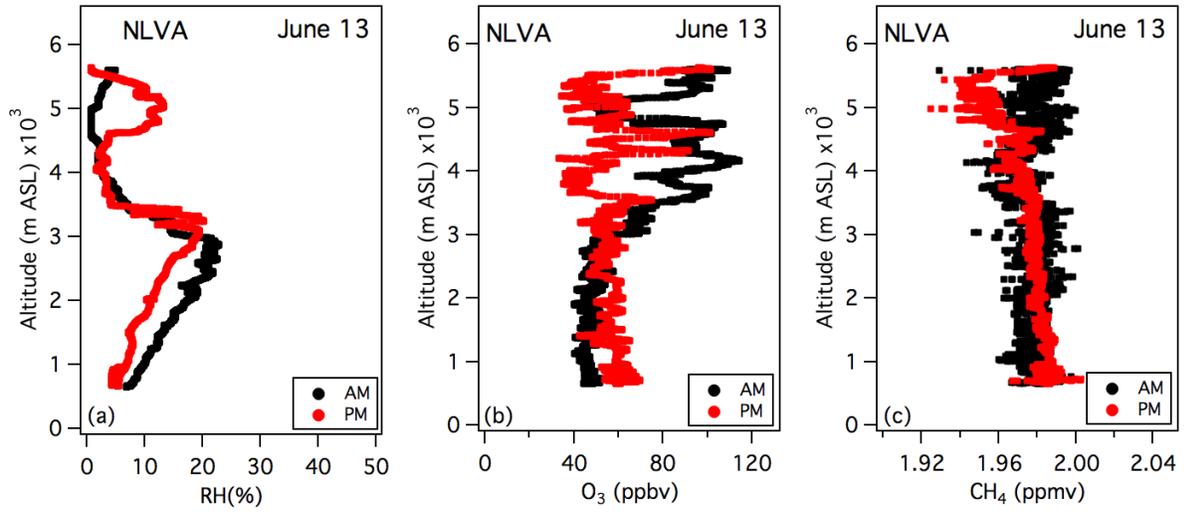
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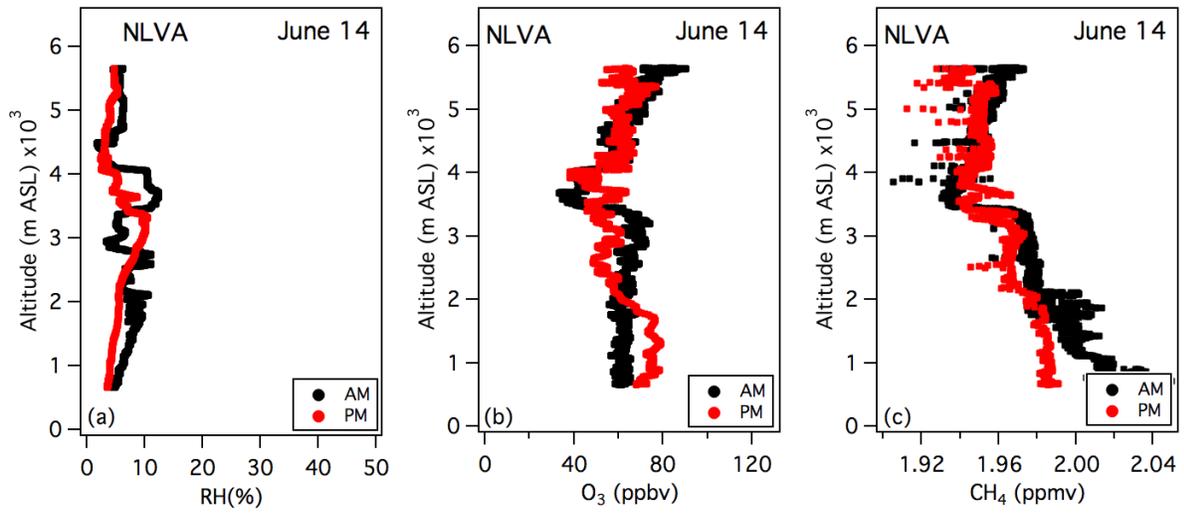
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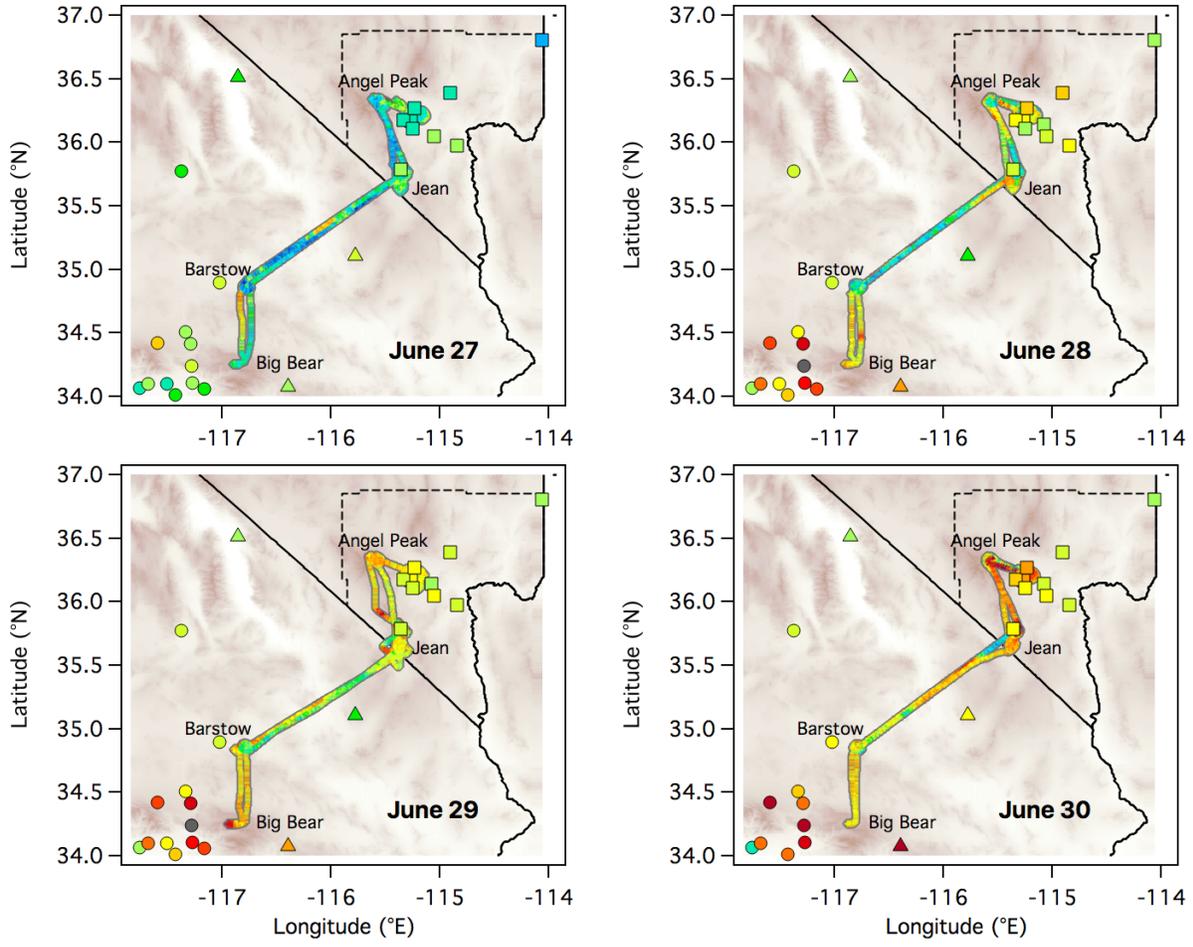
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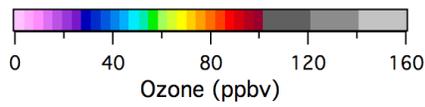
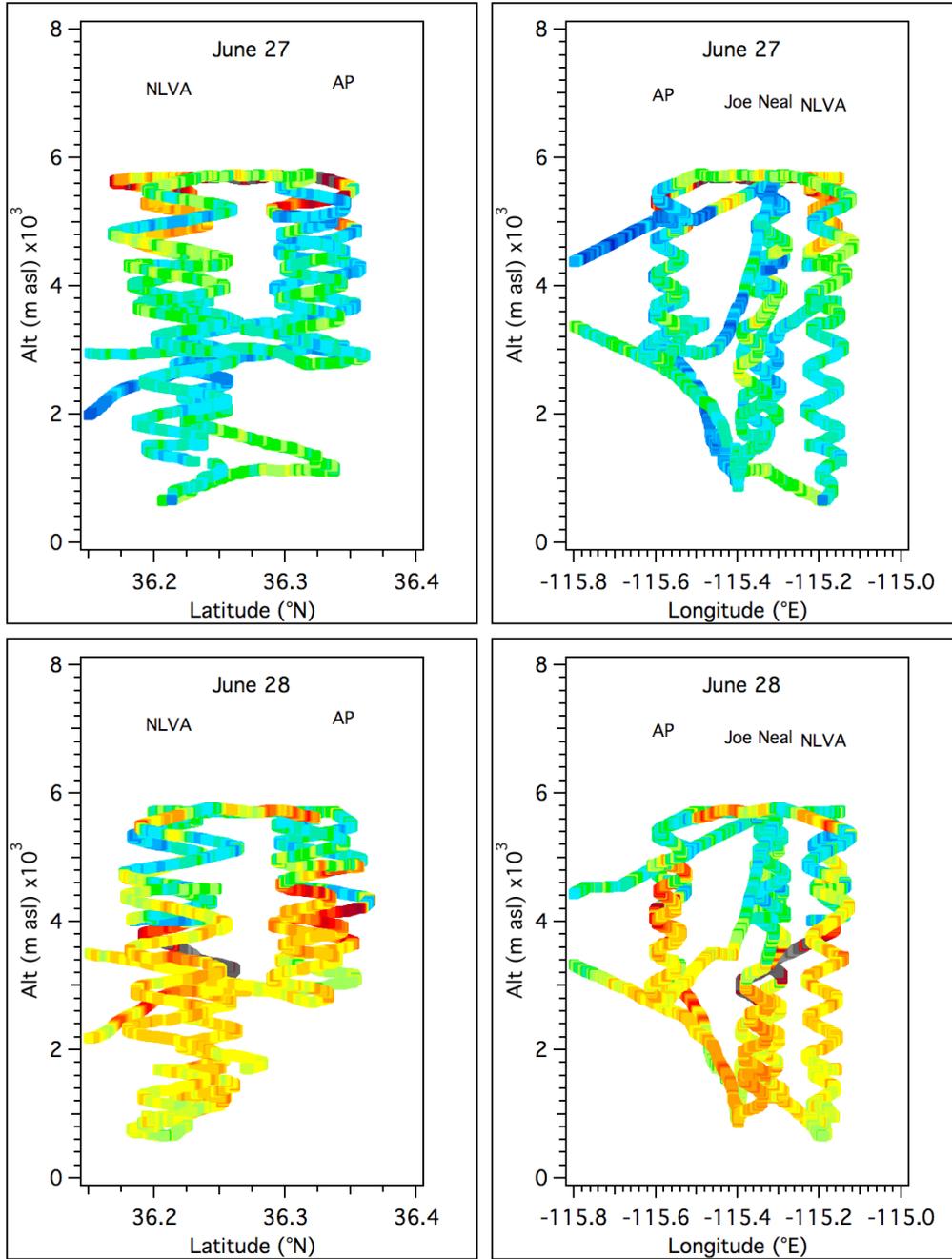
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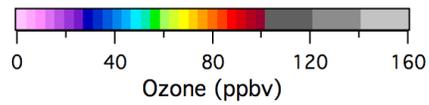
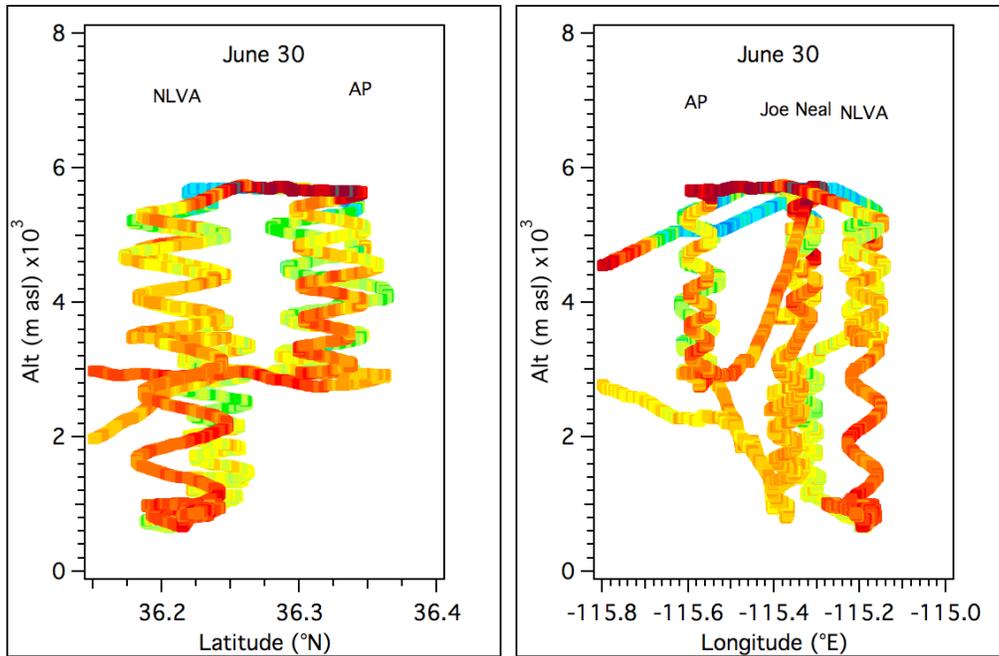
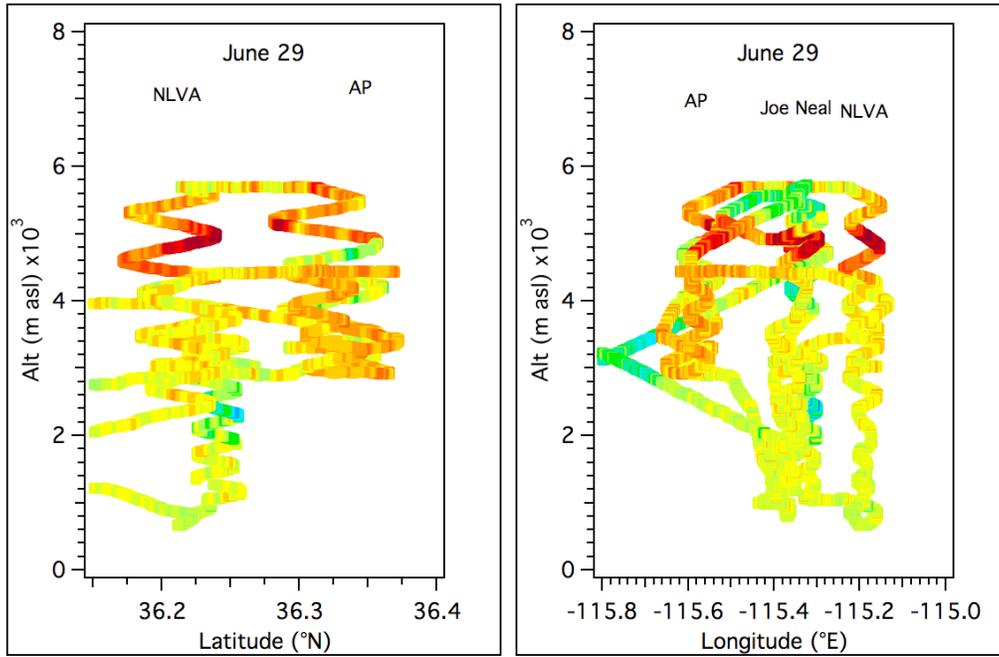
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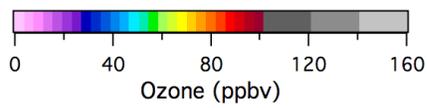
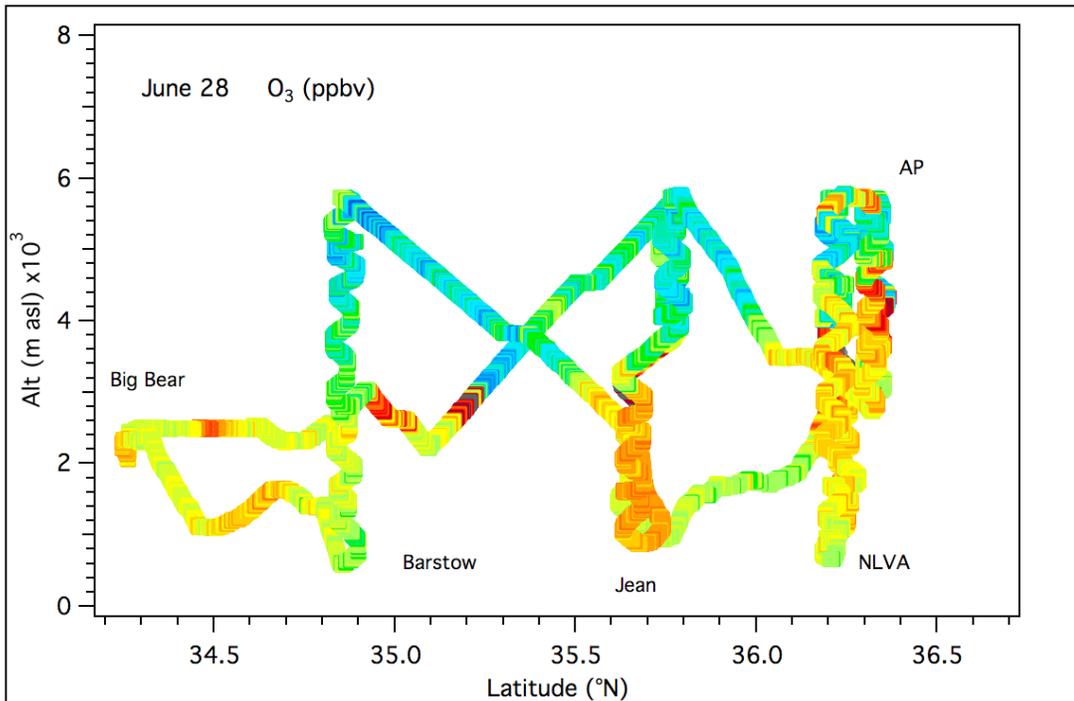
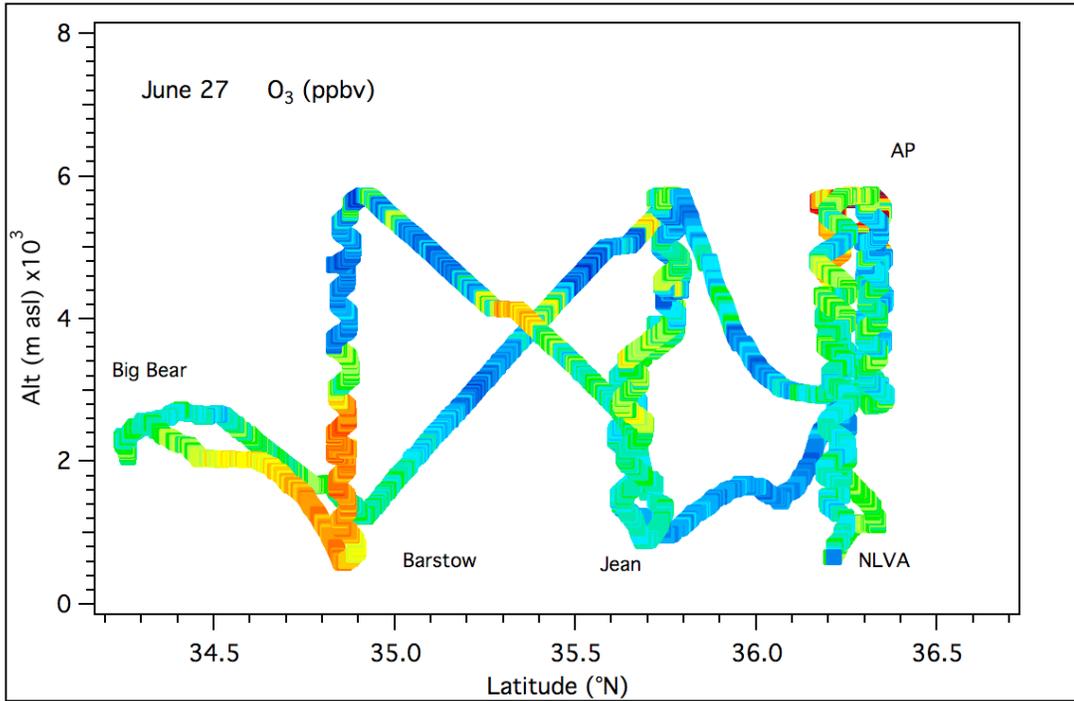
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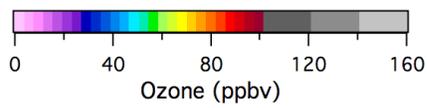
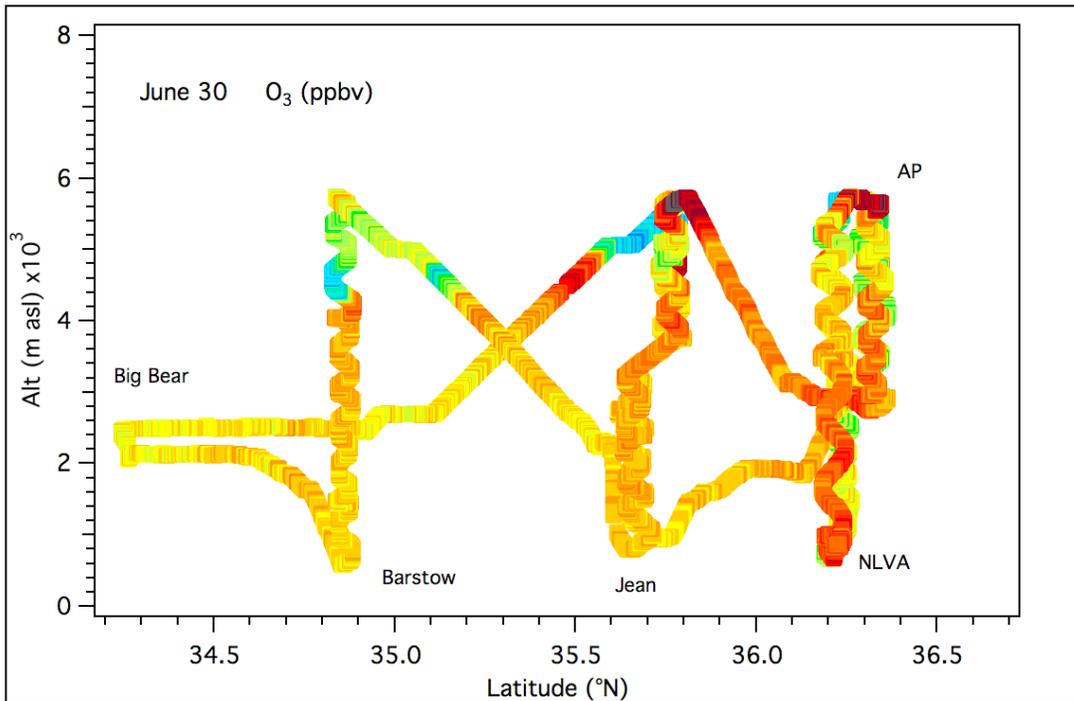
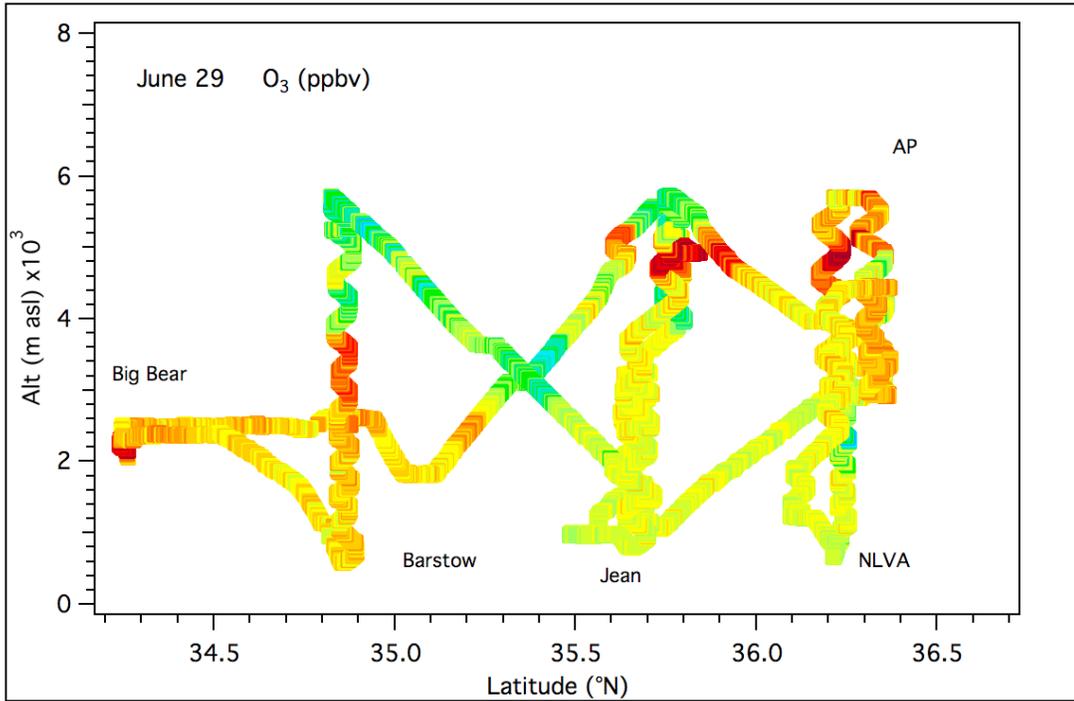
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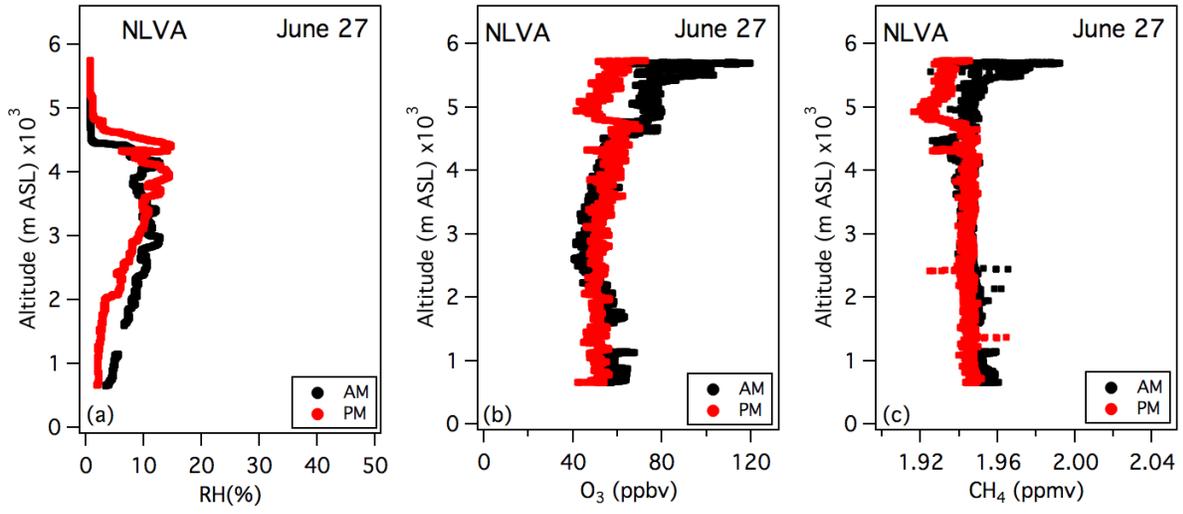
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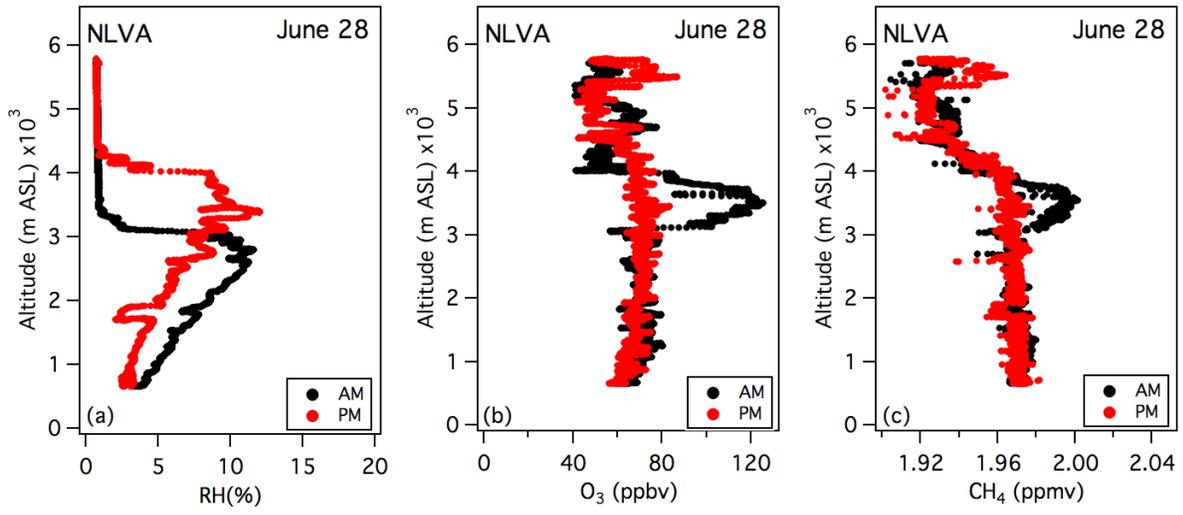


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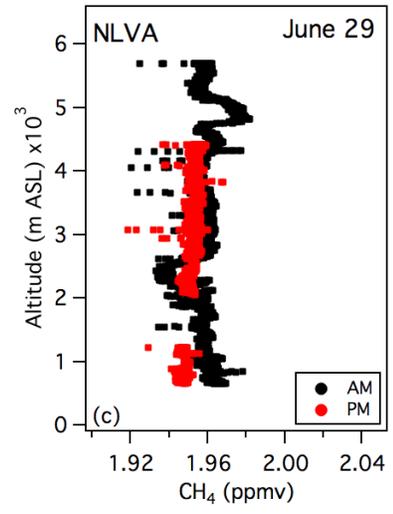
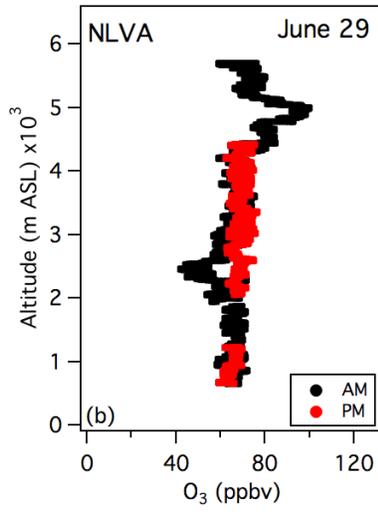
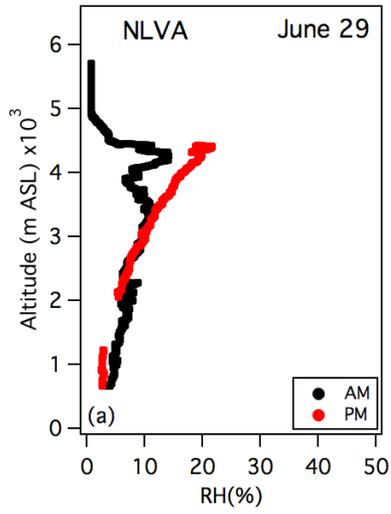
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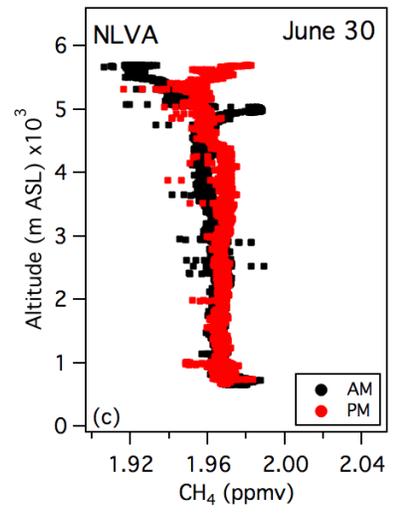
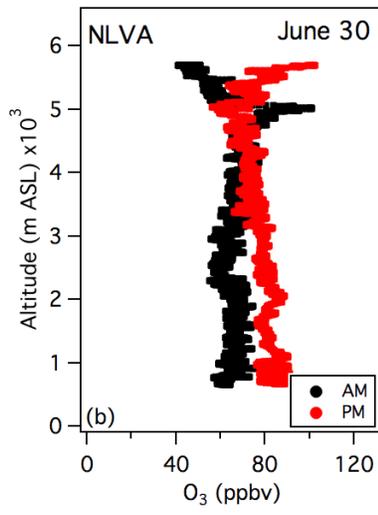
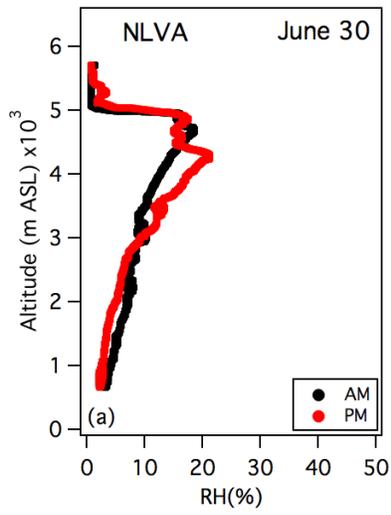
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